

ABSTRACT

Title of Dissertation: EXPLORING REGIONAL EMISSIONS AND TROPOSPHERIC OZONE IN THE EASTERN UNITED STATES USING AIR QUALITY MODELS AND DATA PRODUCTS

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Tropospheric ozone (O_3) is a harmful pollutant regulated by the US Environmental Protection Agency (EPA) through the use of designated air quality standards. Within the United States, approximately 110 million people live within counties designated as in non-attainment of the O_3 standard. In this work, analysis is performed to examine the influence of anthropogenic emissions on tropospheric O_3 production within the framework of the CMAQ regulatory air quality model. Adjustments are recommended to improve emission representation from the largest (class 3) commercial marine vessels (c3 Marine). Model results with the implemented corrections show improved comparison to surface O_3 observations from AQS sites. Characterization of the photochemical O_3 production regime (VOC or NO_x sensitive)

is performed using the ratio of formaldehyde (HCHO) and nitrogen dioxide (NO₂) tropospheric column observations from the satellite borne Ozone Monitoring Instrument (OMI), and whole air sampling canisters in the Long Island Sound (LIS) collected on May 17th and 18th, 2017. Evidence for the importance of anthropogenic VOCs in the New York City pollution plume and their role in tropospheric O₃ production is presented. Aircraft O₃ observations are used to evaluate model performance of the National Oceanic Atmospheric Administration (NOAA) National Air Quality Forecast Capability system CMAQ model for the O₃ event in the LIS. Finally, a series of CMAQ simulations are performed to suggest the likely inventory sector (non-road mobile) most responsible for the significant O₃ production downwind of coastal urban centers like New York and Chicago. Important air quality policy implications of these findings are discussed.

EXPLORING THE IMPACT OF REGIONAL EMISSIONS SOURCES ON
TROPOSPHERIC OZONE IN THE EASTERN UNITED STATES USING AIR
QUALITY MODELS AND DATA PRODUCTS

by

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Dedication

This thesis is dedicated to Gary and Carolyn Ring, my loving and supportive parents. All my life they stimulated my curiosity about the world, fostered my inquisitive mind, and provided me the tools to succeed. Without them, I would not have found the courage within myself to pursue this degree. I am forever grateful for their unyielding love and encouragement. I love you.

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List of Abbreviations

AIS	Automated Identification System
AM8O ₃	Average Maximum 8-hour Ozone
AMF	Air Mass Factor
AQS	Air Quality System
BEIS	Biogenic Emissions Inventory System
BW	Baltimore-Washington
c3 Marine	Class 3 Commercial Marine Vessel
CAA	Clean Air Act
CAMx	Comprehensive Air Quality Model with Extensions
CB05	Carbon Bond mechanism (2005)
CB6	Carbon Bond mechanism (version 6)
CCTM	CMAQ Chemical Transport Model
CEMS	Continuous Emission Monitoring System
CMAQ	Community Multiscale Air Quality model
CMV	Commercial Marine Vessel
CONUS	Continental United States
CRF	Cloud Radiance Fraction
CTM	Chemical Transport Model
DECSO	Daily Emissions estimates Constrained from Satellite Observations model
DISCOVER-AQ	Deriving Information on Surface Conditions from COLUMN and VERTically resolved observations relevant to Air Quality
DOAS	Differential Optical Absorption Spectroscopy
DOMINO	Dutch Ozone Monitoring Instrument NO ₂ retrieval
DV	Design Value
ECA	Emission Control Area
EEZ	Exclusive Economic Zone
EGU	Electrical Generating Unit
EPA	Environmental Protection Agency
ESA	European Space Agency
GEOS-Chem	Goddard Earth Observing System – Chemical transport model
GMI	Global Modeling Initiative
GSFC	Goddard Space Flight Center
HAP	Hazardous Air Pollutant
HRRR	High Resolution Rapid Refresh model
HYSPLIT	HYbrid Single-Particle Lagrangian Integrated Trajectory model
IMO	International Maritime Organization
IUPAC	International Union of Pure and Applied Chemistry
JJA	June, July, August
LIS	Long Island Sound, NY

MARAMA	Mid-Atlantic Regional Air Management Association
MARPOL	International Convention for the Prevention of Pollution from Ships
MCIP	Meteorology-Chemistry Interface Processor
MDA8	Maximum Daily Average 8-hour Ozone
MDE	Maryland Department of the Environment
MEGAN	Model of Emissions of Gases and Aerosols from Nature
MIR	Maximum Incremental Reactivity
MODIS	Moderate Resolution Imaging Spectroradiometer
NAAQS	National Ambient Air Quality Standards
NAM	North American Mesoscale model
NAQFC	National Air Quality Forecast Capability
NASA	National Aeronautics and Space Administration
NEI	National Emissions Inventory
NOAA	National Oceanic and Atmospheric Administration
NYC	New York City
PARANOX	PARAMetrization of emitted NOX model
PBL	Planetary Boundary Layer
PM	Particulate Matter
POCP	Photochemical Ozone Creation Potential
OAQPS	Office of Air Quality Planning and Standards
OTC	Ozone Transport Commission
RAMMPP	Regional Atmospheric Measurement Modeling and Prediction Program
RF	Research Flight
RRF	Relative Response Factor
SAO	Harvard-Smithsonian Astrophysical Observatory
SAPRC	Statewide Air Pollution Research Center
SCD	Slant Column Density
SCIAMACHY	Scanning Imaging Absorption spectrometer for Atmospheric CHartographY
SIP	State Implementation Plan
SPC	Storm Prediction Center
SLT	State, Local, Tribal agencies
SMOKE	Sparse Matrix Operator Kernel Emissions model
STS	Stratosphere-Troposphere Separation
TROPOMI	TROPospheric Ozone Monitoring Instrument
UN	United Nations
UV	Ultraviolet radiation
VCD	Vertical Column Density
VOC	Volatile Organic Compound
WRF	Weather Research Forecasting model

Chapter 1: Introduction

1.1: Ozone in the Troposphere

In the atmosphere, ozone (O_3) is required to create a habitable environment for life on Earth [Finlayson-Pitts and Pitts, 2000; Seinfeld and Pandis, 2006] and yet it is also responsible for causing health problems including premature death [Bell *et al.*, 2006; Cohen *et al.*, 2017; Guarnieri and Balmes, 2014; Kampa and Castanas, 2007]. Stratospheric O_3 , known as “good ozone,” acts as a shield by protecting life on Earth from harmful ultraviolet radiation, making the existence of life possible [Finlayson-Pitts and Pitts, 2000; Seinfeld and Pandis, 2006]. Tropospheric O_3 , known as “bad ozone,” is produced in the atmosphere from precursor species emitted as byproducts of anthropogenic and biogenic activity.

In the early 1950s, Arie Haagen-Smit from the California Institute of Technology discovered that the pollution plaguing Los Angeles, CA (LA) at the time, contained O_3 [Jacobson, 2002; Seinfeld and Pandis, 2006]. He exposed plants and other materials to O_3 in a series of chamber experiments and found they all experienced the same detrimental effects as when exposed to the LA pollution. Further experiments confirmed O_3 was produced from a combination of nitrogen oxides, reactive organic gases, and sunlight, commonly found in the summertime troposphere [Haagen-Smit *et al.*, 1953]. Based on his research, LA and other regions in the United States (US) began imposing air regulations on local businesses in an attempt to alleviate pollution episodes. By the mid-1960s, congress started passing air quality control legislation, however it lacked enforcement capability [Jacobson, 2002]. In 1970, Congress passed

the Clean Air Act (CAA) and established the Environmental Protection Agency (EPA) as the enforcing entity [EPA, 1970; 2017c; Jacobson, 2002; OAQPS, 2007]. Major revisions to the CAA were passed in 1977 and 1990. The amendments of 1977 focused primarily on regions that were not achieving national air quality objectives and provided further guidance for the prevention of significant deterioration of air quality [EPA, 2017c]. The amendments of 1990 were passed to provide expanded control of acid rain, urban air, and toxic air pollution [EPA, 2017c]. More discussion about environmental regulation in the United States (US) will follow in Section 1.1.2.

It is important to remember that air pollution is a global problem. Exposure to ambient concentrations of O₃ and particulate matter (PM) served as the leading source of global disease in 2015, disproportionally impacting underdeveloped countries [Cohen *et al.*, 2017; Lelieveld *et al.*, 2015]. Chronic exposure to O₃ causes impairment to lung function, eye irritation, and is linked to increased risk of mortality [Bell *et al.*, 2006; Kampa and Castanas, 2007; Lippmann, 1989]. As a secondary pollutant, O₃ is formed from precursors emitted into the atmosphere like carbon monoxide (CO), nitrogen oxides (NO_x = NO + NO₂), and volatile organic compounds (VOCs). The mechanisms describing O₃ production from these precursor species are detailed in Sections 1.1.1 and 1.1.2. Understanding the complex mechanisms and the emission sources of the precursors involved in producing tropospheric O₃ is vital to developing strategies to mitigate O₃ formation.

1.1.1: Ozone Production from CO and NO_x

Emissions of CO and NO_x are predominantly from combustion sources associated with anthropogenic activity and biomass burning [EPA, 2017b; e].

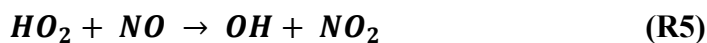
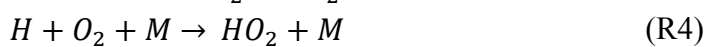
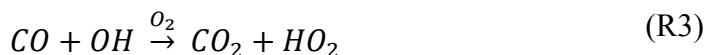
Emissions of CO result from the incomplete combustion of hydrocarbons from industrial processes, vehicle emissions, wood and fuel burning, and wildfires. Vehicle emissions alone can contribute up to 95% of the total CO measured in cities [EPA, 2017b]. Emissions of NO_x are from fossil fuel combustion, power generation, industrial sources, soil microbes, and lightning [EPA, 2017e].

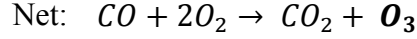
Earth has an oxidizing atmosphere which means it is rich in oxygen (O₂), but the primary oxidant in the troposphere, and therefore the most important species for tropospheric chemistry, is the hydroxyl radical (OH) [Levy, 1971]. Formation of the OH radical begins with the reaction sequence below [Seinfeld and Pandis, 2006]:



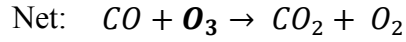
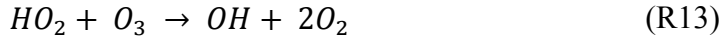
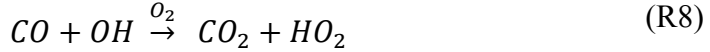
Photolysis of O₃ that leads to formation of the excited state oxygen atom O(¹D) occurs at wavelengths (λ) < 320 nm (reaction R1). This O(¹D) is critical for the formation of OH via reaction R2. Atmospheric concentrations of OH are sustained at $\sim 10^6$ molecules/cm³, maintained via reaction with OH and other compounds [Perner *et al.*, 1987].

While some O₃ is transported from the stratosphere into the troposphere, most is produced directly in the troposphere [Fishman and Crutzen, 1978; Jacob, 1999; Seinfeld and Pandis, 2006]. The mechanism for tropospheric O₃ production involving CO and NO_x is described below:





If NO reacts with O₃ rather than HO₂ (R10), the net reaction would consume tropospheric O₃:



Therefore, the process by which NO oxidizes is critical for tropospheric O₃ production. In this mechanism, the important rate-limiting reaction (R5) determines the production of tropospheric O₃. It is represented as the equation below:

$$P_{O_3} = k[HO_2][NO] \quad (14)$$

Concerning reactions R6 and R11, NO₂ will photodissociate at $\lambda < 420$ nm; high energy radiation with some wavelengths just outside the ultraviolet (UV) range. *Dickerson et al.* [1997] used the Variable-grid Urban Airshed Model (UAM-V) model to show that the rate of NO₂ photolysis was amplified by the aerosol scattering of solar radiation; particularly impactful in the planetary boundary layer. This process accelerates the rate of tropospheric photochemical production of O₃ in polluted regions.

1.1.2: Ozone Production from VOCs and NO_x

Tropospheric O₃ is also produced through atmospheric reactions of Volatile Organic Compounds (VOCs). These organics oxidize with OH, beginning the reaction sequence that produces O₃. VOCs have biogenic sources (mainly trees) and anthropogenic sources such as industrial activity, paints and solvents, evaporative emissions from gasoline, and engine exhaust [EPA, 2017f]. Figure 1 shows the relative contribution of anthropogenic and biogenic VOC emissions in 2011 for the US (a), and the relative contribution of four source sectors to the total anthropogenic VOC emissions in 2011 (b).

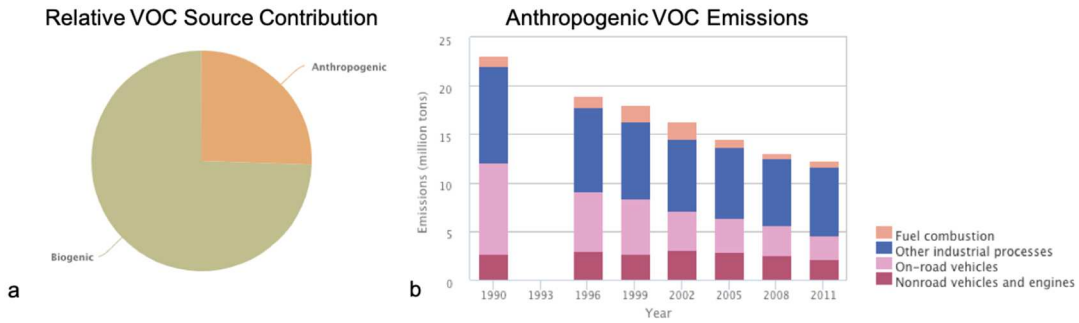
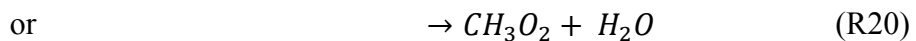
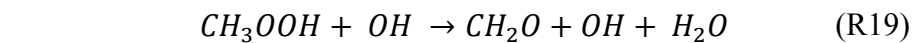
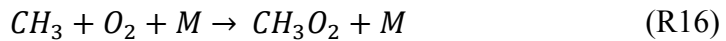
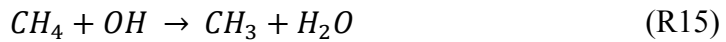
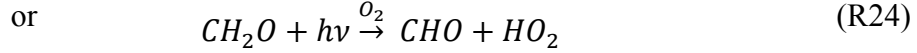
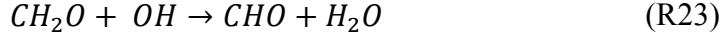
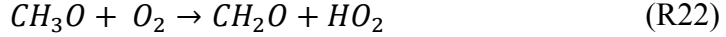
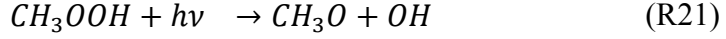


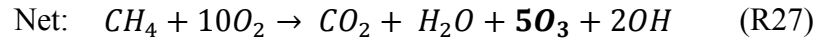
Figure 1: VOCs in the US: (a) relative source contribution from 2011 of anthropogenic and biogenic VOCs in the US (b) VOC emissions from four source sectors between 1990 and 2011. The on-road sector consists of cars, trucks, buses, and motorcycles. For this plot, the non-road sector consists of farming equipment, lawn equipment, ships, aircraft, etc. This figure is from EPA [2017f].

A representative mechanism for VOC oxidation is provided below. The example shown is for methane (CH₄) [Jacob, 1999]:

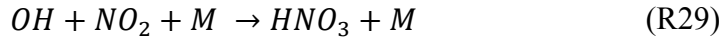
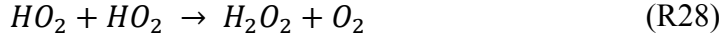




Begin CO oxidation described in Section 1.1.1



Based on the meaningful net reaction (R27), in the presence of NO_x, five O₃ molecules are produced during the oxidation process of CH₄ to CO₂. Termination of these production mechanisms requires radical reaction to produce a nonradical:



In relatively clean environments, or regions with low NO_x concentrations, radical termination occurs through R28. In regions of high NO_x, radical termination occurs through R29.

1.1.3: The NO_x-O₃-VOC Relationship

Tropospheric O₃ production relies upon the non-linear relationship between NO_x and VOCs [Jacob, 1999; Jacobson, 2002; Lippmann, 1989; Sillman, 1990]. Effective emission control strategies are developed according to the local photochemical O₃ production regime: either NO_x-sensitive or VOC-sensitive. The NO_x-sensitive regime refers to atmospheric conditions where concentrations of NO_x are relatively low compared to VOCs. In this regime, a reduction of NO_x leads to a reduction of O₃. The VOC-sensitive regime describes atmospheric conditions where

concentrations of VOCs are relatively low compared to NO_x . In this regime, a reduction of NO_x leads to an *increase* of O_3 .

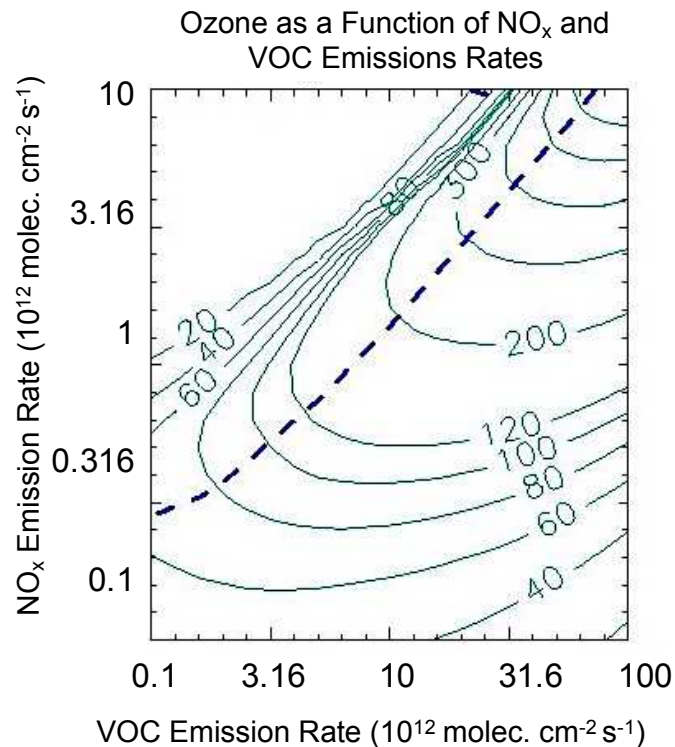


Figure 2: Isopleths of O_3 (ppb) as a function of NO_x and VOC emissions simulated by a photochemical box model. The blue dotted line represents NO_x and VOC emission rates that correspond to a ridge of peak O_3 production. NO_x -sensitive conditions are represented below the ridge line, VOC-sensitive conditions are represented above. (Figure from <http://www-personal.umich.edu/~sillman/ozone.htm#OZO1.2>)

Figure 2 shows isopleths of O_3 calculated by a photochemical box model as a function of VOC and NO_x emission rates. The blue dotted line represents the ridge of maximum O_3 production. Regions below the ridge line are considered NO_x -sensitive. Regions above the ridge line are considered VOC-sensitive. In this regime, reductions in NO_x alone could have the *local* adverse effect of increasing O_3 . It is never advisable to increase NO_x emissions, therefore controlling VOCs and NO_x emissions together are necessary to achieve the desired O_3 reduction.

Development of effective air quality policy depends heavily on the regulated region's $[\text{VOC}]/[\text{NO}_x]$ ratio. Creation of controls targeting the appropriate pollutants (NO_x and/or VOC) is necessary to effectively reduce surface O_3 . Determination of the photochemical production regime, however, can be challenging. A variety of factors including meteorology, biogenic emissions, temporal and spatial variation of local anthropogenic activity, and transported pollution from upwind sources all contribute to the $[\text{VOC}]/[\text{NO}_x]$ ratio, and therefore the instantaneous rate of O_3 production for a given region. In an effort to develop effective air quality strategies, state agencies often rely on air quality modeling simulations. Therefore, considerable effort is made to create model frameworks that represent the actual atmosphere.

1.2: Air Quality Regulation and Health

Consequential amendments to the original CAA passed in 1970 were passed in 1977 and 1990, which gave the EPA more authority to enforce air pollution regulation [EPA, 1970; 1990; 2017c; OAQPS, 2007]. Passage of the CAA established the National Ambient Air Quality Standards (NAAQS), which sets atmospheric abundance limits for six criteria pollutants [EPA, 1970; 1990; 2015c; 2017c; 2018d]. Current standards are described in Table 1.

The criteria pollutants were chosen based on their known health risks and environmental degradation capability [EPA, 1990; 2018d]. Each pollutant has a primary and secondary standard. The primary standard is set to protect the US population from health risks, especially for sensitive populations like the elderly, children, and asthmatics. The secondary standard is set to protect the environment

(ecosystems, crops, visibility, climate, etc.) from adverse effects of the pollutant [EPA, 1990; 2018d].

Table 1: Current NAAQS standards for criteria pollutants [EPA, 2016b]

Pollutant	Primary or Secondary Designation	Time Frame	Standard	Meeting the Standard	
Particulate Matter	PM _{2.5}	Primary	1 year	12 µg/m ³	Annual mean (3-year average)
	Secondary	1 year	15 µg/m ³	Annual mean (3-year average)	
	Primary and Secondary	1 day	35 µg/m ³	98 th percentile (3-year average)	
	PM ₁₀	Primary and Secondary	1 day	150 µg/m ³	No more than 1 exceedance/year on average for 3 years
Ozone (O ₃)	Primary and Secondary	8 hours	70 ppb	Annual 4 th -highest max. daily avg. 8-hr O ₃ (MDA8) averaged for 3 years	
Sulfur Dioxide (SO ₂)	Primary	1 hour	75 ppb	99 th percentile of daily 1-hour max. conc. average for 3 years	
	Secondary	3 hours	0.5 ppm	No more than 1 exceedance/year	
Nitrogen Dioxide (NO ₂)	Primary	1 hour	100 ppb	99 th percentile of daily 1-hour max. conc. average for 3 years	
	Secondary	1 year	53 ppb	Annual mean	
Carbon Monoxide (CO)	Primary	8 hours	9 ppm	No more than 1 exceedance/year	
		1 hour	35 ppm		
Lead (Pb)	Primary and Secondary	running 3-month average	0.15 µg/m ³	No exceedances	

Air quality has improved since inception of the CAA in 1970, with significant advancements since establishment of the NAAQS in 1990 [EPA, 2018b; He *et al.*, 2013b; Lamsal *et al.*, 2015; Tong *et al.*, 2015]. Northeast regional (from Maryland to Maine) average concentrations of the six criteria pollutants have declined over the past two decades (Figure 3), such that all are now below the national standard, except for O₃ (as of 2017) [EPA, 2018b]. In the Northeast, tropospheric O₃ is produced in the atmosphere from local emissions of VOCs as well as NO_x, and O₃ is also transported into the region from upwind out-of-state sources due to the prevailing westerly winds

[Endlich et al., 1984; He et al., 2016a; Loughner et al., 2011; Moghani et al., 2018; Ryan et al., 1998; Sillman et al., 1993; Wolff et al., 1977].

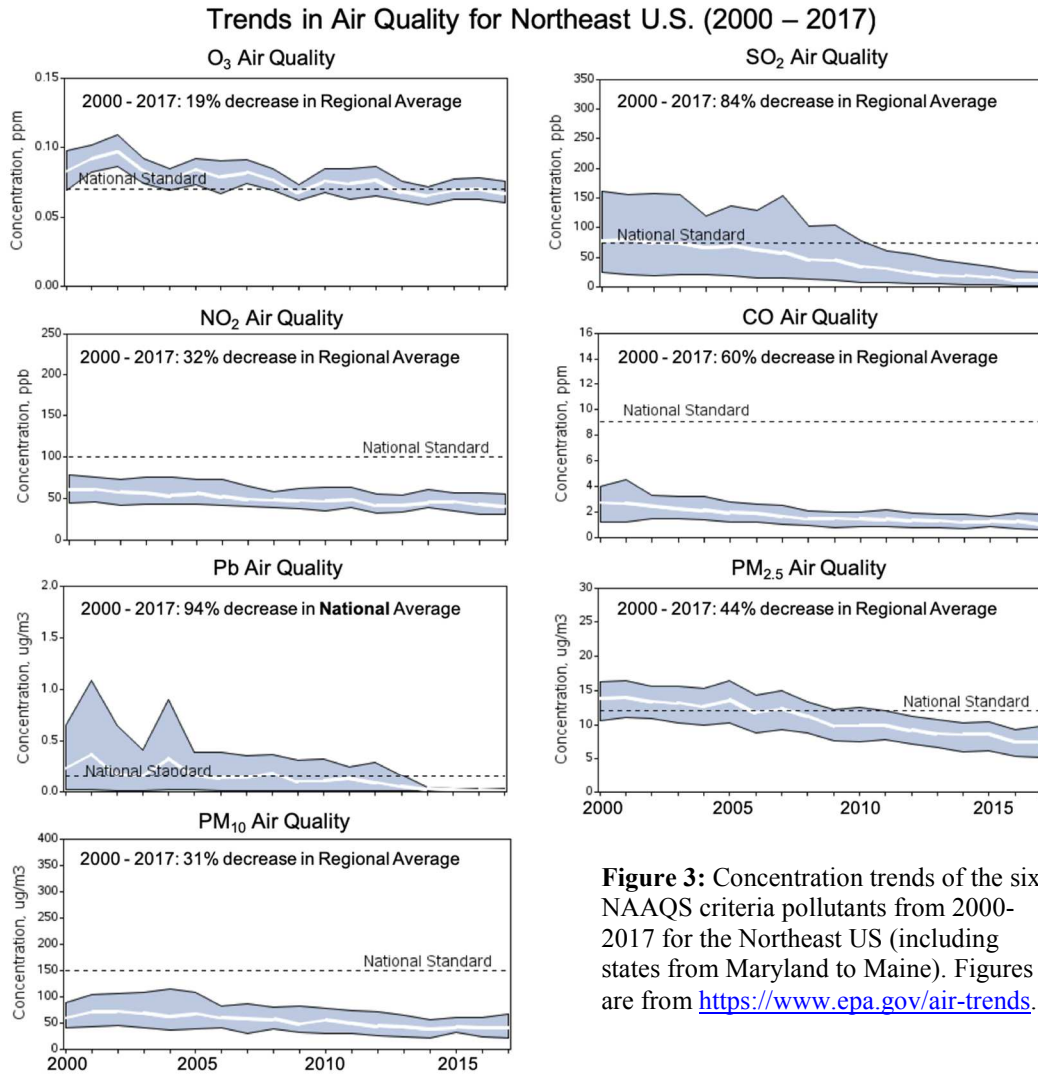


Figure 3: Concentration trends of the six NAAQS criteria pollutants from 2000-2017 for the Northeast US (including states from Maryland to Maine). Figures are from <https://www.epa.gov/air-trends>.

Emissions from large populations coupled with typical synoptic meteorology contribute to the continued challenges in meeting the O₃ NAAQS faced by the Northeast US. Figure 4a shows the national map of 8-hour O₃ non-attainment areas and Figure 4b shows the population living within these non-attainment areas (as of 2017). With about one third of the total US population subjected to unhealthy levels of

pollution, emission control of the criteria pollutants and specifically O₃ precursors are essential to improving air quality. As shown in Figure 4, 8-hour O₃ non-attainment areas are mostly concentrated in or downwind of heavily populated urban areas. In other words, people living in urban centers are subjected to poor air quality and related health problems, more than those in rural areas. Studies have shown that incidences of hospital visits coincide with poor air quality episodes [Bell *et al.*, 2007; Fann *et al.*, 2018; Fauroux *et al.*, 2000]. Symptoms of impaired health related to air quality include episodic respiratory distress, aggravation of conditions like asthma, heart complications, and in some cases death [Bell *et al.*, 2007; Fann *et al.*, 2018; Fann *et al.*, 2015; Fauroux *et al.*, 2000; Pope and Dockery, 2006; Schnell and Prather, 2017; Thompson *et al.*, 2014; Watts *et al.*, 2015]. Additionally, particulate matter (PM_{2.5}) can move into the blood once inhaled into the lungs, causing problems throughout the body [Fann *et al.*, 2018; Pope and Dockery, 2006; Watts *et al.*, 2015].

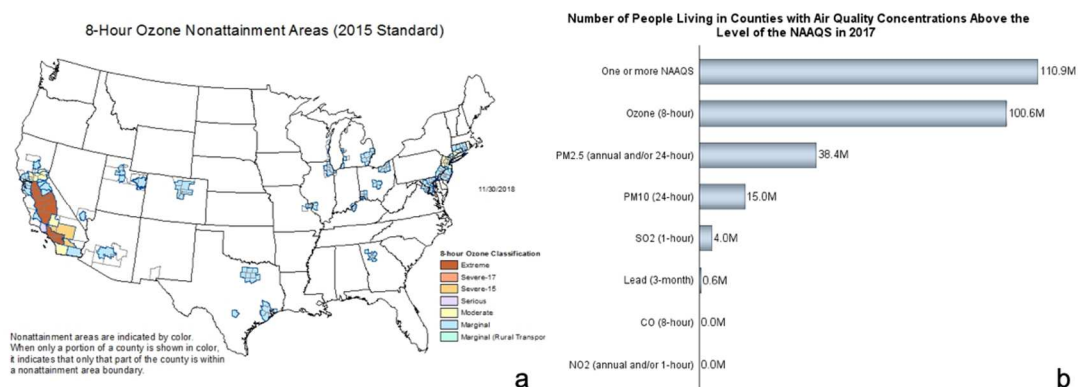


Figure 4: EPA designated (a) non-attainment areas for the 2015 8-hour O₃ standard and (b) the number of people living in counties with concentrations of pollutants above the NAAQS standard, designated as non-attainment. Image (a) is from <https://www.epa.gov/air-trends/air-quality-national-summary>. Image (b) is from <https://www.epa.gov/air-trends/air-quality-national-summary>.

In a warming climate, air quality related health issues are expected to rise [Bell *et al.*, 2007; Fann *et al.*, 2015; Watts *et al.*, 2015]. A study by Bell *et al.* [2007] found

the average number of days per summer that will exceed the 8-hour O₃ standard (75 ppb at the time of this study) will increase 68% when using the Intergovernmental Panel on Climate Change (IPCC) A2 climate scenario. This scenario holds anthropogenic emissions constant while allowing fluctuation for the temperature induced changes to chemical reaction rates and biogenic emissions [Bell *et al.*, 2007]. In summary, in order to mitigate health issues for the growing human population, pollution regulations to limit poor air quality episodes will continue to be necessary.

1.3: Overview of Regulatory Air Quality Modeling

As required by the CAA, state and local governments must develop strategies to ensure air quality monitoring sites within the State meet the various standards [EPA, 1990; 2015c; 2018d; OAQPS, 2007]. These State Implementation Plans (SIPs) outline proposed emission reduction strategies to meet the federal regulations. States with areas that do not meet the air quality standards (non-attainment areas) must submit specific SIPs for each area designated as non-attainment to the EPA for approval [EPA, 2017a]. The SIP should include the control strategies to be implemented and the costs and benefits of enforcing the control measures to attain the national standard. It should also explain all monitoring procedures, methods for addressing any violations, and designate the enforcing agency for the proposed control measures [EPA, 1990; 2018d]. States must submit detailed summaries including:

1. the methods used to determine necessary emissions reductions;
2. all tools used to analyze the proposed controls including data, emissions and air quality models;

3. the expected attainment results to demonstrate adequacy of the proposed SIPs [EPA, 1990; 2018d].

1.3.1: Modeling in the Mid-Atlantic Region

Maryland is a member of the Ozone Transport Commission (OTC), a multi-state organization that assists in the development and implementation of strategies to improve air quality with respect to surface O₃ (<https://otcair.org/about.asp>). The OTC includes states from Virginia to Maine, encouraging collaborative efforts from all the member state environmental departments to resolve issues related to surface O₃.

For the modeling centers in the Eastern US, a baseline model platform for the year 2011 was developed by the EPA to assist states with the creation of SIP quality attainment strategies [EPA, 2014; 2015a]. Creation of a model platform requires generation of anthropogenic and biogenic emissions, model domain initial and boundary conditions, and meteorology representative of the base year. The EPA released the first version of the 2011 model platform in early 2014 called “CAP_BAFM 2011-Based Platform, version 6” (2011v6) [EPA, 2014]. This platform models Criteria Air pollutants and Precursors (CAPs) and Benzene, Acetaldehyde, Formaldehyde (HCHO), and Methanol (BAFM). Known hazardous air pollutants for the model year 2011 version 6 are improved from the previous base year of 2007 or “version 5” [EPA, 2014]. The 2011v6 platform includes three improvements that warranted new version numbers. The most recent release in 2016 is called 2011v6.3 [EPA, 2016c], however most modeling performed in the following analyses uses the 2011v6.2 model platform [EPA, 2016c].

To create the model-ready emissions, the EPA used the Sparse Matrix Operator Kernel Emissions (SMOKE) Modeling System [CMAS, 2014b]. The primary use of SMOKE is to create gridded, speciated, hourly emissions files to be used in air quality modeling. National Emissions Inventories (NEIs) are developed every 3 years and consist of total emissions estimates gathered from the EPA and State, Local, and Tribal (SLT) agencies for a wide variety of chemical species for the CONtinental US (CONUS) [EPA, 2015a]. SMOKE temporalizes and assigns source locations (and heights) for the emissions based on the surrogate location files provided by the EPA (gridded at 12km) [EPA, 2014; 2015a; d]. A schematic describing the major components of the SMOKE model is shown in Figure 5.

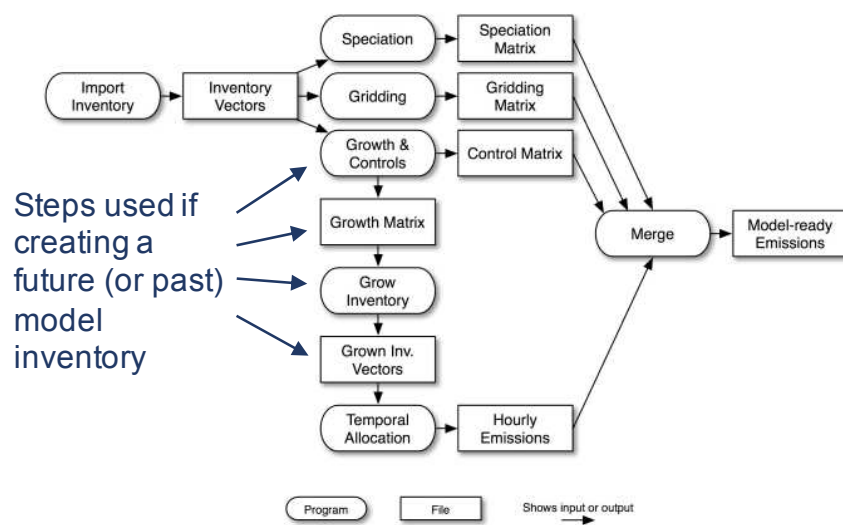


Figure 5: Major components of the SMOKE model. (Figure 2.4 from [CMAS, 2014b])

SMOKE uses a parallel approach (Figure 5) to emissions processing improving the computational efficiency of producing model-ready emissions [CMAS, 2014b]. Highlighted below are the major SMOKE steps to producing the model-ready emissions used for the 2011 base case modeling described in Chapters 3 and 4 [CMAS, 2014b]:

1. Import the NEIv2 files
2. Speciate the inventory into appropriate species for the air quality model chemical mechanism
3. Apply emissions controls (if necessary)
4. Temporally distribute emissions (including meteorology)
5. Spatially distribute emissions
6. Merge emissions into sector files with common characteristics (model-ready)

Applying the general steps outlined above, SMOKE is also used to drive the Biogenic Emissions Inventory System version 3.61 (BEISv3.61) model which produces biogenic emissions used in this air quality modeling [CMAS, 2014b]. This version of BEIS incorporates improved land use inputs from Moderate Resolution Imaging Spectroradiometer (MODIS) and the 2006 National Land Cover Database. Improved vegetation speciation inputs from the US Forest Service Forest Inventory Analysis and the US Dept. of Agriculture Biogenic Emissions Land-use Database version 4 (BELD4) are also used [Bash *et al.*, 2016]. BEIS produces emissions estimates for 33 species, but is re-specified to 18 for compatibility with the air quality chemical mechanism used in regulatory modeling (described in Section 1.5.1) [Bash *et al.*, 2016].

Boundary conditions for the model domain were generated from the Goddard Earth Observing System (GEOS) Chemical (GEOS-Chem) transport model for the year 2011 [EPA, 2014]. Gridded meteorology for 2011 was generated from the Weather Research and Forecasting (WRF) model version 3.4 [Skamarock *et al.*, 2008] using high resolution sea surface temperature data from the Group for High Resolution Sea Surface Temperature (<https://www.ghrsst.org/>) [EPA, 2014]. To produce model-ready meteorology inputs, WRF output is processed through the Meteorology-Chemistry Interface Processor version 4.3 (MCIPv4.3) to convert the WRF coordinates into

CMAQ compatible horizontal and vertical coordinates [CMAS, 2015; Otte and Pleim, 2010]. This processing is a critical step to ensure mass consistency within CMAQ [Otte and Pleim, 2010].

The modeling platform for the base year 2011 described above, and used for the model analysis throughout this thesis, is used by all modeling centers in the OTC to preserve modeling result consistency. Until recently, attainment strategies for the various states within the OTC have been developed based on this modeling platform. Faculty in the Department of Atmospheric and Oceanic Science at the University of Maryland are intimately involved with the air quality modeling performed for the State of Maryland. The work in this thesis has contributed to those efforts.

1.4: Overview of Data Products

1.4.1: Air Quality System (AQS)

The Air Quality System (AQS) is an EPA and Office of Air Quality Planning and Standards (OAQPS) network of more than 4,000 surface monitoring stations throughout the US that measure ambient air pollution. The principal goal of the network is to examine the exposure of the US population to a variety of pollutants. AQS data are available from 1980 to the present for approximately 500 species, must pass several quality control checks before distribution, and are publicly available at <https://www.epa.gov/outdoor-air-quality-data> [EPA, 2016a]. These data are used in the Commercial Marine Vessel analysis discussed in Chapter 3.

1.4.2: Ozone Monitoring Instrument (OMI)

The Ozone Monitoring Instrument (OMI) is a space-borne instrument supported by three countries. The instrument was developed by a team of scientists from the Netherlands and Finland and is deployed aboard the National Aeronautics and Space Administration (NASA) Aura satellite (launched July 15, 2004), which orbits with a suite of other satellites in the NASA A-train. Aura has a polar, sun synchronous orbit providing daily global coverage with an overpass time of approximately 13:30 local solar time (LST) at the equator [Levelt *et al.*, 2006a; Levelt *et al.*, 2006b]. The OMI instrument measures solar radiation backscattered from the atmosphere and Earth's surface within the UV/Visible wavelength range of 270-500 nm, with a spatial resolution of $13 \times 24 \text{ km}^2$ at nadir and total swath coverage of 2600 km. The OMI science team retrieves column O_3 , NO_2 , and SO_2 , three of the six criteria pollutants named by the EPA, as well as column BrO , OCIO , and HCHO [Levelt *et al.*, 2006a; Levelt *et al.*, 2006b].

A detector row anomaly appeared in the data on June 25, 2007 and impacted approximately 50% of OMI pixels in 2011, significantly reducing data density [Bucsela *et al.*, 2013]. All pixels affected by the row anomaly in the OMI products described below are filtered out for this analysis [Boersma *et al.*, 2011; Bucsela *et al.*, 2013]. We use a total of two OMI data products: daily tropospheric column NO_2 and column HCHO and three retrieval methods (two for NO_2 and one for HCHO).

For the analysis in Chapter 2, we use the Dutch Ozone Monitoring Instrument NO_2 (DOMINO) near real-time tropospheric column NO_2 retrieval [Boersma *et al.*, 2007]. As OMI passes overhead, the instrument detects a slant column of the

atmosphere using a Differential Optical Absorption Spectroscopy (DOAS) approach [Boersma *et al.*, 2004; Boersma *et al.*, 2007; Bucsela *et al.*, 2006]. Stratospheric NO₂ is estimated by a stratospheric Chemical Transport Model (CTM) called TM4 and then removed from the total slant column to produce a residual tropospheric NO₂ slant column [Boersma *et al.*, 2007]. The tropospheric slant column is converted to a vertical column using an Air Mass Factor (AMF) calculated by the TM4 global tropospheric CTM. This model simulates the expected profile of NO₂ in the troposphere, producing the necessary AMFs for the algorithm [Boersma *et al.*, 2007; Eskes and Boersma, 2003]. Data from the DOMINO retrieval, along with data for other species and satellites, can be found at <http://www.temis.nl/airpollution/no2.html>.

For the analysis in Chapter 3 we use the NASA Goddard Space Flight Center (GSFC) version 3 level 2 tropospheric NO₂ column product [Bucsela *et al.*, 2013; Krotkov *et al.*, 2017; Krotkov and Veeffkind, 2016] gridded to 0.25° latitude × 0.25° longitude resolution. The GSFC retrieval uses the Stratosphere-Troposphere Separation (STS) algorithm to calculate the tropospheric column NO₂ by removing the stratosphere component [Bucsela *et al.*, 2013]. This method assumes that over much of the world, the vast majority of detected NO₂ is in the stratosphere. Therefore, the stratospheric Vertical Column Density (VCD) from these clean or cloud obscured pixels is used to interpolate the stratospheric component of neighboring polluted or cloud free pixels [Bucsela *et al.*, 2013]. To calculate the stratospheric and tropospheric AMFs, monthly NO₂ profile climatology are derived from Global Modeling Initiative (GMI) CTM [Strahan *et al.*, 2007] simulations for 2005-2007 [Bucsela *et al.*, 2013]. For the clean pixels, an initial total VCD is calculated by dividing the total Slant

Column Density (SCD) by the stratospheric AMF. A small model estimated tropospheric NO₂ column representing the tropospheric component of the VCD for the clean region is subtracted from the total VCD. Pixels are masked for highly polluted areas where the modeled tropospheric component is larger than the stratospheric VCD [Bucsela *et al.*, 2013]. Masked pixels are removed from the initial stratospheric VCD calculation. The stratospheric component over polluted tropospheric regions is interpolated from its neighboring clean (or cloudy) pixels. Smoothing of any “hot spots” in the stratospheric VCD are done by replacing the values with the mean if the “hot spot” exceeds the mean by more than 1.5 standard deviations [Bucsela *et al.*, 2013]. Finally, the stratospheric component is removed from the total column SCD. The remaining tropospheric SCD is divided by the tropospheric AMF to yield the tropospheric VCD. This STS method is a very different approach to calculating the stratosphere NO₂ contribution than the DOMINO retrieval described earlier, yet the two retrievals produce similar results [Bucsela *et al.*, 2013].

We also use the HCHO version 3 level 2 reference sector corrected swath product from the Harvard-Smithsonian Astrophysical Observatory (SAO) retrieval [González Abad *et al.*, 2015] on a 0.25° latitude × 0.25° longitude grid. For both the GSFC OMI products (i.e., NO₂ retrieval and the SAO HCHO retrieval), we only use pixels that satisfy quality and row anomaly flags, have a cloud fraction less than 30%, and a solar zenith angle less than 70°. Additionally, data from the two outer-most pixels are removed due to their large footprint (28 × 150 km²) compared to the nadir view. The GSFC NO₂ and SAO HCHO retrievals are publicly available at <http://mirador.gsfc.nasa.gov>. Since HCHO concentrations in the stratosphere are

negligible [Millet *et al.*, 2006], the HCHO retrieval does not incorporate a stratosphere-troposphere separation algorithm. Instead, *a priori* HCHO profiles are generated using GEOS-Chem simulations. These vertical profiles are used to develop the air mass factor (AMF) calculations at various latitudes. This profile extends up to 100hPa where the concentrations of HCHO have dropped to near 0 ppb [González Abad *et al.*, 2015].

1.4.3: RAMMPP – Long Island Sound 2017

For the past two decades the University of Maryland has conducted a multifaceted effort to understand air quality in the Eastern US called the Regional Atmospheric Measurement Modeling and Prediction Program (RAMMPP) (<http://www.atmos.umd.edu/~rammpp>). RAMMPP consists of four major components: ozone forecasting, *in situ* measurements, mesoscale modeling, and chemical transport modeling. Valuable *in situ* measurements have been collected each year (except 2006) during the O₃ season (April – October) since 1997 providing information about long-term trends in the Mid-Atlantic and Eastern U.S. [He *et al.*, 2013b; Taubman *et al.*, 2004]. For the work in this thesis, a small portion of the RAMMPP data are used to analyze VOCs in the Long Island Sound (LIS) region of New York on May 17 and 18, 2017. These data were collected during a total of three research flights.

Trace gas and aerosol measurements were collected aboard the University of Maryland's Cessna 402B research aircraft. A detailed description of the aircraft instrumentation can be found in Ren *et al.* [2018]. During this campaign, O₃ was measured with a modified UV absorption Thermal Electron Model 49C analyzer; NO₂ was measured with a Los Gatos Research Model RMT-200 Cavity Enhanced Absorption Spectroscopy (CEAS) analyzer; SO₂ was measured with a modified pulsed

fluorescence Thermal Electron Model 43C analyzer; CO, CH₄, CO₂, and H₂O were measured with a Picarro Cavity Ring-Down Spectrometer (CRDS) Model G2401-m [Ren *et al.*, 2018]. Additional aerosol optical properties were measured, and grab cans were used to sample atmospheric air at various intervals throughout the research flights. All VOCs were analyzed by the Maine Department of Environmental Protection using a Gas Chromatography – Mass Spectrometer (GC-MS) [Ren *et al.*, 2018]. A map of the three research flights is shown in Figure 6. Research Flight (RF) 1 was from 14:50 to 18:10 LST on May 17, RF2 was from 7:40 to 10:25 LST on May 18, and RF3 was from 13:10 to 16:40 LST on May 18, 2017.

A total of 35 canisters sampling ambient air were collected, and 50 VOC species were reported during the RAMMPP LIS 2017 field campaign. The list of species can be found in Appendix A. These data are used in the VOC analysis over Long Island Sound discussed in Chapter 4.

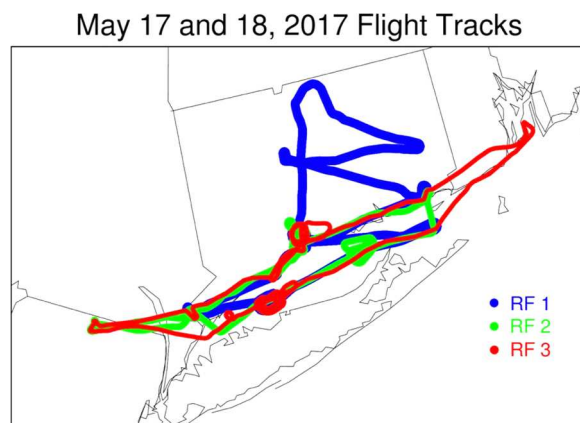


Figure 6: Flight tracks for the 3 research flights over the Long Island Sound during May 17th and 18th, 2017. RF1 (blue) was in the afternoon on May 17th, RF2 (green) was in the morning on May 18th, and RF3 (red) was in the afternoon on May 18th.

1.5: Modeling and Emissions Inventories

1.5.1: CMAQ version 5.0.2 – UMD

For the analysis in Chapters 3 and 4 we use the EPA Community Multiscale Air Quality (CMAQ) model version 5.0.2 [Byun and Schere, 2006; CMAS, 2014a]. CMAQ is a 3-dimensional Eulerian chemical transport model simulating chemical processes in the troposphere [EPA, 2017d]. Mathematical calculations describing transport across grid cells and the chemical reactions that occur within them are based upon conservation of mass for the given time step [EPA, 2017d]. A basic schematic of the various components of CMAQ is described in Figure 7. Generation of the required inputs to the CMAQ Chemical Transport Model (CCTM) was detailed previously in Section 1.3.1.

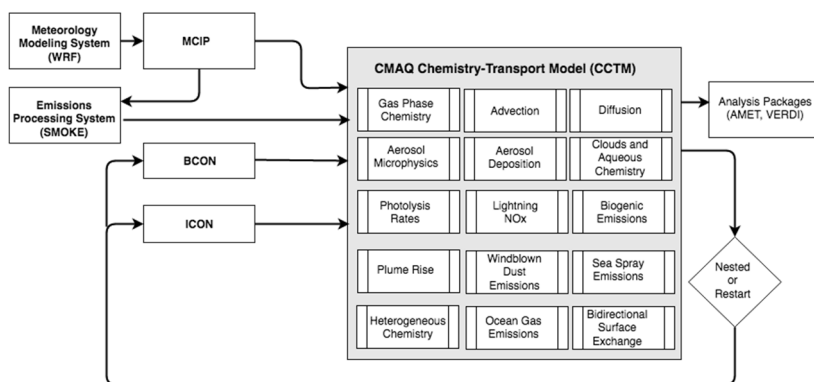


Figure 7: The CMAQ modeling system detailing the necessary emission and the various modules of the CMAQ Chemistry-Transport Model (CCTM). (Figure 4-6 from EPA [2017d])

The core 2005 Carbon Bond (CB05) chemical mechanism consists of 156 gas-phase reactions with 51 species representing the photochemistry of the troposphere [Yarwood *et al.*, 2005]. The CMAQ platform used in this analysis incorporates a modified version of CB05 consisting of 172 gas phase reactions and ~70 species called

CB05cltx_AE5_aq. This chemical mechanism includes the CB05 extended toxics reactions, updates to all rate reactions as of 2008, and the addition chlorine reactions for certain VOCs. All reactions and species for chemical mechanism used in this CMAQ model platform are detailed in Appendix B. It is important to note that this is the regulatory-approved CMAQ platform used by state and local agencies to develop attainment strategies. The modifications to the base CB05 mechanism were made by the developers and approved by the regulatory agencies.

CB05 is a simplified chemical mechanism designed to model the oxidative chemistry of the atmosphere for important pollutants like O₃, NO_x and particulate matter [Yarwood *et al.*, 2005]. The Carbon Bond chemical mechanisms are representative of the numerous reactions that occur in the atmosphere yet simple enough to be computationally feasible for use in air quality models [Yarwood *et al.*, 2005]. A much more expansive mechanism called the Master Chemical Mechanism is available (<http://mcm.leeds.ac.uk/MCM/>) but is best used in 0-dimensional box modeling analyses.

An updated version of CB05 called the Carbon Bond Mechanism version 6 (CB6) was released in 2010 [Yarwood *et al.*, 2010]. This mechanism consists of 218 gas-phase reactions with 77 species [Yarwood *et al.*, 2010]. The additional reactions and species improve secondary organic aerosol and oxidant formation from long-lived VOC species in the model [Yarwood *et al.*, 2010]. Further improvements were released in 2013 as the Carbon Bond version 6 revision 2 (CB6r2) mechanism [Hildebrandt Ruiz and Yarwood, 2013; Yarwood *et al.*, 2014]. This mechanism has 80 distinct species, the same gas-phase reactions as CB6, all reactions are updated to incorporate

the 2013 International Union of Pure and Applied Chemistry (IUPAC) reaction rates, and partitioned organic nitrates more appropriately which improved O₃ simulations [Hildebrandt Ruiz and Yarwood, 2013; Yarwood *et al.*, 2014]. In 2014 an optional halogen mechanism suite consisting of 88 reactions and 41 species called CB6r2h was released [Yarwood *et al.*, 2014]. Halogens such as iodine, chlorine, and bromine are important to tropospheric O₃, especially in coastal regions influenced by sea-salt. Under some conditions, halogens produce tropospheric O₃; under others, they destroy it [Chang *et al.*, 2002]. This detailed mechanism aims to represent this important halogen chemistry, however it is computationally expensive and therefore remains an optional addition to CB6r2 [Yarwood *et al.*, 2014]. A mechanism update to improve representation of wintertime O₃ formation due to snow cover was released in 2015 (CB6r3) [Emery *et al.*, 2015]. A new version called CB6r4 was released in 2016, incorporating the wintertime O₃ improvements from CB6r3 and the 16 most important halogen reactions for O₃ from CB6r2h [Emery *et al.*, 2016]. These reactions are from the iodine mechanism in CB6r2h and are the most important for ozone depletion by iodine [Emery *et al.*, 2016]. Improvements to reaction product yields and IUPAC reaction rates were also performed for the CB6r4 update [Emery *et al.*, 2016].

These updated chemical mechanisms were not incorporated into the modeling performed in this thesis; however, discussion of their benefits is given in Section 5.3 and references the relationship between halogens and tropospheric O₃. A new version of CMAQ (v5.2) was released in 2017 and is capable of implementing the CB6r3 mechanism (<https://www.cmascenter.org/>) [EPA, 2017d]. At the time of analysis, CMAQv5.0.2 was the only appropriate model platform available.

1.5.2: CMAQ version 5.0.2 – ARL

For the analysis in Chapter 4, the National Oceanic and Atmospheric Administration (NOAA) National Air Quality Forecast Capability (NAQFC) system was used to forecast an air quality O₃ event in the Long Island Sound, NY (LIS) region [Lee *et al.*, 2017; Tong *et al.*, 2015; Tong *et al.*, 2016]. The NAQFC system includes three major components: chemistry, meteorology, and emissions. The NAQFC uses the CMAQv5.0.2 and CB05 to predict O₃, PM_{2.5}, and their precursors concentrations over the CONTiguous US (CONUS) states, northern Mexico and southern Canada. Hourly meteorological data from the NOAA National Center for Environmental Prediction's North American Mesoscale (NAM) forecast model are used in the CMAQ simulations. The NAM relies on the meteorology dynamic core of WRF-NMM (Non-hydrostatic Mesoscale Model) on the B grid (WRF-NMMB) with upgraded tracer advection scheme [Janjic, 2003]. Boundary conditions used in the simulation are monthly averaged profiles extracted from the 2006 simulation with Harvard University's GEOS-Chem model [Zhang *et al.*, 2011]. The NAQFC emission dataset includes gaseous and particulate emissions from anthropogenic sources (area, mobile, and point) and natural sources (biogenic, soil, and sea salt). For sources outside the US, the Mexico NEI from 2012 and the Environment Canada and Climate Change Emission Inventory from 2010 are used to represent emissions Mexico and Canada. Inside the US, area and mobile sources are generated from a mixture of the EPA 2011 NEI and projected 2012 NEI using the Cross-State Air Pollution Rule (CSAPR) projection data [Pan *et al.*, 2014]. Point sources, including both Electrical Generating Units (EGUs) and non-EGUs, are generated using the 2011 NEI as baseline, with the EGU sources

further updated to 2016 with the Continuous Emission Monitoring System (CEMS) data. The performance of this modeling system to predict O₃ and PM_{2.5}, and their key precursors has been extensively evaluated with ground and field campaign data [*Lee et al.*, 2017; *Tong et al.*, 2016].

1.5.3: Alpha2 Emissions Inventory

As mentioned in Section 1.3, modeling is essential for developing and testing strategies to improve compliance with air quality regulation. Reliable emissions are required to effectively model various attainment strategies. For modeling efforts in the OTC, the Mid-Atlantic Regional Air Management Association (MARAMA) develops emissions based on the EPA NEIs [*EPA*, 2015a; *McDill et al.*, 2015]. For the analysis in Chapter 3, we use the MARAMA Alpha2 version of emissions, released in November 2015 and based primarily on the 2011 NEI version 2 (2011 NEIv2). The non-road inventory is the only emissions sector to use input from the 2011 NEI version 1 data [*McDill et al.*, 2015]. Emissions inventories are also projected to future years to aid in the development of effective air quality policy (SIPs). In the Alpha2 inventory, emissions for all sectors (excluding biogenic emissions) are projected from the base year (2011) to the future year (2018). For most sectors, emissions inventories for model year 2018 are generated by applying growth factors based on input from State, Local, and Tribal (SLT) agencies to the 2011 base case emissions. Projected emissions for the non-road and on-road mobile sectors are created using input from version 1 of the 2018 EPA's National Emissions Inventory [*McDill et al.*, 2015]. The Alpha2 inventory also projects emissions to 2028 using growth factors, however these emissions were not used in the analysis in Chapter 3.

All emissions are categorized into seven major source sectors [McDill *et al.*, 2015]:

1. **Point sources:** emission sources that have a specific latitude and longitude location, emitting at a single source, like a smoke stack.
2. **Aircraft, GSE, APU:** emissions sources from aircraft engines, airport ground support equipment (GSE), and auxiliary power units (APUs).
3. **Non-point sources:** emissions sources that are too small or too numerous to model as individual point sources. Emissions are calculated at the county level.
4. **Non-road mobile sources:** generated by the NONROAD model in the EPA National Mobile Inventory Model (NMIM). Represents emissions from a variety of combustion sources using four different fuel types.
5. **Rail/CMV sources:** locomotive and commercial marine vessel emissions.
6. **Fire sources:** emissions related to biomass burning.
7. **Biogenic sources:** emissions from vegetation and soils, computed using the Biogenic Emissions Inventory System (BEIS) model.

Further partitioning of these seven main categories into sub-categories is implemented to facilitate comparison to the EPA NEI and provide more detailed sector-specific analysis [McDill *et al.*, 2015]. Daily Alpha2 emissions files for 48 sub-categories are created using the SMOKE model described in Section 1.3.1. These files allow specific emissions adjustments to be made to reflect targeted controls on individual sectors. Adjustments we implement for the Alpha2 inventory files are discussed in Chapter 3. Additional analysis with the Alpha2 inventory is performed in Chapter 4. Further information about the inventory can be found at <http://www.marama.org/technical-center/emissions-inventory/2011-inventory-and-projections>.

1.5.4: Beta2 Emissions Inventory

An update to the Alpha2 inventory called the Beta2 inventory was released in July 2017 [McDill *et al.*, 2017]. This inventory is structurally organized in the same manner as the Alpha2 inventory, but includes updated emissions information based on the EPA version 6.2 modeling platform [OAQPS, 2015], and the version 6.3 modeling platform [OAQPS, 2016] for on-road mobile emissions [McDill *et al.*, 2017]. The most relevant improvement for the analysis in Chapter 4 reflected in the Beta2 inventory are to the mobile inventory sector and include [McDill *et al.*, 2017]:

- Non-road mobile: Emissions from the 2011 NEIv2 with additional corrections from Delaware.
- On-road mobile: using a new version of the MOtor Vehicle Emission Simulator (MOVES) model (version 2014a) to generate emission factors using CB05 chemistry. Corrections were applied for 15 counties in New Jersey.

More information about the Beta2 inventory and its specific improvements can be found at: <http://www.marama.org/technical-center/emissions-inventory/2011-2017-beta-regional-emissions-inventory>.

We worked closely with Eric Zalewsky from the New York Department of Environmental Conservation to generate the merged Beta2 emissions file for use in CMAQ. Throughout this process, we found an error in the Private Fuel Containers (PFC) emissions files and developed a series of merge scripts (used in SMOKE) to create the final merged emissions file. The error in the PFC files was announced to the modeling community and the corrected files were provided to the various modeling centers in the winter of 2018.

While most sector files were updated from the Alpha2 inventory for 2011, a total of 17 files remained un-adjusted for the Beta2 inventory [McDill *et al.*, 2017]. Merging files from the Alpha2 and Beta2 inventories resulted in inconsistent units for two variables, requiring a convoluted process of batch merging compatible individual sector files, renaming the problem variables, and finally merging the partly-merged files. Complete merged files for modeling years 2011 and 2017 were created. Analysis using these files is discussed in Chapter 4.

1.6: Research Objectives

The goal of this research is to answer four overarching questions, and provide this information to the air quality community:

1. Does the regulatory CMAQ modeling platform accurately represent the atmosphere? If not, what modifications can be made to improve this representation? How does this impact modeled O₃?
2. Can we use CMAQ and relevant data sources to determine photochemical O₃ production regimes across the model domain?
3. How does the photochemistry in the New York City pollution plume impact levels of O₃ in heavily populated areas downwind?
4. What are the important emissions sectors contributing to O₃ production in heavily populated, near-coastal, urban centers within the model domain?

The research presented in this thesis will provide the air quality modeling community answers to these questions, and contribute important insight into CMAQ model performance, a necessary tool for attainment strategy development.

Chapter 2: Developing a Top-Down Emissions Inventory

2.1: Findings from previous studies

As discussed in the introduction, emissions inventories used in regulatory air quality modeling are developed based on the EPA NEI. For some sectors, inventory estimates use *in situ* data collected at frequent intervals to derive values. For others, the estimates are based upon data that were collected years (sometimes decades) earlier and grown to more recent years [EPA, 2015a]. Development of emissions inventories is time consuming, requiring extensive inter-agency cooperation, making it difficult to quickly produce accurate and reliable inventories. Creating a robust method to generate inventories based on recent, reliable satellite data would allow inventories to reflect current emissions more accurately, and significantly reduce the time required to gather the necessary data. Additionally, inventories would be based on observed data for the year, rather than grown from earlier years.

In the US, the Commercial Marine Vessel (CMV) inventory is based on data collected in 2002 which is grown to the inventory year [EPA, 2002; 2009b]. The CMV sector is unique because most ship emissions are released in the open ocean, in regions that are relatively pristine along well traveled trading routes. This is indicated as elevated NO₂ along distinct ship tracks relative to the background when viewed from space. A few studies have explored the possibility of developing a top-down emissions inventory in various parts of the world specifically for the CMV sector. Their work will be discussed in the next few sections.

2.1.1: NO₂ Measurements from SCIAMACHY

In 2004, a study by *Richter et al.* discussed the uncertainty in emissions estimates from available inventories. *Corbett and Koehler* [2003] published an inventory based on total fuel burn for internationally registered ships that carry goods weighing greater than 100 tons, and found a total annual fuel burn of 289 Mt [*Richter et al.*, 2004]. This was nearly twice that of the fuel consumption estimate published by the international marine bunker industry [*Richter et al.*, 2004], underlining the difficulty of producing an accurate and reliable bottom-up inventory.

The coarse resolution of satellite imagery can make developing a reliable top-down emissions inventory difficult. For their study, *Richter et al.* [2004] used tropospheric column NO₂ observations from the SCanning Imaging Absorption spectrometer for Atmospheric CHartographY (SCIAMACHY) instrument to characterize commercial shipping emissions. This instrument is a nadir viewing instrument that uses the DOAS technique to measure NO₂ [*Bovensmann et al.*, 1999; *Eichmann et al.*, 2004]. Backscattered light from the Earth's surface and atmosphere between wavelengths of 425 and 450nm is used to measure the total SCD for NO₂ [*Burrows et al.*, 1999]. First, the stratosphere is measured in limb mode which requires observing the atmosphere tangentially to the Earth's surface forward along the flight path, then the SCD is measured in nadir mode by observing the atmosphere perpendicular to the Earth's surface [*Bovensmann et al.*, 1999; *Eichmann et al.*, 2004; *Noel et al.*, 2002; *Richter and Burrows*, 2002] The stratospheric component measured in limb is then subtracted from the total column measured in the nadir mode to return a tropospheric SCD [*Bovensmann et al.*, 1999; *Eichmann et al.*, 2004; *Noel et al.*,

2002]. For SCIAMACHY, a radiative transfer program called the SCIAMACHY Radiative TRANSfer Model (SCIATRAN) is used to calculate tropospheric AMFs. These are quantities that relate the tropospheric profile of absorption of NO₂ to the measured SCD and are necessary to calculate the tropospheric VCD, the desired satellite measurement [Bovensmann *et al.*, 1999; Burrows *et al.*, 1999; Rozanov *et al.*, 2002].

SCIAMACHY has a ground resolution of $30 \times 60 \text{ km}^2$ and a total swath length of 960 km [Bovensmann *et al.*, 1999; Noel *et al.*, 2002]. Over the course of the last decade, satellite technology has improved immensely, and current retrievals have much higher resolution. The recent launch of the Copernicus Sentinel-5 Precursor satellite by the European Space Agency (ESA) in October of 2017 has a ground resolution of $7 \times 7 \text{ km}^2$ at nadir [Veeffkind *et al.*, 2012], significantly improving spatial detail from space.

This study by Richter *et al.* [2004] found a commercial shipping signal in the SCIAMACHY tropospheric column NO₂ in the Red Sea, Indian Ocean, and South China Sea regions. These areas were chosen due to their large volume of ship traffic and localization along narrow shipping lanes [Franke *et al.*, 2009; Richter *et al.*, 2004]. Comparison of the satellite data to an available commercial shipping inventory at the time of the study shows good spatial agreement (Figure 8a and b respectively).

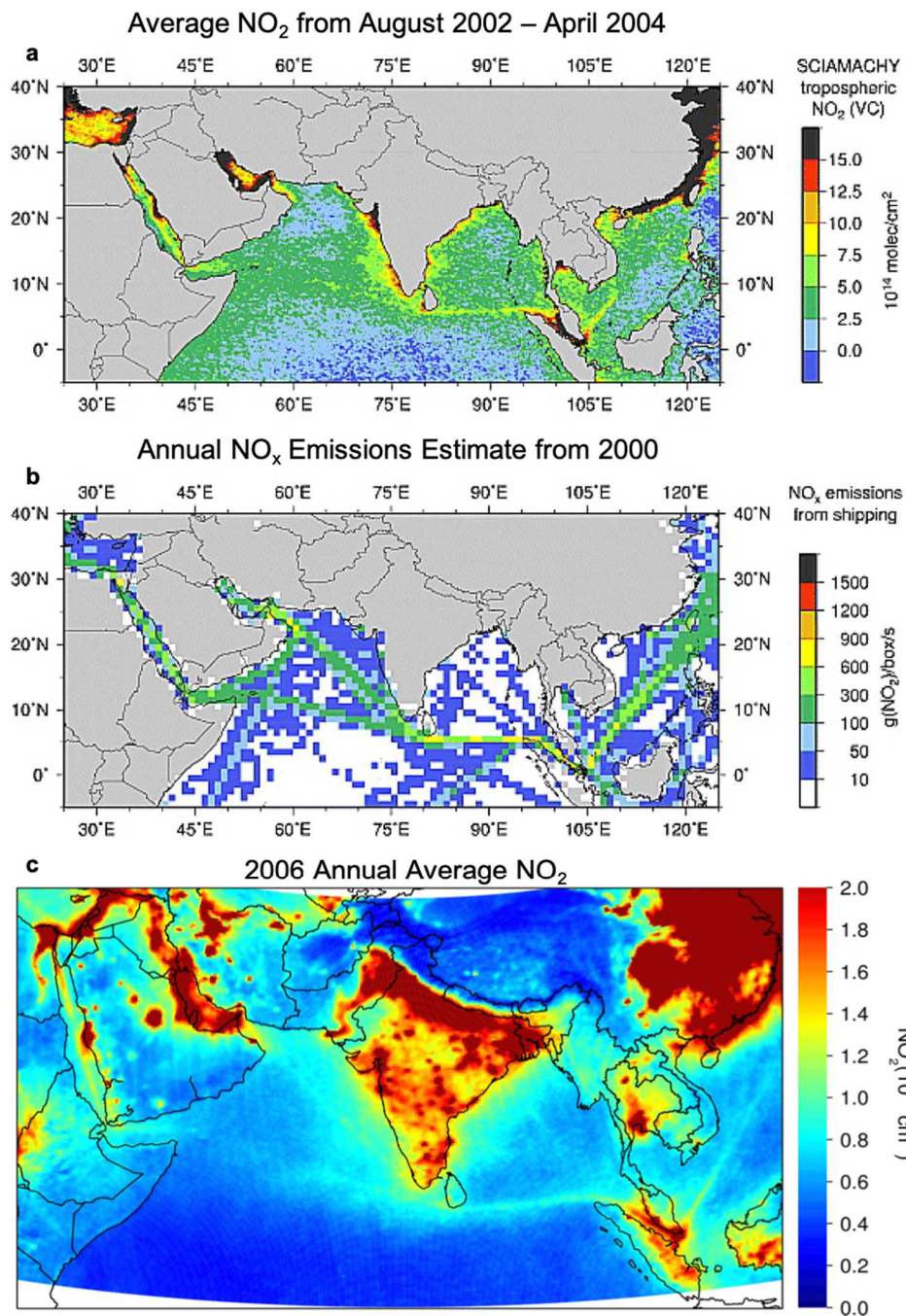


Figure 8: NO₂ from commercial shipping activity from the Red Sea through the Indian Ocean and South China Sea. (a) Tropospheric NO₂ columns derived from SCIAMACHY data from August 2002 to April 2004 (Figure 1a from Richter *et al.* [2004]), (b) NO_x emissions estimated for the year 2000 from the Corbett and Koehler [2003] inventory and AMVER vessel distribution data from Endresen *et al.* [2003] (Figure 1b from Richter *et al.* [2004]) (c) Annual average tropospheric NO₂ from OMI for the year 2006.

Using simple assumptions about NO₂ loss processes in the troposphere, *Richter et al.* [2004] were able to calculate lower bound estimates of annual shipping emissions for the various regions. In some locations, satellite-derived emissions were over a factor of 2 lower than inventory estimates [*Richter et al.*, 2004]. While this method is crude, it does provide evidence for creating a lower-bound estimate for commercial shipping emissions, especially in heavily trafficked areas. This is further supported by Figure 8c, which shows elevated levels of NO₂ compared to the background in the same locations shown by *Richter et al.* [2004]. We use the DOMINO level 2 data (described in Section 1.4.2) from the TEMIS website (www.temis.nl/airpollution/no2.html), create daily 0.25° latitude × 0.25° longitude level 3 gridded files, and averaged over the year 2006. All level 2 data were filtered for quality flags provided with the retrieval and require each unflagged pixel to have a Cloud Radiance Fraction (CRF) < 30% and a surface albedo < 30%. The outer 15 pixels on each side of the OMI swath are removed so only the highest resolution near-nadir pixels are used (Figure 9). Clearly defined shipping lanes through the Red Sea into the Arabian Sea, Indian Ocean, and up into the South China Sea are evident.

Figure 8a and c show that satellites can detect shipping emissions in the NO₂ retrievals on a large scale, especially in clean environments with few other anthropogenic sources of NO_x. Additionally, global emissions inventories seem to accurately represent the locations of heavily traveled shipping lanes in the Indian Ocean and Red Sea regions of the world. However, generating a top-down emissions inventory from satellites that is usable for modeling purposes was not possible in 2004.

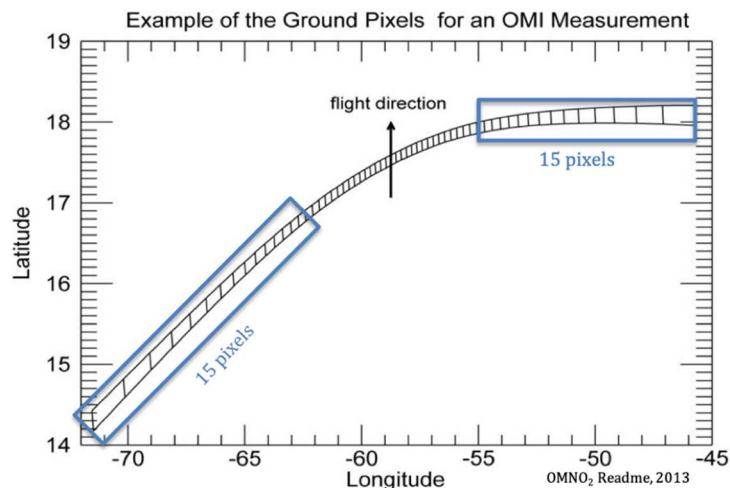


Figure 9: Description of the OMI swath (figure from *Krotkov* [2013]). The boxed OMI pixels are removed, to ensure only the highest resolution near-nadir pixels are used.

2.1.2: A Different Approach to the Top-Down Inventory using SCIAMACHY data

Development of bottom-up emissions inventories is time-intensive and results in delayed release, often years past the relevant model year. The goal of the *Lamsal et al.* [2011] study is to provide near-real time constraints on bottom-up inventories using GEOS-Chem, a global atmospheric chemistry model, and SCIAMACHY satellite data [Lamsal et al., 2011; Walker et al., 2010]. First, the relationship between changes in surface emissions of NO_x and column NO_2 values must be established. To do this, [2011] performs two model simulations using GEOS-Chem version 8-01-04 at a 1° latitude \times 1.25° longitude resolution, driven by GEOS-4 assimilated meteorology. The first simulation uses a global NO_x emissions inventory for 2006 (created by [2011] from a variety of sources), and the second uses emissions where the global NO_x inventory for 2006 is perturbed by an arbitrarily chosen amount of 15%. Results from the two model simulations are then used in the equation from *Lamsal et al.* [2011]:

$$\frac{\Delta E}{E} = \beta \times \frac{\Delta \Omega}{\Omega}. \quad (30)$$

Equation (30) relates changes in surface NO_x emissions to simulated column NO₂ values. In Equation (30), E is the 2006 global NO_x emissions, ΔE is the imposed perturbation to the NO_x emissions, Ω is the simulated NO₂ column from GEOS-Chem based on the 2006 NO_x emissions, and $\Delta \Omega$ is the change in simulated NO₂ column based on the perturbed NO_x emissions. The β is a constant calculated for each model grid cell that describes how the change in the NO_x surface emissions affects the NO₂ column [Lamsal *et al.*, 2011]. In order to apply Equation (30) to generate a top-down NO_x emissions inventory, one must make some (likely unrealistic) assumptions:

- (1) GEOS-Chem realistically simulates the global atmosphere.
- (2) A modification of 15% to the global NO_x emissions is representative of how NO_x emissions are adjusted from the base year (2006) to the projected year.
- (3) Changes in the NO₂ column in response to the adjustments of surface NO_x emissions detected in the GEOS-Chem model and represented by the β term in Equation (30), are representative of the differences in the tropospheric column NO₂ measured by satellites for the reference and future inventory years.

Lamsal et al. [2011] then applied this relationship to average tropospheric column NO₂ from SCIAMACHY. To do this, Equation (30) is modified to solve for the adjusted NO_x emissions inventory rather than the β term. This equation is found in *Lamsal et al.* [2011]:

$$E_j = \left(1 + \beta \frac{(\Omega_j - \Omega_i)}{\Omega_i} \right) E_i. \quad (31)$$

In Equation (31), E_j represents the monthly NO_x emissions inventory for future year j, E_i represents the corresponding monthly 2006 NO_x emissions inventory, Ω_i represents the corresponding monthly average tropospheric NO₂ column from SCIAMACHY for 2006, Ω_j represents the SCIAMACHY corresponding monthly average tropospheric NO₂ column for future year j, and β is carried over from Equation (30). To minimize error, only grid boxes where anthropogenic emissions dominate and large tropospheric NO₂ columns are detected are used. This limits the grid boxes used in this analysis to heavily populated cities, which account for approximately 80% of the global anthropogenic NO_x emissions [Lamsal *et al.*, 2011]. In this analysis, Equation (31) is applied to calculate a global NO_x inventory for 2009.

While this study outlines a fairly simple method to calculate a near-real time NO_x emissions inventory, it requires many assumptions and is best used to calculate annual anthropogenic emissions across all sectors. Since the goal is to develop a top-down emissions inventory for the commercial marine vessel inventory sector, an expanded and more specific version of this method is necessary.

2.1.3: Providing Constraints for Commercial Marine Vessel Emissions in Europe

The previously discussed studies show that satellite NO₂ data can provide detailed geographic information about commercial shipping lanes and can provide constraints to annually temporalized anthropogenic NO_x emissions estimates, however

neither study provides a robust reliable method for creating a top-down emissions inventory for the CMV sector.

Two studies published by *Vinken et al.* [2011] and [2014] use the DOMINO data, along with GEOS-Chem and a Gaussian plume dispersion model called the PARAmetrization of emitted NO_x (PARANOX) [*Vinken et al.*, 2011] to provide constraints to the commercial shipping inventory in Europe [*Vinken et al.*, 2014]. An “instant dilution” approach is widely used to represent the complicated chemistry from ship emissions within a model grid cell [*Vinken et al.*, 2011; *Vinken et al.*, 2014]. This approach assumes localized emissions are immediately diluted to the entire grid cell and results in lifetimes for NO_x emissions from ships that are up to a factor of 2.5-10 too long. This simplification produces unrealistically high concentrations of NO_x and O₃ within the model [*Chen et al.*, 2005; *Davis et al.*, 2001; *Vinken*, 2010; *Vinken et al.*, 2011; *von Glasow et al.*, 2003].

The “instant dilution” method or something similar used by CMAQ may be a contributing factor to the very high values of O₃ produced over large bodies of water in model simulations [*Canty et al.*, 2015; *Loughner et al.*, 2014; *Trail et al.*, 2014; *Travis et al.*, 2016; *Vinciguerra et al.*, 2017]. With a typical CMAQ grid cell of 12 km, the “instant dilution” method immediately distributes very localized NO_x emissions from ships over the entire grid cell, diluting important localized chemical and dynamical impacts of high NO_x emissions near the source. This method is an unrealistic representation of how CMV emissions (and all point-source emissions) are dispersed in the actual atmosphere. Enhanced representation of the important chemistry on smaller spatial scales that occurs near the emission source, and in the few hours

immediately after emission will improve model production of atmospheric pollutants and surface O₃.

To address the limitations of the “instant dilution” method, *Vinken et al.* [2011] suggests a plume-in-grid approach to represent chemical aging of ship emissions in the 3 hours after emission. This method simulates production of the secondary pollutants that result from ship emissions [*Vinken et al.*, 2011]. Briefly, the PARANOX model assumes dispersion represented by 10 concentric rings extending from the pollution source. Chemical reactions and species aging are carried out within each ring in response to meteorologically driven dispersion. This method simulates the rapid chemical conditions that occur within the model grid cell immediately after emission from the shipping vessel [*Vinken*, 2010; *Vinken et al.*, 2011].

Figure 10 shows how the PARANOX model compares to measurements from the NOAA 2002 Intercontinental Transport and Chemical Transformation (ITCT) field campaign conducted ~100 km off the coast of California [*Chen et al.*, 2005]. At time of emission and extending to ~40 min after emission, NO_x concentrations are very high as they represent the initial emissions from commercial marine vessels (Figure 10a), translating to initially low O₃ concentrations due to NO titration (Figure 10b). As OH is created by reaction of NO + HO₂, concentrations increase (Figure 10c) and eventually contribute to a shorter modeled lifetime of NO_x using the plume-in-grid approach than is modeled using the “instant dilution” method. This reaction also produces NO₂ which leads to the net photochemical production of O₃ as shown in Figure 10b [*Vinken et al.*, 2011]. Figure 10d shows that the instantaneous lifetime of NO_x using this plume-in-grid approach is actually on the order of 2-4 hours rather than

~1 day as assumed by models using the “instant dilution” method [Vinken *et al.*, 2011]. This metric is calculated by dividing the NO_x concentration by the HNO_3 production rate for a given time because the dominant sink for NO_2 during the day is the production of HNO_3 via reaction with OH [Vinken *et al.*, 2011].

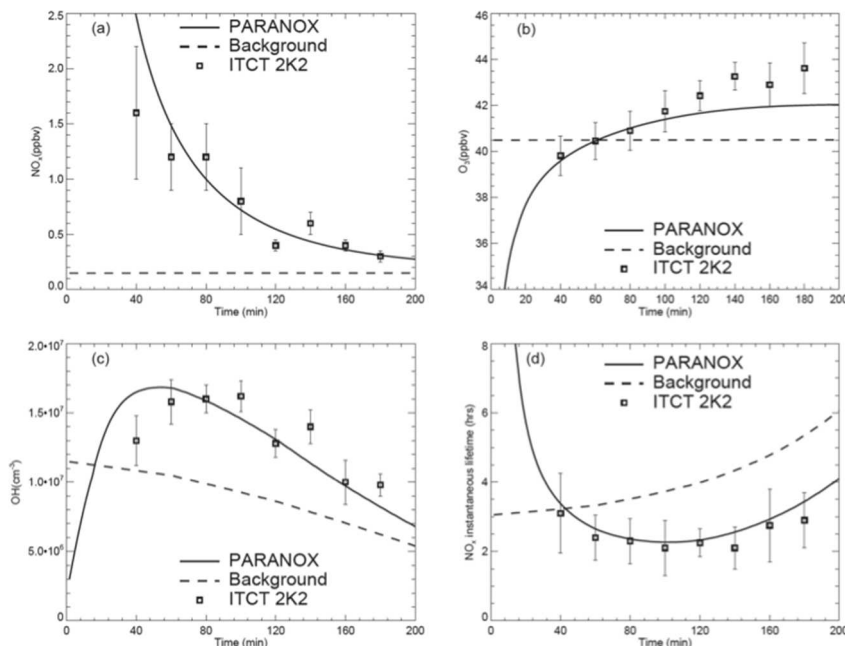


Figure 10: Comparison of observations from the 2002 ITCT field campaign [Chen *et al.*, 2005] and simulated concentrations of (a) NO_x (b) O_3 (c) OH and (d) NO_x instantaneous lifetime, from the PARANOX model. Concentrations are cross-sectional averages of ship plumes, released at 12:00 LT and depict how the species age in the 3 hours after emission. The dotted line represents background concentrations for each species plotted. This is Figure 1 from [Vinken *et al.*, 2011].

When compared to the 2002 ITCT data, PARANOX is able to accurately reproduce the concentrations of chemically important species. To reduce computational expense, Vinken *et al.* [2011] produce look up tables used in GEOS-Chem modeling to represent the small-scale complex non-linear chemistry that takes place in the 5 hours after initial emission, and within the GEOS-Chem grid cell.

Utilizing the plume-in-grid method detailed in Vinken *et al.* [2011] to correctly represent modeled ship emissions, Vinken *et al.* [2014] apply this tool to develop

constraints on NO_x emissions from ships in Europe using satellite observations and the GEOS-Chem model. Modeled and observed tropospheric NO₂ columns from GEOS-Chem and the DOMINO retrieval are compared. The OMI NO₂ retrievals are adjusted to use GEOS-Chem *a priori* column profiles rather than the TM4 profiles provided with the DOMINO retrieval to preserve consistency when comparing modeled and observed tropospheric NO₂ columns. *Vinken et al.* [2014] apply the general principals outlined by *Lamsal et al.* [2011] and add terms to account for the sensitivity of OMI NO₂ columns to the model *a priori* profile represented by the equation from *Vinken et al.* [2014]:

$$\gamma = \frac{\Delta\Omega_{OMI}/\Omega_{OMI}}{\Delta\Omega_{GC}/\Omega_{GC}}. \quad (32)$$

In Equation (32), Ω_{OMI} is the retrieved tropospheric NO₂ column, Ω_{GC} is the simulated tropospheric column using the same *a priori* profile as the OMI retrieval, $\Delta\Omega_{OMI}$ represents the change in the OMI retrieval due to changing the model *a priori* profile, and $\Delta\Omega_{GC}$ represents the change in simulated column NO₂ due to perturbed CMV emissions (same method for calculating β) [*Vinken et al.*, 2014]. Values of γ equal to 0 indicate no model sensitivity, and values of 1 indicate the retrieved column is entirely dependent upon the model *a priori* profile [*Vinken et al.*, 2014]. Incorporating the γ term into Equation (31) results in the following equation from *Vinken et al.* [2014]:

$$E_{top-down} = E_{a\ priori} + \left(\left(\frac{\Omega_{OMI} - \Omega_{GC}}{\Omega_{GC}} \right) \times \beta \times E_{a\ priori} \right) + \left(\left(\frac{\Omega_{OMI} - \Omega_{GC}}{\Omega_{GC}} \right) \times \gamma \times \beta \times E_{a\ priori} \right). \quad (33)$$

In Equation (33), the first term on the right side is the *a priori* bottom-up emissions used in GEOS-Chem modeling, the second term represents the scaling of the *a priori*

emissions to match the retrieved and modeled tropospheric NO₂ columns, and the third term describes the influence of the *a priori* model profile on the necessary *a priori* emissions adjustment [Vinken *et al.*, 2014]. This method produces GEOS-Chem modeled NO₂ columns that correlate well with OMI observed NO₂ columns in the Baltic Sea, Bay of Biscay, Mediterranean Sea, and North Sea study regions [Vinken *et al.*, 2014].

The major findings from Vinken *et al.* [2014] include:

- (1) Tropospheric NO₂ columns observed from satellites are sensitive to induced changes in shipping emissions over shipping lanes, therefore correct representation of this emissions sector is vital to accurately simulate chemistry and pollution transport.
- (2) Satellite observed columns are influenced by the choice of model *a priori*. Use of consistent emissions, dynamical schemes, etc. between the model used to calculate the satellite *a priori* profiles (TM4 for DOMINO) and the comparison model (GEOS-Chem in this case) is necessary to reduce error.
- (3) A calculated top-down annual emissions inventory for the CMV inventory sector is possible. Top-down estimated emissions closely match bottom-up annual emissions estimates for shipping lanes in waters around Europe, especially emissions from the AMVER-ICODS inventory [Wang *et al.*, 2008].

Based on these findings, it is evident that Vinken *et al.* [2014] have developed a sound method to calculate an annual estimate of commercial marine vessel emissions using satellite observations. This method can ensure the annual EPA NEI values are

within a reasonable uncertainty of satellite observations, but it cannot produce constraints for daily emissions values that are used to conduct air quality modeling. The goal is to apply this method to estimate a data-derived top-down emissions inventory value for the CMV sector for the near-coastal US region comparable to the EPA 2011NEI for the O₃ season (April – October).

2.2: Creating a Commercial Marine Vessel Inventory for the East Coast US

2.2.1: Signal from OMI

We establish that shipping lanes are visible in the tropospheric column NO₂ (DOMINO) retrieval and are coincident with the shipping lanes found in *Vinken et al.* [2014] (Figure 11). The white boxes in panels a and b of Figure 11 highlight the consistent signal in the tropospheric column NO₂ from OMI, an annually persistent feature. While it is established that a shipping signal can be found in European waters and the Indian Ocean [*Richter et al.*, 2004], is there evidence of a signal from commercial shipping emissions in the satellite data off the coast of the Eastern US?

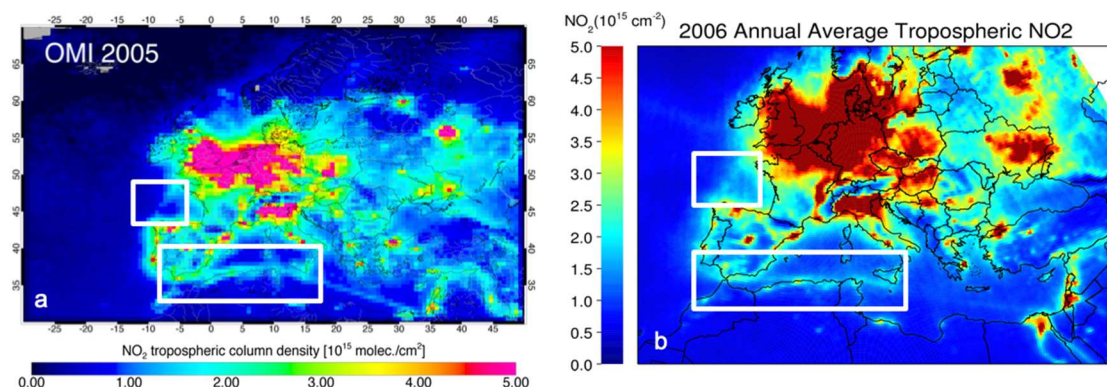


Figure 11: Annually averaged DOMINO tropospheric column NO₂ for 2005 gridded on $0.5^\circ \times 0.66^\circ$ grid (a). Pixels used have a Cloud Radiance Fraction (CRF) < 50% and an albedo of < 20%. The figure comes from Figure 3 of [*Vinken et al.*, 2014]. Annually averaged DOMINO tropospheric column NO₂ for 2006 gridded on a $0.25^\circ \times 0.25^\circ$ grid (b). Pixels are filtered the same way as *Vinken et al.* [2014].

Figure 12a shows what we perceive to be a signal from shipping activity across the Atlantic Ocean. Elevated NO_2 emissions are shown in areas across the Atlantic that are known to be heavily trafficked shipping lanes (Figure 12b). Unfortunately, the continental outflow from the US (dark red color in Figure 12a) extends well into the Atlantic Ocean, obscuring any NO_2 signal from CMV emissions in the satellite data.

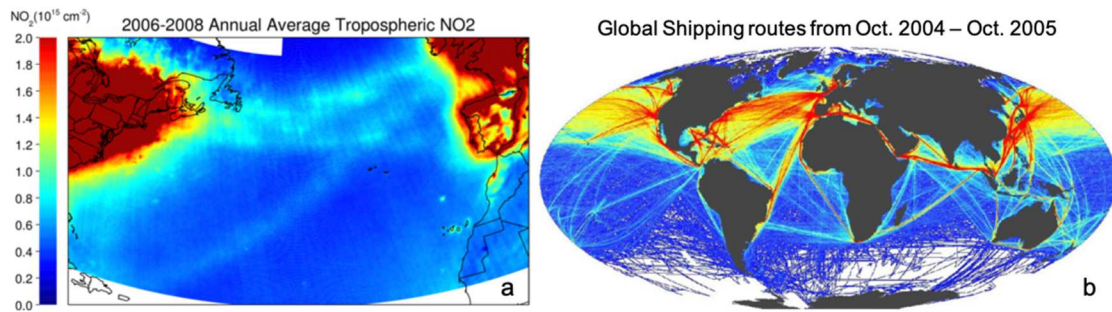


Figure 12: (a) Annually averaged tropospheric NO_2 from the DOMINO retrieval for 2006-2008. OMI pixels are filtered the same as in *Vinken et al.* [2014] (b) Ship traffic estimated from Voluntary Observing Ships Scheme data collected from October 2004 – October 2005. This figure is found in the supplementary material of [*Halpern et al.*, 2008].

Further analysis of the OMI retrieval reveals the features attributed to shipping lanes in Figure 12a may actually be artifacts in the data as a result of the model used in the retrieval. Based on personal correspondence with K. Folkert Boersma from the Royal Netherlands Meteorological Institute (KNMI) at the 2017 OMI Science Team Meeting in Greenbelt, MD, we learned the elevated NO_2 in the satellite observations across the Atlantic is due to the ship emissions used in the TM4 model to derive the *a priori* NO_2 profiles. The TM4 model is a 2° latitude \times 3° longitude global Chemistry-Transport Model (CTM) with 35 vertical levels that calculates the *a priori* NO_2 profiles used in the DOMINO retrieval algorithm [*Boersma et al.*, 2007]. Anthropogenic and biogenic NO_x emissions are based on values from the Precursors of Ozone and their Effect on the Troposphere (POET) European Union project for the year 1997 [*Boersma*

et al., 2007; *Olivier et al.*, 2003]. The erroneous signal across the Atlantic Ocean is likely due to this dated emissions inventory and coarse resolution of the TM4 model.

For the TROPOspheric Monitoring Instrument (TROPOMI) (<http://www.tropomi.nl/>) the update to OMI launched on October 13, 2017, the retrieval algorithm is based on the updated TM5-MP model which has a resolution of 1° latitude \times 1° longitude and uses the Monitoring Atmospheric Composition and Climate (MACC) emissions inventory to calculate the NO_2 *a priori* profiles [*Granier et al.*, 2011; *van Geffen et al.*, 2018; *Williams et al.*, 2017]. Analysis with TROPOMI NO_2 data will likely show a more realistic NO_2 signature across the Atlantic assuming the updated emissions inventory and TM5-MP model resolve modeling artifacts in the data. Analysis has not been completed to examine this possibility.

A recent study by *Ding et al.* [2018] used an inverse modeling algorithm called the Daily Emissions estimates Constrained from Satellite Observations (DECSO) [*Ding et al.*, 2017] and OMI observations between 2007 and 2016 to derive monthly shipping emissions estimates off the China coast. DECSO has a spatial resolution of $0.25^\circ \times 0.25^\circ$ and calculates the β term in Equation (30) (the sensitivity of column NO_2 to shipping emissions) [*Ding et al.*, 2017]. The transport of NO_2 across the model domain is calculated from an ensemble of 150 isotropic 2-D trajectories using the European Center for Medium-Range Weather Forecasts (ECMWF) forecasts. Trajectory calculations are performed for each model grid cell [*Ding et al.*, 2017]. DECSO assumes forecasted NO_2 emissions to be equal to emissions from the previous day, therefore updates to the emissions from satellite tropospheric column NO_2 measurements that vary from the mean by more than 3σ are considered incorrect and

a total adjustment of 3σ is used for the model grid cell [Ding *et al.*, 2017]. The trajectory analysis coupled with the Automated Identification System (AIS) data to monitor ship tracks of individual ships along the eastern China coast allow the DECSO algorithm to filter out transported NO₂ and reveal shipping lanes [Ding *et al.*, 2018]. In the future, application of the DECSO algorithm along with TROPOMI satellite data to the eastern US Coast is likely to improve monthly CMV inventory estimates, yielding similar results to Ding *et al.* [2018].

2.2.2: Using HYSPLIT

Our goal is to develop a top-down CMV inventory to compare with the EPA NEI inventory. This requires a discernable NO₂ signature in the 200 nautical miles off the US coast, known as the Exclusive Economic Zone (EEZ) [EPA, 2015a], and use of satellite data from the year 2011. Based on Figure 12a, contamination of the NO₂ column due to continental outflow must be addressed for the East Coast of the US due to the prevailing westerly winds. We use the HYbrid Single-Particle Lagrangian Integrated Trajectory (HYSPLIT) model [Stein *et al.*, 2015] developed by the Air Resources Laboratory at NOAA to find days during the O₃ season in 2011 that indicate Easterly wind flow. These days are assumed to have reduced anthropogenic pollution in the atmospheric column over the coastal Atlantic Ocean, allowing the NO₂ from shipping activity to be detected.

The ensemble back-trajectories end at 13:30 LST, approximately coincident with the local OMI overpass time, providing information about the origin of the air in the atmospheric column observed by OMI. All trajectories are run for three days using the NAM 12 km composite forecasts aggregated to 1-hour daily files for the CONUS

meteorology and terminate when rain is detected in the individual trajectory. Additionally, all trajectories are set to end at 50% of the model's mixed layer depth so ending trajectory heights vary from day to day. Figure 13 shows examples of the typical westerly flow (a and b), and the desired easterly flow (c and d) over the course of the O₃ season.

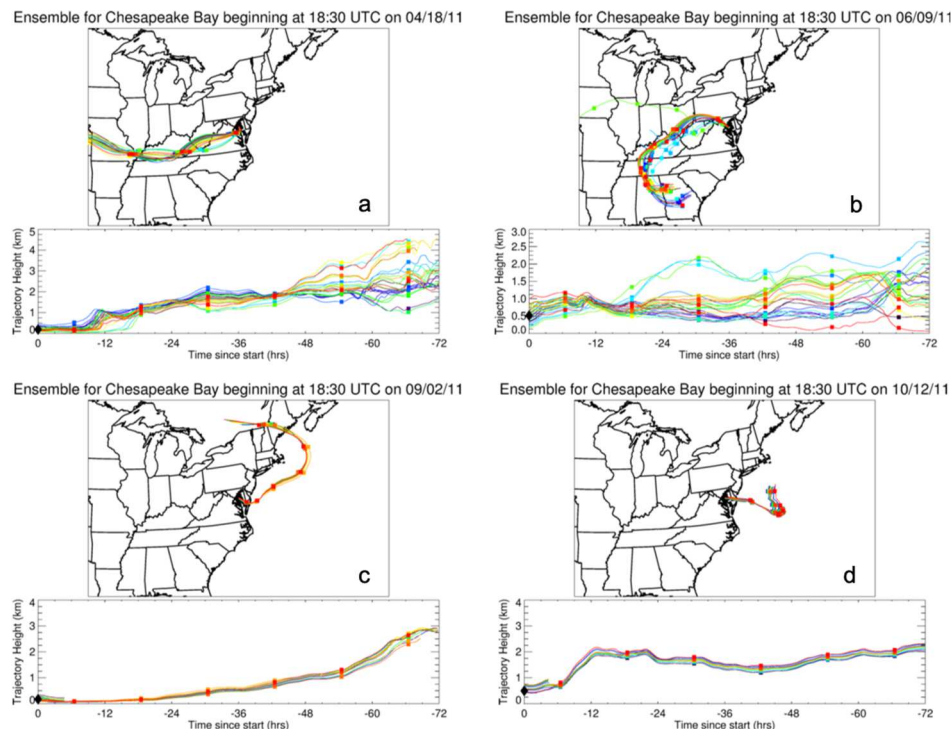


Figure 13: Ensemble 3-day back-trajectories performed from the Chesapeake Bay using meteorology from the North American Model. All trajectories terminate if rain is encountered. All trajectories begin at 13:30 LT, which coincides with OMI-overpass time. Panels (a) and (b) are examples of typical westerly atmospheric flow for the region. Panels (c) and (d) are examples of Easterly flow.

Using this filtering method, we isolate eleven days in the 2011 O₃ season that satisfy the criteria of having easterly flow and minimal influence from the US continent. In Figure 14, panel a shows the tropospheric column NO₂ averaged for the O₃ season, and panel b shows the average of the identified easterly flow days only. The striping feature in both panels results from the row anomaly in the OMI data discussed

previously in Section 1.4.2 and due to the removal of the outer most 15 pixels (Figure 9). Continental outflow is again shown off the East Coast of the US (Figure 14a) for the 2011 O₃ season, a persistent feature in the retrieved tropospheric column NO₂. In addition to the imposed filtering criteria of CRF < 30%, albedo < 30%, and no quality flags, we now are using data from just a handful of days in the study period. While the average shown in panel b reduces the continental NO₂ outflow distance from shore, we find this method to be unreliable to derive a top-down shipping emissions inventory for 2011. It is possible that this method could be useful if applied to earlier years, before the OMI row anomaly effected almost half of the pixels, however we are interested in developing a top-down inventory for 2011, and therefore did not pursue this prospect. Methodology suggested by the recent paper by *Ding et al.* [2018] should also be explored for application to the US east coast.

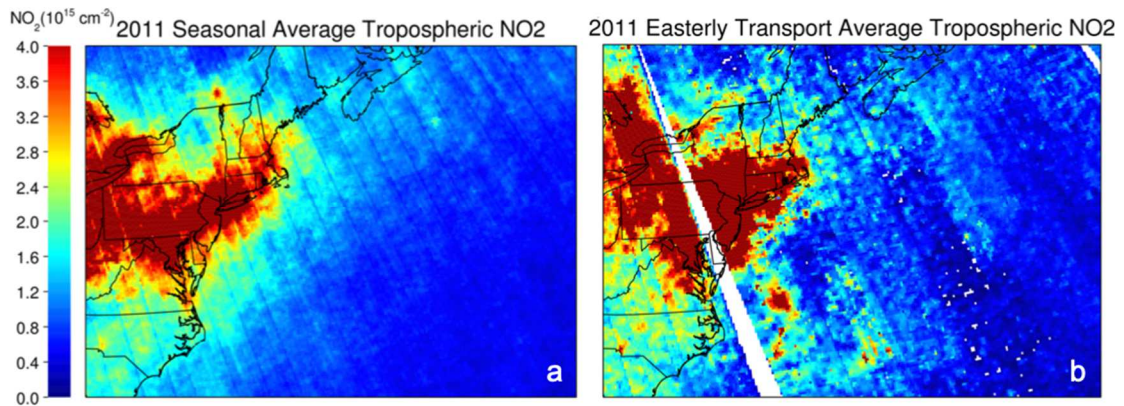


Figure 14: DOMINO tropospheric column NO₂ (a) averaged from April – October 2011 and (b) averaged easterly transport days only. Only pixels without a quality flag, and pixels with a Cloud Radiance Fraction (CRF) < 30% and an albedo < 30% are used in the average.

2.2.3: Issues with the Commercial Marine Vessel Inventory

The method to derive a top-down emissions inventory for the CMV sector described in *Vinken et al.* [2014] requires an atmospheric chemical model to provide

the *a priori* profile used in the adjusted satellite retrieval. This requires the model, and its emissions inputs, to be accurately represented in the model framework. In addition to the difficulties with the satellite data described in Section 2.2.2, we found that the CMV inventory is not represented accurately in the CMAQ model framework. Due to the increasing number of limitations to developing a reliable top-down emissions inventory for CMVs, we decided to suspend this research direction and pursue improving representation of the CMV emissions inventory in the CMAQ model framework. This effort is discussed in the next chapter.

Chapter 3: Emissions from Commercial Marine Vessels and their Role in Air Quality Policy (Published as Ring et al., 2018)

3.1: Introduction

Many studies have shown that elevated levels of pollution exist around heavily trafficked waterways and harbors [Cooper, 2003; Goldberg et al., 2014; Lawrence and Crutzen, 1999; Murphy et al., 2009; Pirjola et al., 2014; Williams et al., 2009], an important air pollution source for coastal states and states with shipping ports. Additionally, higher levels of criteria pollutants are observed at Air Quality System (AQS) monitoring sites near these coastal regions [Gégo et al., 2007; Mazzuca et al., 2016; Stauffer et al., 2015; Yu et al., 2006] than areas more inland.

In our current globalized society, it is estimated that approximately 90% of all traded goods are transported via CMVs [Pirjola et al., 2014; UN, 2017], with emissions from CMVs contributing to approximately 15-30% of the total global anthropogenic NO_x budget [Corbett et al., 2007; Eyring et al., 2005; Williams et al., 2009]. Due to the immense number of ships used to transport goods, and the international identity of CMVs, NO_x regulation and enforcement are difficult, even when ships are operating in near-shore shipping lanes and harbors [Eyring et al., 2005; Pirjola et al., 2014].

To manage the growing global use of CMVs, the United Nations (UN) proposed creation of the Inter-Governmental Maritime Consultative Organization (IMCO) in 1948, which was then officially established almost a decade later. This organization was charged with developing international guidelines to optimize maritime safety,

navigational efficiency, and to ensure all ships regardless of country flag registration can partake in international trade [IMO, 1948]. The original charter had no mention of controlling pollution due to shipping activity, but after the Torrey Canyon oil spill of 1967 [IMO, 1973; Wells, 2017], Article I of the original establishing document was amended to include specific reference to “prevention and control of marine pollution due to ships” [IMO, 1973]. Additionally, the name was eventually changed to the International Maritime Organization (IMO), to emphasize the organizations ability to act and enforce its regulations [IMO, 1973; 1975]. Beyond marine pollution, the IMO also realized the importance of regulating atmospheric pollution due to the ever-increasing international shipping activity. As a result, the International Convention for the Prevention of Pollution from Ships (MARPOL) was created [IMO, 1973]. To regulate marine pollution near land, the IMO has instituted controls for diesel engine vessels in specified Emission Control Areas (ECA); zones that extend 200 nautical miles off the coast of participating countries in their sovereign waters [EPA, 2008; IMO, 1998]. The US petitioned the IMO to include the North American continent in MARPOL Annex VI, which regulates air pollution from large ocean-going vessels, allowing the US and Canada to regulate CMV emissions within ECAs [EPA, 2008; 2009b]. The US was successfully added to the list of IMO participating countries in 2008, requiring all Class 3 Commercial Marine Vessels (c3 Marine) operating within US coastal waters to comply with IMO regulations [EPA, 2008].

Vessel age is important when determining emissions regulations. As of 2015, approximately 50% of the c3 Marine global fleet is over 20 years old, approximately 20% is between 10-20 years old, and approximately 30% is less than 10 years old [UN,

2015]. This means that ships younger than 20 years old (about half of the global fleet) are required to meet the Tier I IMO emissions regulations which require emission rates to be no larger than 17 g/kWh of NO_x when at idle [IMO, 2014]. As more of the global fleet is retired, new vessels must meet the more stringent global Tier II regulations (14.4 g/kWh of NO_x at idle) and Tier III regulations (3.4 g/kWh of NO_x at idle) if operating within ECAs [IMO, 2014].

The main engines of most commercial marine vessels are run solely for propulsion while auxiliary engines are run continuously to meet all other energy demands for ship operation. NO_x emissions rates from these two engine types are highly dependent upon fuel composition and engine temperatures. The slower the ships are moving, the longer and hotter these engines are running, resulting in higher NO_x emissions [Cooper, 2003]. CMVs are also a significant source of large particles called giant cloud condensation nuclei, contributing to enhanced boundary layer cloud formation [Sorooshian *et al.*, 2015] as well as particulate matter with an aerodynamic diameter less than 2.5 µm (PM_{2.5}), a criteria pollutant that contributes to hundreds of thousands of premature deaths globally [Cohan and Chen, 2014; Cohen *et al.*, 2005; Corbett and Koehler, 2003; Pope *et al.*, 2002].

Elevated levels of tropospheric O₃ over CMV traveled bodies of water have been measured by both in situ and remote techniques [Cleary *et al.*, 2015; Goldberg *et al.*, 2014]. Advection of this polluted air over coastal land and cities may contribute to the elevated pollution over these regions [Loughner *et al.*, 2011; Loughner *et al.*, 2014; Mazzuca *et al.*, 2017; Stauffer *et al.*, 2015] and at coastal AQS monitoring sites. For major metropolitan cities near heavily trafficked CMV areas like Baltimore, MD,

Philadelphia, PA, and New York, NY, the development of attainment strategies addressing CMV emissions will become increasingly important as global shipping activity is projected to increase in the future [EPA, 2009b; McDill *et al.*, 2015].

In this Chapter, we discuss the role of CMV emissions on regional air quality, and the representation of this pollution source within the CMAQ model [Ring *et al.*, 2018]. We review the necessary adjustments made to the vertical distribution of emissions from the largest (class 3) CMVs and examine the effect of this adjustment on surface O₃ production for various model simulations conducted for years 2011 and 2018. Comparisons of modeled surface O₃ to AQS data, and modeled column formaldehyde (HCHO) and NO₂ to satellite measurements, are used to evaluate model performance for 2011. Finally, we quantify the effect of an improved model framework for a 2018 SIP attainment strategy developed by the Maryland Department of the Environment (MDE) [Ring *et al.*, 2018].

3.2: Commercial Marine Vessel (CMV) Emissions Adjustment

3.2.1: Using the Community Multiscale Air Quality (CMAQ) Model

For this analysis, we use an EPA approved regulatory air quality model used by state and federal agencies to develop surface O₃ attainment strategies called CMAQ (described in Section 1.5.1). The model domain is represented by the outlining black box in Figure 15.

All anthropogenic emissions used for this study are the MARAMA Alpha 2 version described in Section 1.5.3. Various sector emissions are gridded to CMAQ

resolution and merged into daily 3-D, temporally varying input files using SMOKE version 3.6 [CMAS, 2014b] described in Section 1.3.1. Model output from CMAQ is generated for June 1st through August 31st for 2011 and 2018, the three hottest and most important months for photochemical O₃ production during the O₃ season (April-October). We compare CMAQ output from the 2011 base case simulation to observations for the same period. The 2018 CMAQ output is used to examine the effectiveness of future emissions reductions and proposed policies, discussed later in Chapter 3 [Ring *et al.*, 2018].

3.2.2: CMV Emissions Inventory Development

This study focuses on the Class 3 Commercial Marine Vessel (c3 Marine) inventory, which represents emissions for marine diesel engines with fuel displacement of 30 liters/cylinder or larger [EPA, 2002; 2009b; 2015a]. This sector mainly consists of cruise ships and international open ocean vessels, used in transporting consumer goods and resources.

The EPA generates the total c3 Marine emissions values for the NEI2011 by applying growth factors to ship emissions data acquired in 2002 [EPA, 2002; 2009b]. These growth factors account for a variety of policies and changes in shipping activity that affect the 2011 inventory year, including the ECA regulations described earlier and global NO_x and sulfur dioxide (SO₂) controls [EPA, 2015a; McDill *et al.*, 2015]. The spatial distribution of CMV emissions are generated using data from the Research and Innovative Technology Administration's Bureau of Transportation Statistics National Transportation Atlas Database, the US Census Bureau, GIS shapefiles provided by ports, and satellite imagery from sources such as Google Earth [EPA, 2009a; ERG,

2010]. This method was refined for developing the 2014 NEI, improving county designation and grid point locations within state waters [EPA, 2015b]. Similar studies have used data from AIS, which provides detailed information about ship location, speed, ship activities, navigational plans, and many other parameters [Chen *et al.*, 2017a; Chen *et al.*, 2017b; Chen *et al.*, 2016]. Due to the large data volume, computational intensity, complexity and scale of US coastal operations, AIS was not used to develop 2011 or 2014 CMV emissions in the NEIs, however it is recommended by the EPA along with State and Local agencies for future inventory development [EPA, 2015b].

For the c3 Marine sector, the emissions inventory is divided between off-shore and near-shore emissions. The EPA is responsible for the development of off-shore c3 Marine emissions, designated as outside state waters but within the country's established Exclusive Economic Zone (EEZ) and therefore the ECA. Conversely, near-shore emissions are designated to be within county (and therefore state) waters extending approximately 3 miles off the coast, and thus the EPA requests that states gather and report this information [EPA, 2015a; McDill *et al.*, 2015]. If near-shore emissions from states are incomplete, EPA generated emissions are used to supplement.

Figure 15 shows the spatial extent of near-shore (a) and off-shore (b) emissions within our CMAQ modeling domain. The off-shore emissions are designated as point sources, allowing for the vertical distribution of emissions above the surface, reflecting the stack heights of the ships, along with plume rise and dispersion of the emissions. Near-shore c3 Marine emissions are designated as area sources, so all emissions occur only in the surface layer of the model. This different designation results in inconsistent

modeling of c3 Marine emissions and erroneously places high levels of NO_x and other c3 Marine pollution from the near-shore emissions files at the surface of heavily trafficked, near coastal waterways like the Chesapeake and Delaware Bays, New York Harbor, and Long Island Sound. It also prevents CMAQ from accurately simulating vertical mixing and transport of these emissions because they are emitted too close to the surface.

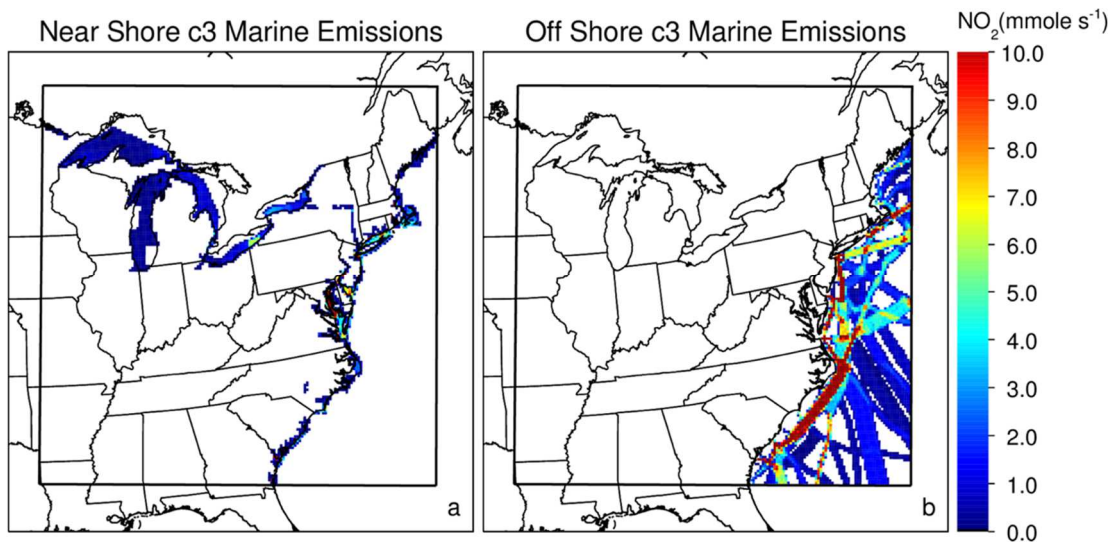
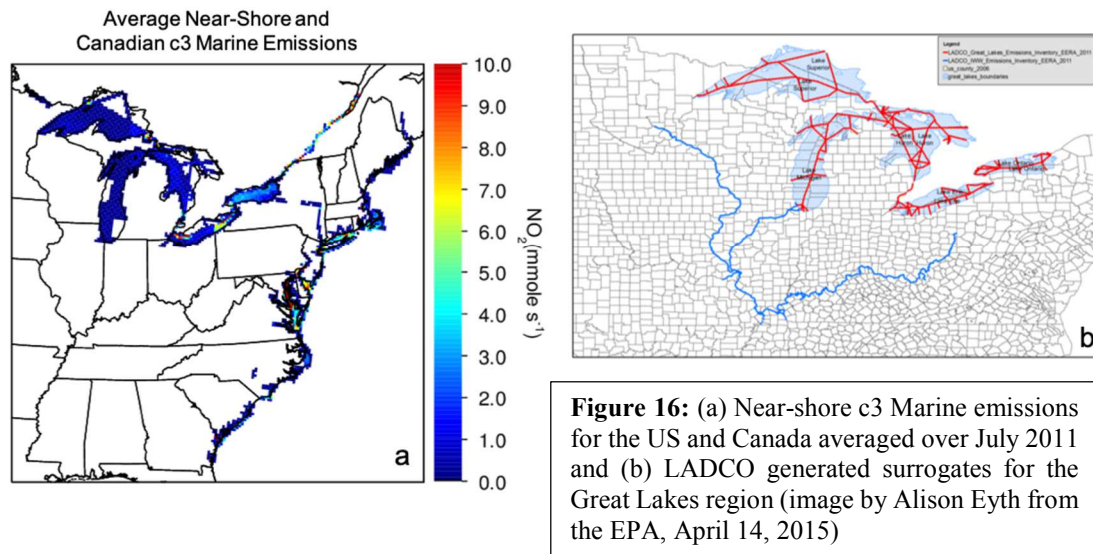


Figure 15: Average class 3 commercial marine vessel (c3 Marine) NO₂ inventory emissions at 2pm LST for June, July, and August 2011 are shown for (a) near-shore and (b) off-shore emissions files. Black boxes in (a) and (b) outline the CMAQ model domain.

Additionally, in the Great Lakes region the near shore emissions are spread out across the surface of the lakes in US waters. This is because the area source designation distributes the emissions out across the county boundaries that extend over water, stopping at the US-Canada border as shown in Figure 16. Comparing the US near-shore and Canadian near-shore c3 Marine emissions, it is evident that shipping lane surrogates were used to develop the Canadian emissions but were not used, or were used incorrectly, to develop the US emissions shown in Figure 16a. After further

research, we found surrogates were developed for commercial marine traffic in the Great Lakes (Figure 16b), however they were not present in the EPA emissions inventory ftp database: <ftp://ftp.epa.gov/EmisInventory/2011v6/v2platform/training>.



We believe the white geometric areas in Lake Michigan (Figure 16a) representing no near-shore emissions indicate that surrogates did influence inventory development but were not used correctly. As shown in Figure 16b, any county that has a shipping lane extend through it has near-shore area source emissions attributed to the entire county. If the ship tracks do not extend through the county, no emissions are designated in those areas, which explains the strange shape of the missing near-shore emissions. The issues of missing surrogate files in the EPA inventory ftp database and the clear difference between the Canadian and US near-shore c3 Marine emissions distribution were communicated to the appropriate people at the EPA. These should also be rectified in future inventories.

3.2.3: Class 3 Commercial Marine Vessel Emissions Adjustment

As discussed in the previous section, near-shore c3 Marine emissions occur within the surface layer of the model while off-shore c3 Marine emissions are vertically distributed in model layers above the surface. We attempt to rectify this discrepancy by using an appropriate average profile calculated based on nearby, off-shore c3 Marine emissions, and apply this profile to the near-shore c3 Marine emissions.

First, we must determine how the near-shore and off-shore emissions interface with each other in order to adjust the near-shore emissions most accurately. Figure 17 shows how the near-shore and off-shore emissions meet along the entire Eastern US coast. In all 7 panels (a-g) of Figure 17, overlap between the near-shore emissions (blue circles) and off-shore emissions (red squares) are identified and highlighted by yellow circles. This overlap is an oversight in the emissions development process and must be rectified. Further analysis shows that the emissions from the near-shore and off-shore files for the overlapping points are not identical, meaning the sources for the emissions estimates at these grid points are not the same. This supports the notion that this mistake likely results from combining input from SLT agencies with EPA emissions estimates.

Previous modeling performed with unadjusted emissions files therefore used c3 Marine emissions that were too high at these overlapping grid points, allocating a gross overestimation of emissions within the surface layer of the model. To correct this issue, we removed all overlapping grid points from the near-shore c3 Marine inventory file. This might not be the best method, since the EPA generally defers to SLT inputs, however, it is these SLT emissions that are erroneously allocated to the surface layer of the model and need further adjustment. We elected to keep the off-shore emissions

wherever possible because they required no modifications, and therefore match most closely to the government provided emissions [*Ring et al.*, 2018].

In addition to accentuating the overlapping emissions, Figure 17 shows the spatial extent of the near-shore emissions. In all 7 panels, it is evident that the emissions extend well over land, beyond where any large tanker or shipping vessel could travel. It is likely some of these grid points contained emissions that were incorrectly designated as c3 Marine, however further investigation is needed to determine how best to correct this problem. With the resources available, it was not feasible to adjust the horizontal spatial distribution of the near-shore emissions, but this issue was made apparent to the inventory development team at the EPA and modifications should be made for future emissions inventories.

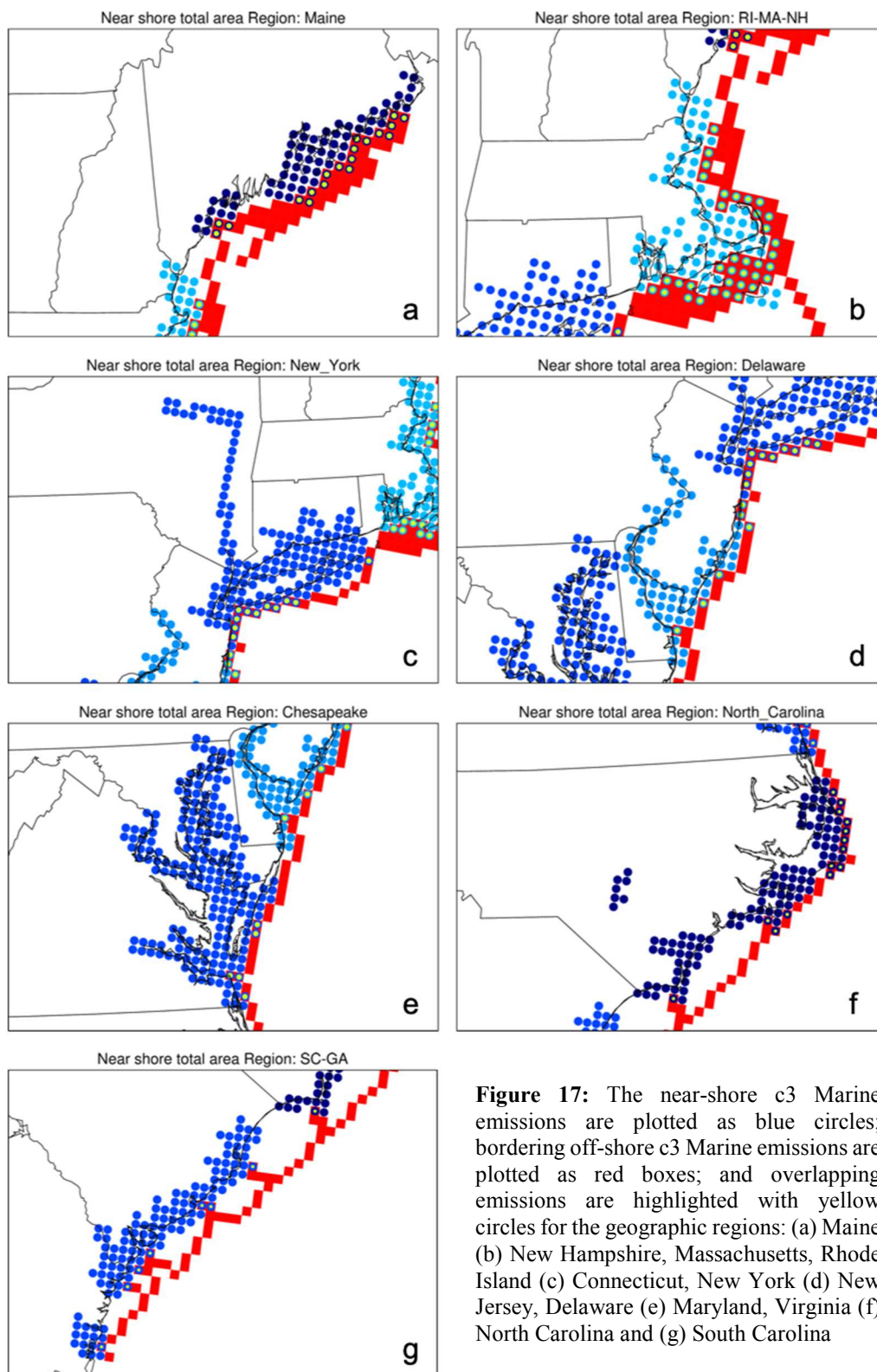


Figure 17: The near-shore c3 Marine emissions are plotted as blue circles; bordering off-shore c3 Marine emissions are plotted as red boxes; and overlapping emissions are highlighted with yellow circles for the geographic regions: (a) Maine (b) New Hampshire, Massachusetts, Rhode Island (c) Connecticut, New York (d) New Jersey, Delaware (e) Maryland, Virginia (f) North Carolina and (g) South Carolina

After removing overlapping emissions, we had to determine how best to adjust the near-shore emissions to reflect the vertical distribution. Since the goal of this work is to represent the entire c3 Marine inventory sector (both off-shore and near-shore emissions) in a consistent manner, determining how the off-shore inventory files vary from day to day, hour by hour, is crucial to creating a sound method for adjusting the near-shore file. To determine if there were temporal and geographic variations in the off-shore emissions vertical distribution, we chose test grid points near large shipping ports to analyze (Figure 18). These test points highlight areas that are highly trafficked by class 3 vessels and represent off-shore emissions that border the near-shore c3 Marine emissions.

To create a formula for adjusting the near-shore inventory file, we needed to determine the percent of the total emissions in each layer for the off-shore c3 Marine inventory file. For example, in Figure 19, we see that for all four test locations, the largest percent of the emissions are placed in layer 3 of the model for all hours (royal blue circles in panels a-d). Throughout the day, anywhere between 60-80% of the emissions are put into layer 3, with varying amounts in the other layers depending on geographic location. Based on this finding, it is evident that preserving geographic and temporal variation found in the off-shore emissions is necessary to adjust the near-shore emissions in a sound manner.

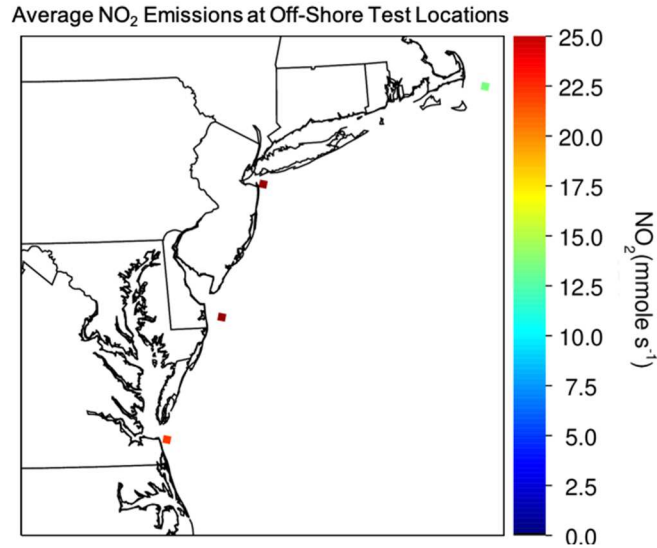


Figure 18: Off-Shore c3 Marine grid points chosen as test locations to determine temporal variation of emissions. The four grid points are colored by the average NO_2 emissions in July 2011.

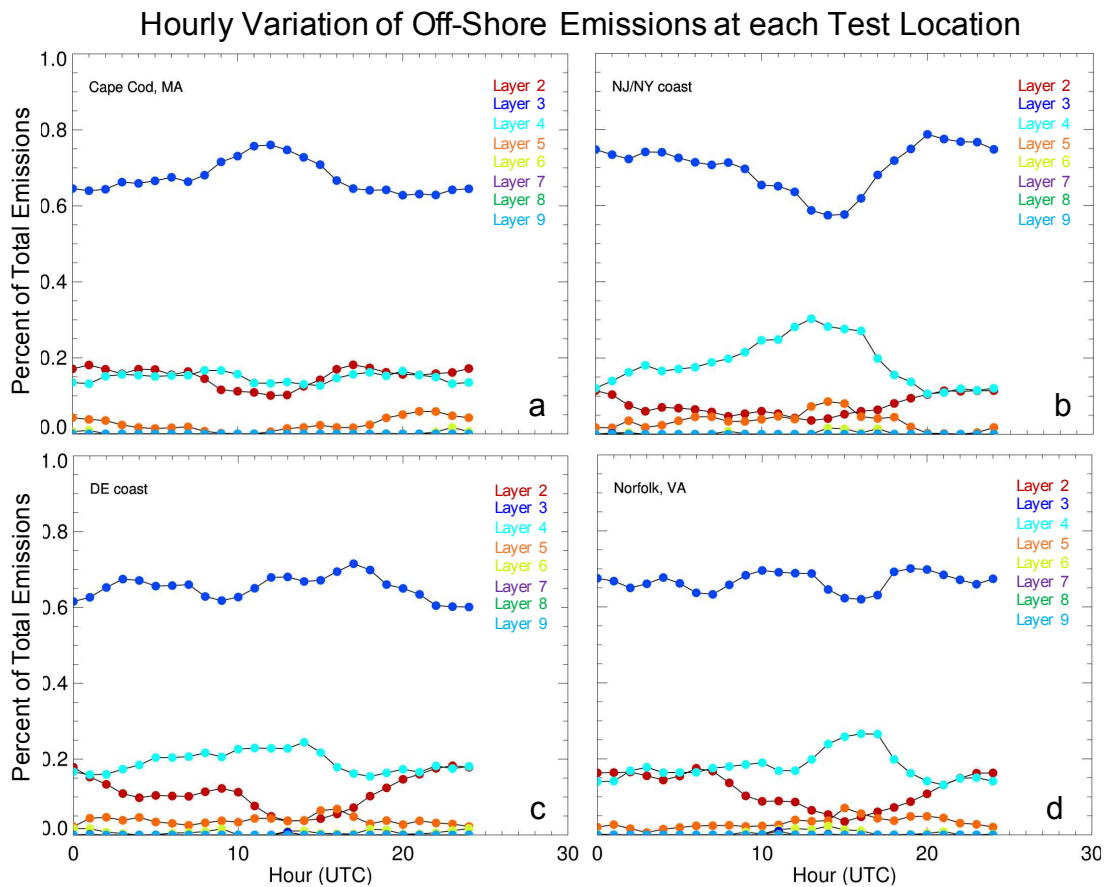


Figure 19: Hourly variation in the percent of total emissions that are present in each model layer for test locations (a) Cape Cod, MA, (b) NJ/NY coast, (c) DE coast, (d) Norfolk, VA. The exact geographic locations of the test grid points are shown in Figure 18.

After determining geographic variation, it became evident that regional division of the near-shore files is necessary. As shown in Figure 17 (a-g), natural breaks in the near-shore c3 Marine inventory created 7 logical adjustment regions. Figure 20a shows how the near-shore c3 Marine emissions are divided into 7 adjustment regions. Since the majority of the near-shore inventory grid points do not directly border off-shore inventory grid points, we create an average distribution profile for each adjustment region. This average profile is created from all off-shore grid points that border the specified adjustment region. An example is shown as gray boxes in Figure 20b for the MD/VA adjustment region [Ring *et al.*, 2018].

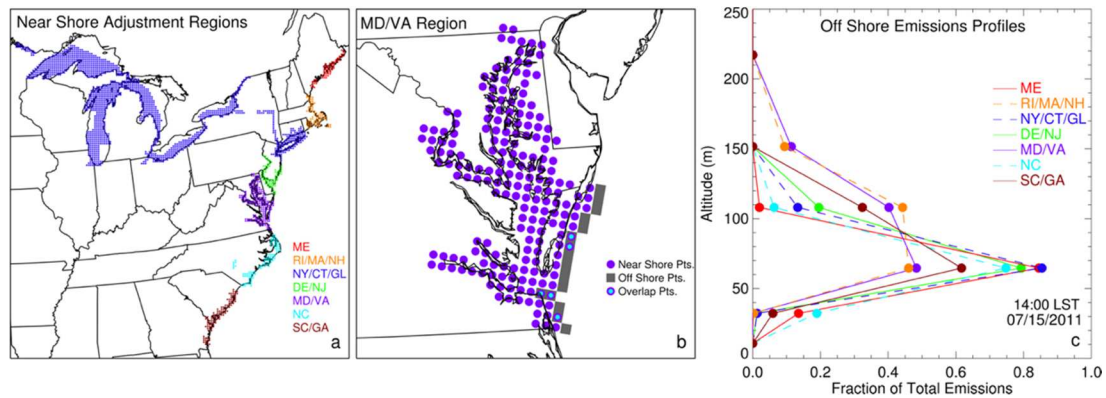


Figure 20: (a) Near-shore c3 Marine emissions geographically divided into 7 adjustment regions. All near-shore c3 Marine emissions grid points are color coded to show regional designation. (b) The MD/VA region near-shore c3 Marine emissions grid points (purple filled circles), off-shore emissions grid points (gray squares), and overlapping grid points (purple and blue circles). (c) Average vertical fractional distribution of off-shore c3 Marine emissions for each region at 14:00 LST on July 15, 2011 plotted at the half-height of the emissions layer.

Importantly, the Great Lakes near-shore emissions do not border any off-shore grid points, therefore, we attribute the profile for the NY/CT region, the adjustment region that most closely connects, to the Great Lakes (Figure 20a).

The calculated average profile for each adjustment region was then used to vertically distribute all near-shore emissions within the corresponding adjustment

region. Due to the temporal variation shown in Figure 19, the average profile was calculated for each hour of each day and applied to the appropriate near-shore region. As an example, average vertical distribution profiles calculated for each adjustment region at 14:00 EST on July 15, 2011 are shown in Figure 20c [Ring *et al.*, 2018].

Finally, all adjusted near-shore c3 Marine emissions were removed from the daily area source near-shore file and added to the corresponding daily point source offshore file to create a single c3 Marine emissions inventory file. This ensures all c3 Marine emissions data are modeled consistently within CMAQ, as point source emissions, with appropriate vertical distribution [Ring *et al.*, 2018].

3.3: Modeling Impacts for 2011

To analyze the impact of this adjustment to the c3 Marine inventory sector on air quality, CMAQ simulations are performed. In addition to baseline and adjusted c3 Marine model simulations, model runs that incorporate observationally driven adjustments to the on-road mobile emissions inventory [Anderson *et al.*, 2014; Travis *et al.*, 2016], and modifications to the CB05 chemical mechanism [Canty *et al.*, 2015] are completed.

The adjustment to on-road mobile NO_x emissions throughout the modeling domain addresses an overestimation of the NO_x emitted from car tail pipes. Analysis of aircraft observations in the Baltimore-Washington area collected during the Deriving Information on Surface Conditions from Column and Vertically Resolved Observations Relevant to Air Quality (DISCOVER-AQ) 2011 campaign indicate emissions from this sector are high by a factor of two in the EPA inventory [Anderson

et al., 2014; *Travis et al.*, 2016]. While this conclusion from *Anderson et al.* [2014] was originally challenged by the EPA, subsequent studies have confirmed the results for other areas across the US, providing further evidence for reducing NO_x emissions from cars. Consequently, we apply a 50% reduction to on-road mobile NO_x emissions throughout the entire domain. The next adjustment corrects the expectation that the alkyl-nitrate radical group (NTR) in CB05, has a lifetime of ~10 days for loss by photolysis. For simplicity, CB05 groups species that react similarly in the atmosphere. This can lead to over simplifications of important chemical reactions, as in this case. The 10-day lifetime for the NTR group is much longer than the actual photolytic lifetime of most of the chemical species comprising the NTR group, so we reduce the lifetime by a factor of 10 to a more realistic ~1-day lifetime [*Canty et al.*, 2015]. It is important to note that CB05 assumes 100% recycling of NO_x, which contributes to more O₃ production within the model than is realistic [*Canty et al.*, 2015]. The final adjustment made to the various modeling simulations is the change to the vertical distribution of c3 Marine emissions explained in section the previous section. Table 2 details the various model and emissions inventory adjustments made for each modeling scenario discussed in this paper.

Table 2: Description of the four CMAQ modeling scenarios and their appropriate inventory and model adjustments.

MODELING SCENARIO	C3 MARINE EMISSIONS ADJUSTMENT	ON-ROAD MOBILE EMISSIONS INVENTORY ADJUSTMENT	ALKYL NITRATE RADICAL (NTR) LIFETIME ADJUSTMENT
BASELINE	No adjustment	No adjustment	No adjustment
C3 ADJUST	Adjusted emissions	No adjustment	No adjustment
SCIENCE	No adjustment	Adjusted emissions decrease NO _x by 50% (Anderson et al. 2014, Travis et al. 2016)	Reduced NTR lifetime from 10 days to 1 day (Canty et al. 2015)
C3 SCIENCE	Adjusted emissions	Adjusted emissions decrease NO _x by 50% (Anderson et al. 2014, Travis et al. 2016)	Reduced NTR lifetime from 10 days to 1 day (Canty et al. 2015)

We use average maximum 8-hr O₃ (AM8O₃) as a metric to analyze model output and compare with ground-based observations. We follow the EPA guidelines for calculating maximum daily average 8-hr O₃ (MDA8) [EPA, 2015c], which are summarized as follows. A forward running 8-hr mean is calculated for each day at each grid cell within the model. The MDA8 within a 3 × 3 grid is then attributed to the center grid cell. For the June, July, August (JJA) time frame, we average the MDA8 for the top 10 days above 60 ppb at each grid cell. This is known as the AM8O₃ value. To calculate AM8O₃ for corresponding simulations of future air quality, in this case 2018, the MDA8 assigned to the center grid cell of the 3 × 3 grid in 2018 is co-located with the MDA8 cell used in the 3 × 3 grid in 2011. This method highlights air quality on the worst days in the study period. Air quality policy is created using this metric, therefore, the focus of attainment strategies is to reduce poor air quality on the worst days.

For the rest of this analysis, the geographic area is narrowed from the full model domain to a region encompassing the coastal Mid-Atlantic and Northeast states, highlighting locations with large, highly active commercial ports near heavily populated areas with some of the highest levels of observed surface O_3 . The study region is outlined in subsequent figures by a black dashed line.

Figure 21a shows $AM8O_3$ for the baseline model simulation. As shown in Table 2, the 2011 Baseline simulation does not include any model adjustments. The 2011 c3 Adjust scenario (Figure 21b) incorporates the near-shore c3 Marine vertical distribution adjustments. The difference between these two simulations (Figure 21c) reveals the areas within our study domain that show considerable changes in surface $AM8O_3$ due to the improved vertical representation of c3 Marine emissions. The decrease in surface $AM8O_3$ over the Chesapeake Bay and closely surrounding areas is due to the c3 Marine adjustment moving O_3 precursor emissions off the surface and distributing them vertically within the atmosphere. This modification allows some emissions to be vented out of the boundary layer, improving model representation of atmospheric pollution transport, resulting in a significant (~ 6 ppb) decline of surface $AM8O_3$ over the Chesapeake Bay. The increase of surface $AM8O_3$ shown in the New York and Connecticut region is due to increased pollution transport from upwind c3 Marine sources, and the non-linear chemistry that controls O_3 production. This will be discussed further in the next section.

Comparisons of surface O_3 model output to surface O_3 data from AQS locations are shown for the 2011 Baseline (Figure 21d) and 2011 c3 Adjust (Figure 21e) modeling scenarios. Within our study region, we compare AQS $AM8O_3$ calculated

from observations to modeled AM8O₃ from the CMAQ grid point closest to the AQS site. Comparison of Figure 21 panels d and e show the effect of the c3 Marine emissions vertical distribution adjustment. Modeled AM8O₃ at some AQS locations increases, by as much as a factor of two, while at others it decreases, however, almost all the modeled AM8O₃ is too high compared to observations. For instance, the AQS site at Greenwich Point Park, CT has a measured AM8O₃ value of ~80 ppb and a modeled AM8O₃ value of ~150 ppb for both scenarios. A high bias in CMAQ model output of surface AM8O₃ has been shown in prior studies and determining the reasons for this high bias are areas of active research [Canty *et al.*, 2015; Loughner *et al.*, 2014; Trail *et al.*, 2014; Travis *et al.*, 2016; Vinciguerra *et al.*, 2017]. The comparisons shown in Figure 21d and e highlight the limitations of CMAQ to accurately model surface O₃ on specific days, under certain meteorological conditions, particularly for coastal regions.

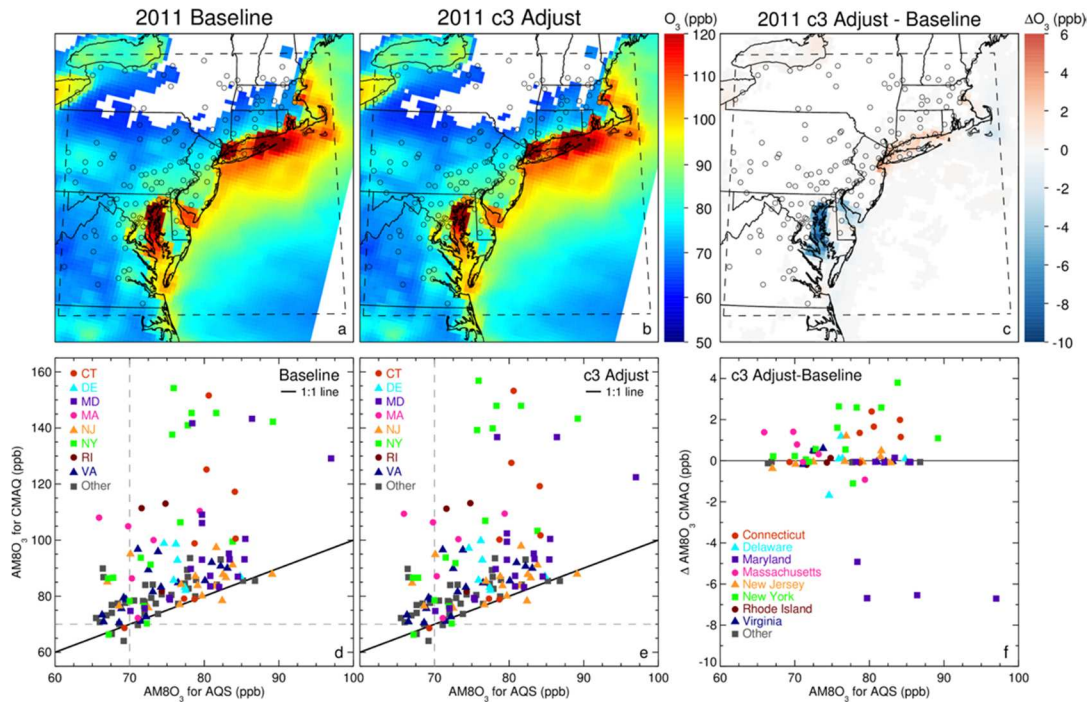


Figure 21: Average maximum 8-hr ozone (AM8O₃) for June, July, and August 2011 for (a) Baseline and (b) c3 Adjust model scenarios. If the AM8O₃ criteria are not met, model output is not shown (white regions within domain). Black dashed lines in upper plots outline the focus region for this study. Open

black circles represent AQS O₃ monitoring station locations. (c) The AM8O₃ difference between Baseline and c3 Adjust. (d) A scatter plot of AM8O₃ for Baseline vs. AQS data, with points color coded by state. (e) A scatter plot of AM8O₃ for c3 Adjust vs. AQS data. (f) The change in AM8O₃ between the c3 Adjust and Baseline model scenarios at AQS locations.

Figure 21f shows the change in modeled surface AM8O₃ at each AQS location when the vertical distribution of c3 Marine emissions is improved. Model values at AQS sites with differences less than 0.05 ppb are not included. We find the largest increases of modeled AM8O₃ at some AQS sites in New York and Connecticut, discussed in the next section, and we see the largest decreases at some AQS sites in Maryland. At the AQS site in Furley, MD, the model shows an AM8O₃ decrease of ~4.9 ppb, while the Edgewood, MD, Essex, MD, Calvert, MD, Blackwater NWR, MD sites all show decreases of over 6 ppb. All five AQS sites border the Chesapeake Bay, with Essex, Edgewood, and Furley located near Baltimore, MD. This reduction of modeled AM8O₃ in Maryland is of considerable magnitude, and an important result for state and federal agencies developing air quality policy [Ring *et al.*, 2018].

3.4: Evaluation of CMAQ Output with Satellite Retrievals

The photochemical production of O₃ is non-linearly dependent upon ambient NO_x and volatile organic compound (VOC) concentrations [Jacob, 2000; Kleinman, 2005; Kleinman *et al.*, 2001; Sillman, 1999; 2002]. This dependence leads to an optimal VOC:NO_x ratio that produces the maximum amount of tropospheric O₃ and represents the transition region between two atmospheric states: one where O₃ production is limited by the concentration of VOCs (VOC-sensitive) and the other where O₃ production is NO_x-sensitive. Understanding the state of the actual atmosphere, along

with how O₃ production is represented within regulatory air quality models, is crucial for developing attainment strategies that will properly inform air quality policy decisions.

If the local atmosphere is NO_x-sensitive, reducing NO_x emissions will have the desired policy effect of decreasing tropospheric O₃ concentrations. In a locally VOC-sensitive environment, reducing NO_x will have the undesired effect of increasing local concentrations of tropospheric O₃, until the reductions of NO_x are large enough to place O₃ production in the NO_x-sensitive regime. Generally, VOC-sensitive conditions are present in megacities and major metropolitan areas such as New York, NY, Houston, TX, and Los Angeles, CA [Duncan *et al.*, 2010; Kleinman, 1994; Kleinman *et al.*, 2000; Madronich, 2014; Mazzuca *et al.*, 2016], making air quality control especially challenging for these areas. In a VOC-sensitive environment, reductions in both VOCs and NO_x are typically necessary to improve surface O₃. Previous studies have shown that the Baltimore-Washington area has successfully transitioned to a NO_x-sensitive regime [Duncan *et al.*, 2010; Goldberg *et al.*, 2016; He *et al.*, 2013a; Liao *et al.*, 2014]. The behavior of surface O₃ in the CMAQ simulations presented in Figure 21c indicate NO_x-sensitive conditions for the Baltimore-Washington region because reductions of surface NO_x emissions, due to the adjustment of vertical profile of c3 Marine emissions, lead to a decrease in modeled surface O₃ (Figure 21c and f). Conversely, the surface O₃ increases shown in the New York City, Connecticut, and Long Island Sound region in Figure 21c are indicative of a region where O₃ production is VOC-sensitive, at least within the model.

Further analysis into the photochemical regime for O₃ production in the New York metropolitan area is performed using OMI satellite observations of column HCHO and NO₂. Daily retrievals of HCHO and NO₂ on a 0.25° latitude × 0.25° longitude grid were used to calculate average HCHO and NO₂ for June, July and August 2011.

We first calculate the standard deviations (σ) of the HCHO and NO₂ data at each grid point. We require there to be at least 10 coincident days of data for both HCHO and NO₂ at each grid point, that fall within two standard deviations (2σ). Finally, we calculate the average HCHO and NO₂ values at each grid point over the period [Ring *et al.*, 2018].

For comparison to CMAQ, we apply the averaging kernel (AK) from the retrievals to the model output for HCHO and NO₂. The AKs were calculated by dividing the Air Mass Factor (AMF) for HCHO and the tropospheric AMF for NO₂ by the scattering weight reported for each retrieval, as described in Gonzalez Abad *et al.* [2015]. Model output from 14:00 LST, was log-linearly interpolated to the satellite pressure level at each model grid point, and then multiplied by the corresponding AK from the closest overlapping satellite pixel. The resulting product was then integrated to calculate a column value for both HCHO and NO₂, comparable to the satellite column observations. These CMAQ-based columns were then averaged onto the same 0.25° latitude × 0.25° longitude grid as the satellite data, based on the criteria used to calculate the average satellite HCHO and NO₂ values, detailed above. This method ensures consistency between the satellite data and model output used in the analysis.

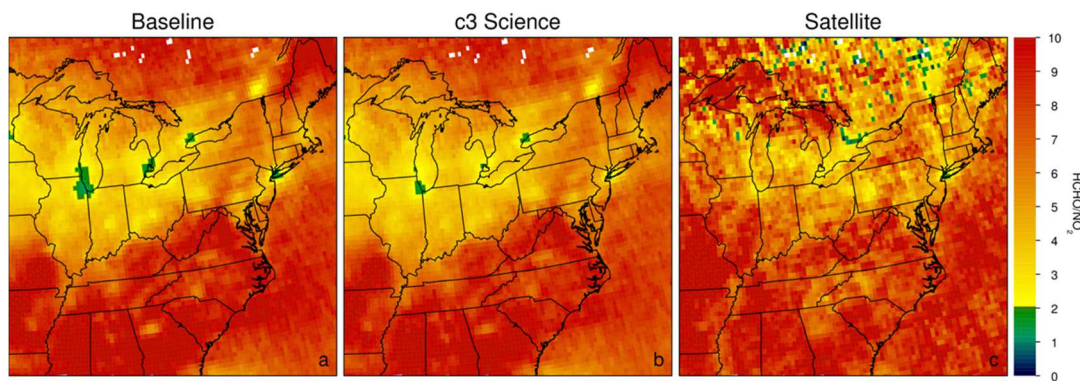


Figure 22: Ratio of average column HCHO to average tropospheric column NO_2 for (a) Baseline model, (b) c3 Science scenario, (c) OMI satellite data, over the model domain for JJA 2011. Only grid points with at least 10 days of satellite data for both HCHO and NO_2 retrievals are used. Model output for the two scenarios are convolved with appropriate OMI SAO (HCHO retrieval) and OMI GSFC (NO_2 retrieval) averaging kernels.

Figure 22 shows the average column HCHO/average tropospheric column NO_2 ratio (hereafter, HCHO/ NO_2 ratio) from CMAQ output (Figure 22a and b) and satellite data (Figure 22c) for the modeling domain. According to *Duncan et al.* [2010], a HCHO/ NO_2 ratio between 1 and 2 indicates the atmospheric column is transitioning between VOC-sensitive and NO_x -sensitive regimes. When the ratio is below 1, the atmosphere is considered to be VOC-sensitive, and when the ratio is above 2, the atmosphere is considered to be NO_x -sensitive. It is important to note that this analysis is for the *atmospheric column*, not the surface, therefore it may not be indicative of the surface photochemical O_3 production regime. This comparison does however provide insight into lower free troposphere composition, an important factor especially in the afternoon, for addressing air quality issues. Mixing of the Planetary Boundary Layer (PBL) with the free troposphere combines pollution transported from upwind sources with local pollution, changing the local atmospheric conditions. Understanding the upwind sources and composition of pollution are paramount to developing effective air quality regulations.

Figure 23 is the same as Figure 22, except it focuses on the New York metropolitan region. Figure 23a shows that for the baseline CMAQ model scenario, the air above New York City (NYC) and Eastern Long Island has a HCHO/NO₂ ratio between 1 and 2 (green boxes), indicating that the atmosphere is transitioning between the VOC and NO_x sensitive regimes [Duncan *et al.*, 2010]. In Figure 23b, the c3 Science model scenario (detailed in Table 2) shows the area downwind of NYC is mostly transitioned to a NO_x-sensitive atmosphere (yellow and orange boxes), leaving a smaller area over NYC within the transition zone. Figure 23c shows the average satellite HCHO/NO₂ ratio, indicating that some areas over NYC are within the transition zone while most of the area is within the NO_x-sensitive regime. Figure 23b shows too much of the NYC region as transitioned to NO_x-sensitive conditions, compared to the satellite ratio (Figure 23c). This is likely due to the 50% reduction of mobile NO_x emissions in the c3 Science model scenario. This reduction was applied domain wide, based on an empirical study that was completed for the Baltimore Washington region [Anderson *et al.*, 2014]. While the 50% reduction in mobile NO_x is a good approximation of the necessary emissions inventory adjustment, further modeling studies should adjust the mobile NO_x emissions inventory on a more localized scale, based on observational data throughout the model domain.

Figure 23d and e are scatter plots of the Baseline CMAQ HCHO/NO₂ ratio to the satellite HCHO/NO₂ ratio, and the c3 Science model HCHO/NO₂ ratio to the satellite HCHO/NO₂ ratio, respectively. Linear least squares fits to the data points, forced to go through the origin (0-intercept) are also shown. These comparisons

indicate that the c3 Science scenario is more closely representative of the observed atmosphere, as evidenced by the 0-intercept slope of 0.97 compared to the 0-intercept

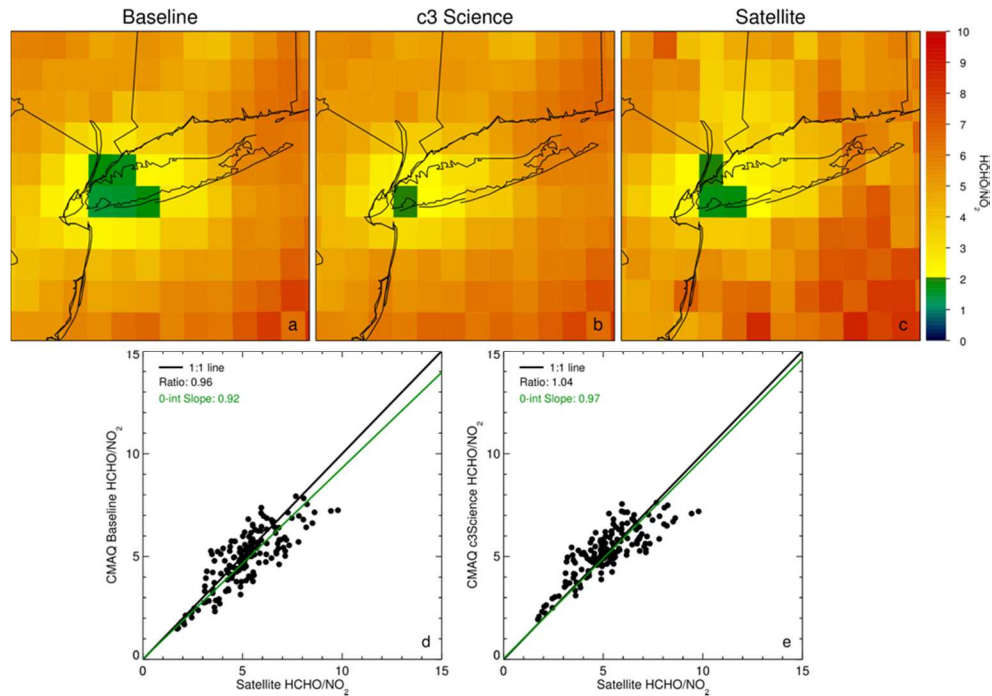


Figure 23: The top row is the same as Figure 22, except focusing on the NY metropolitan region. Scatter plots comparing the satellite derived HCHO/NO₂ ratio and the (d) Baseline scenario ratio, and (e) c3 Science scenario ratio are shown.

slope of 0.92 for the Baseline scenario. In Figure 23d and e, points below the 1 to 1 line indicate the model is producing a lower HCHO/NO₂ ratio than expected. This could mean there is too little HCHO, or too much NO₂ in the model column compared to the satellite retrieval. When the c3 Science model framework is used (Figure 23e), the number of points below the 1 to 1 line are reduced, especially for ratio values < 5. Since the model improvements mostly affect column NO_x concentrations, this suggests that the adjustments are improving the representation of NO_x within the model. Additionally, for the Chesapeake Bay region, both the baseline and c3 science model

frameworks are representative of observed satellite HCHO/NO₂ ratios (Figure 24), accurately representing the area as completely NO_x-sensitive.

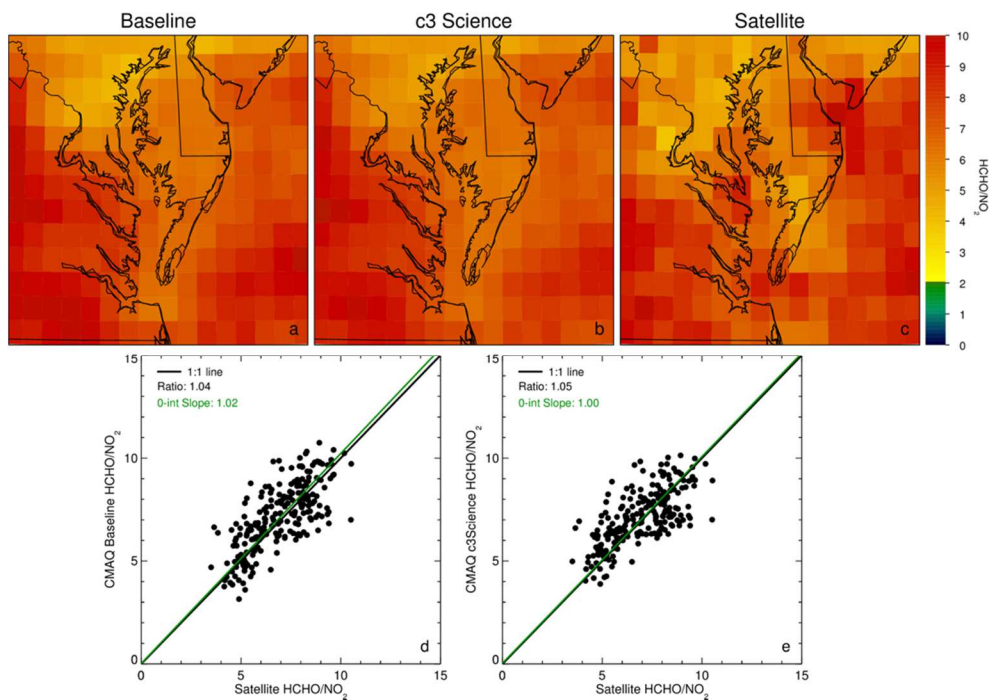


Figure 24: Same as Figure 22 except for the Chesapeake Bay region.

For some regions of the model domain, like the upper mid-west, both model simulations produce HCHO/NO₂ ratios that are lower than observed (see Figure 25 and Figure 26). In more rural regions, we see no improvement in the HCHO/NO₂ ratio between the two model scenarios, which could suggest the low ratios are due to HCHO. This is likely due to the CB05 chemical mechanism used by CMAQ. When the improved chemical mechanism CB6r2 is used [Ruiz and Yarwood, 2013], HCHO concentrations increase within the model, and are more consistent with measured values [Goldberg *et al.*, 2016]. Further improvements to the CB6r2 chemical mechanism, will increase HCHO concentrations within the model [Marvin *et al.*, 2017], and should improve model performance when compared to satellite observations. In the Chicago and Detroit metropolitan centers, we see increases in the HCHO/NO₂ ratio

when model improvements are applied (Figure 25 and Figure 26, respectively). This is expected as decreases of NO_x in urban centers would raise the HCHO/NO_2 ratio.

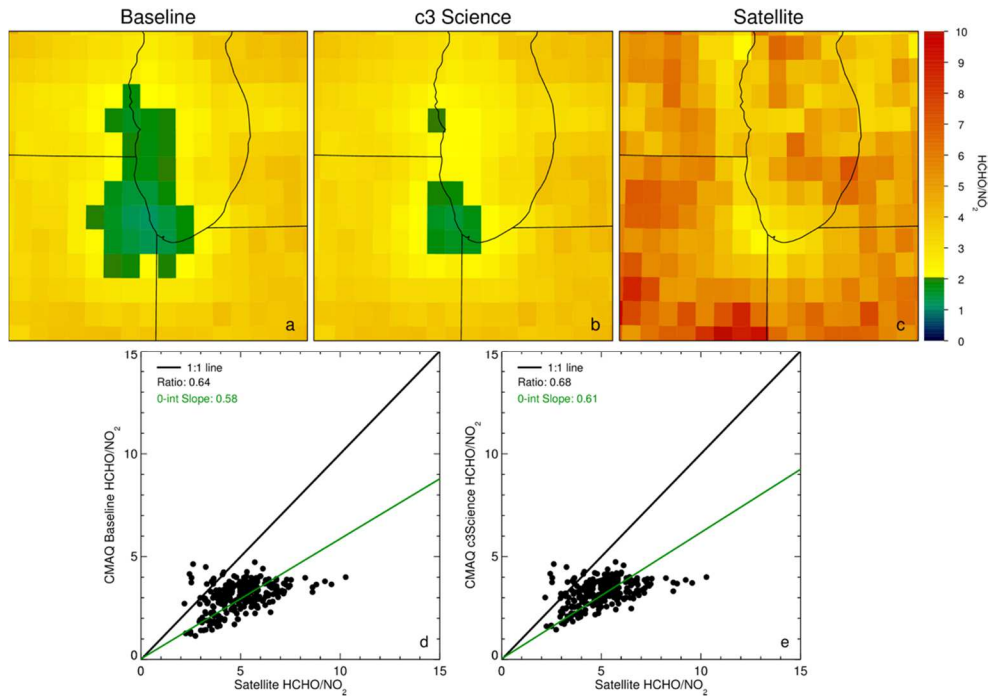


Figure 25: Same as Figure 22 except for the Chicago region.

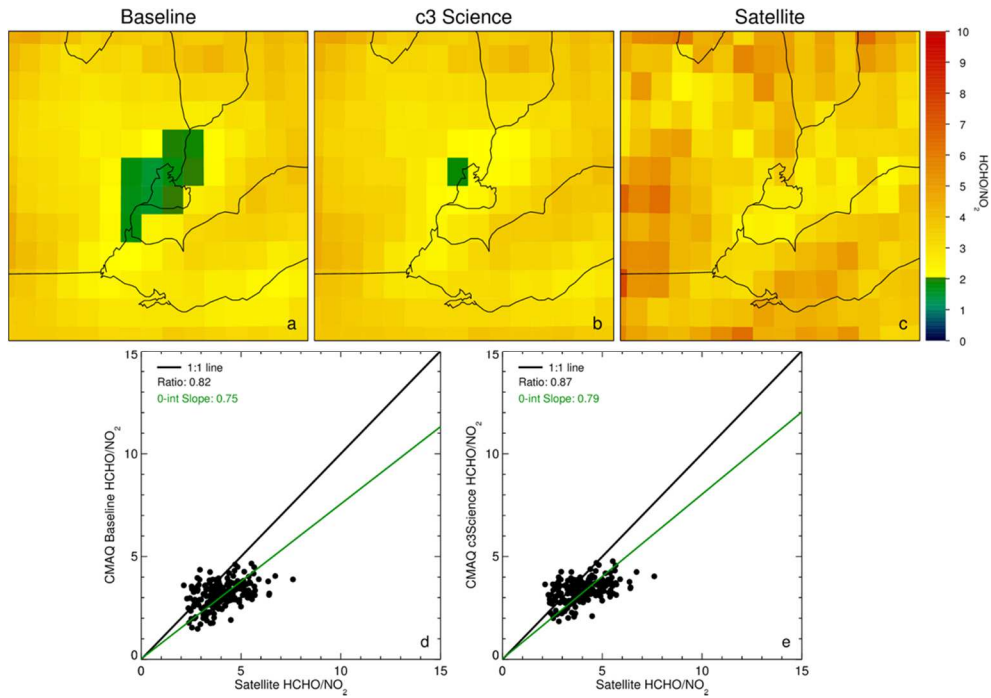


Figure 26: Same as Figure 22 except for the Detroit region.

Based on the satellite comparison, the c3 Science scenario more accurately represents the state of the 2011 atmosphere. While further improvements to HCHO and NO_x representation in rural regions within the model are necessary, the c3 Science scenario improves the representation of O₃ photochemistry in urban centers throughout the modeling domain, creating a more realistic model framework for guiding public policy [Ring *et al.*, 2018].

3.5: Modeling Scenarios for 2018

To assist state agencies in developing air quality attainment strategies, future emissions estimates are generated based on expected economic growth factors, fleet turnover, future air quality regulations, etc. Development of the future emissions inventory aids in providing guidance and justification for emissions standards reductions and proposed government legislation to improve air quality. We use the 2018 projected emissions to examine the efficacy of the c3 Science model scenario in 2018 compared to 2011, and to assess a future attainment strategy developed by the MDE, discussed in the next section.

All 2018 model scenarios discussed in this section use the same meteorology as the 2011 simulations, therefore, all differences in O₃ between 2011 and 2018 are due solely to emissions. AM8O₃ for the 2018 baseline simulation is shown in Figure 27a. Modeled surface O₃ in 2018 is notably lower than in 2011 (Figure 27a), indicative of the expected air quality improvements that are projected to occur by 2018. Values of AM8O₃ for the 2018 c3 Science simulation are shown in Figure 27b. Fewer model

grid points are plotted for the c3 Science scenario than in the Baseline scenario, because some grid points no longer satisfy the criteria of 10 days with maximum 8-hr O₃ above 60 ppb in the 2011 simulation. Elevated O₃ levels are still present in the Chesapeake Bay and Long Island Sound; however, these regions are smaller and have a reduced magnitude when compared to 2011.

A scatter plot of the AM8O₃ at the CMAQ grid points closest to the AQS monitoring sites for the Baseline and c3 Science model scenarios is shown in Figure 27c. Some grid points in NJ, CT, and NY for example, are above the 1 to 1 line, indicating surface AM8O₃ increases in the c3 Science model scenario relative to the Baseline simulation. Conversely, all points in MD, and some in NY, as well as CT lie below the 1 to 1 line, indicating a reduction of AM8O₃ in the c3 Science scenario relative to the Baseline scenario. Sites with the highest AM8O₃ shown in Figure 27c are listed in Table 3. Values in the last column are bolded for sites showing reductions of AM8O₃ between the two model scenarios.

Table 3: AQS sites with highest modeled AM8O₃ shown in Figure 27c.

AQS SITE	2018 BASELINE (AM8O ₃ PPB)	2018 C3 SCIENCE (AM8O ₃ PPB)	ΔAM8O ₃ (PPB)
Pfizer Lab, NY	148.0	146.0	-2.0
Greenwich Point, CT	144.2	141.4	-2.8
Queens College, NY	137.8	136.4	-1.4
White Plains, NY	133.8	131.3	-2.5
Babylon, NY	131.3	126.2	-5.1
Riverhead, NY	128.1	126.5	-1.6
Essex, MD	130.1	116.5	-13.6
Furley, MD	128.8	116.5	-12.3
Sherwood Island, CT	116.2	122.3	6.1
New Haven, CT	114.9	116.8	1.9
Fort Griswold Park, CT	111.9	113.9	2.0
Edgewood, MD	117.7	108.0	-9.7

When comparing AM8O₃ for Baseline and c3 Science scenarios, it is important to remember there are three major modeling framework changes that comprise the c3 Science scenario, as described in Table 2. Decreasing the lifetime of NTR contributes to the domain wide increase (red color) shown in the Figure 27d. Since 2018 emissions are scaled based on 2011 emissions, the 50% reduction of on-road mobile NO_x emissions in 2018 is also necessary. This reduction is less effective in 2018 because the on-road mobile emissions are projected to be cleaner due to national regulations, engine turnover, new fuel requirements, and local control programs despite a projected increase in vehicle miles traveled [*McDill et al.*, 2015]. Essentially, further controls of on-road mobile emissions will be less effective because this NO_x source sector has already been significantly reduced. Large reductions of surface O₃ within the Chesapeake Bay and increases in the coastal NY metropolitan area are shown in Figure 27d. These are due to the c3 Marine emissions inventory adjustments.

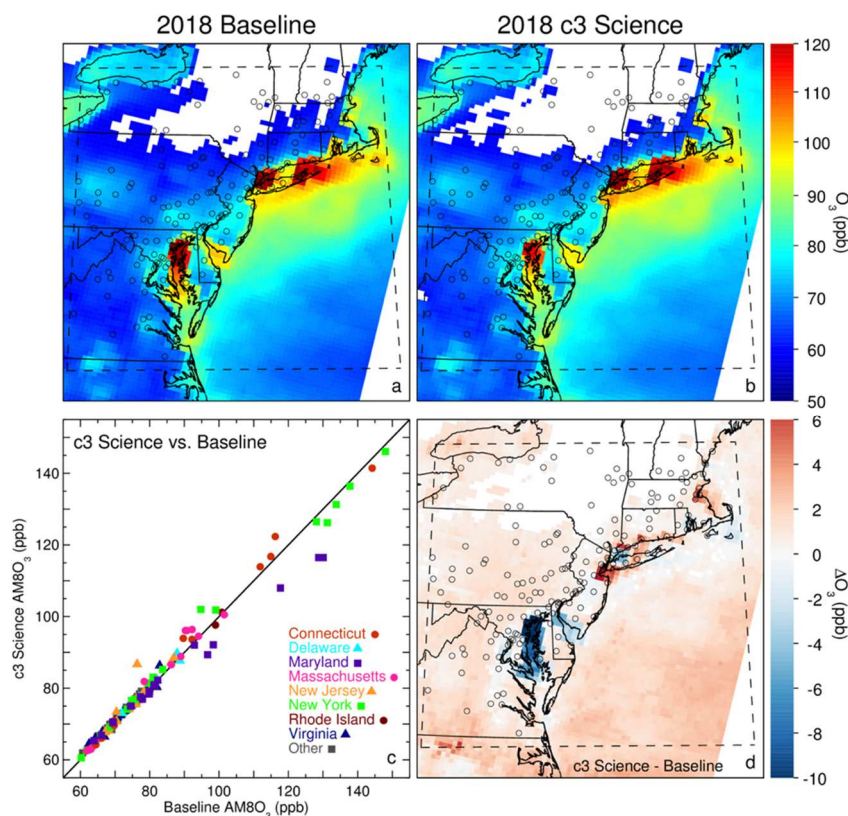


Figure 27: AM8O₃ for JJA 2018 for (a) Baseline and (b) c3 Science model scenarios. (c) A scatter plot of AM8O₃ for model grid points closest to the AQS sites for the Baseline vs. c3 Science scenarios. (d) A difference plot between c3 Science and Baseline model scenarios, highlighting AM8O₃ changes in the Chesapeake Bay and along the NYC metropolitan area coast.

As noted earlier, CMAQ produces extremely high values of modeled AM8O₃ (> 140 ppb) for some locations (Figure 27c). These high values are much larger than measured surface O₃ in 2011 and are therefore unrealistic. Nevertheless, model results at the six AQS sites with the highest AM8O₃ in 2018 show greater reductions of surface O₃ in the c3 Science scenario. This demonstrates the worst days for modeled surface O₃ in 2018 are more improved in the c3 Science model scenario when compared to the Baseline scenario.

We now return to the impact of c3 Marine emissions on NO_x and VOC limits of the photochemical O₃ production. AM8O₃ decreases of ~5-15 ppb are shown in

Figure 27d for grid points in the Chesapeake Bay region. Conversely, increases of ~2-6 ppb are shown for grid points in New York/New Jersey Harbor and Long Island Sound. We see regions of AM8O₃ decrease downwind of New York City, a different result from 2011 (Figure 27c). To further emphasize this point, Figure 28 highlights the impact of ship emissions on surface O₃ production along coastal regions within the 2011 and 2018 Science model framework (Table 2). In both Figure 28a and Figure 28b we see the expected decreases in AM8O₃ within the Chesapeake and Delaware Bay regions. In 2018 (Figure 28b), the AM8O₃ reductions are larger, showing more dramatic effects on surface O₃ production when large precursor sources (c3 Marine emissions) are lifted off the surface.

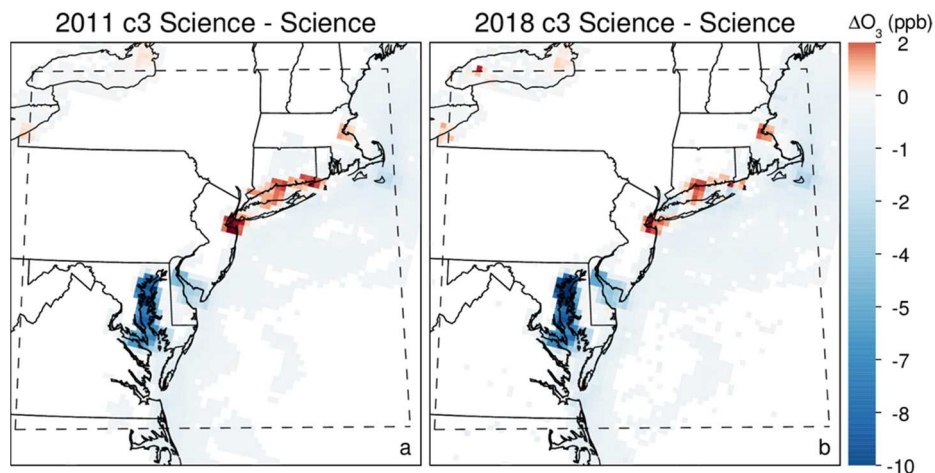


Figure 28: Difference plots of AM8O₃ between the c3 Science and Science modeling scenarios for (a) 2011 and (b) 2018. This highlights the effect of the c3Marine adjustment on surface O₃ production in the Science model framework.

Additionally, Figure 28a and b show smaller increases in AM8O₃ in the NJ/NY/CT area in 2018 than in 2011. Comparing Figure 28a and b highlights the reduction of AM8O₃ in the New York metropolitan area and along the Connecticut coast in 2018, perhaps showing parts of the region are transitioned and others have yet

to transition from a VOC to NO_x-sensitive region for photochemical O₃ production within the model.

3.6: Air Quality Attainment Strategy Analysis

State agencies use air quality models to quantify the effect of proposed legislation on future air quality. Here we examine one attainment strategy developed by MDE called “Scenario 4A”. This approach assumes that in 2018, emissions from EGUs will be at the best observed rates between 2005 and 2012 using existing emissions control equipment. To examine the impact of this attainment strategy on reducing surface O₃ production, Figure 29 shows the effect of implementing scenario 4A regulations within the c3 Science model framework. Implementing the scenario 4A attainment strategy reduces AM8O₃ domain wide, with the largest reductions occurring in areas significantly affected by power plants.

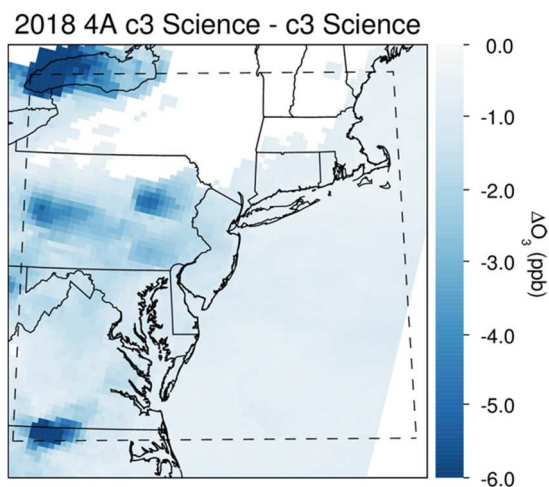


Figure 29: The difference of modeled AM8O₃ between c3 Science simulations both with and without the scenario 4A attainment strategy adjustments.

To quantify the AM8O₃ reduction in 2018 due to various scenario changes, we use Design Values (DV) as required by the EPA. An observed yearly design value is the 3-year running average of the observed fourth highest daily peak 8-hr average O₃

at an AQS site. The base design value (DVB) is a weighted average of yearly design values over a 5-year period [Wayland, 2014]. The second column of Table 4 provides 2011 DVBs for select AQS monitoring sites.

To assess attainment strategies for future years, a surface O₃ Relative Response Factor (RRF) is calculated. This metric represents the fractional change in modeled surface O₃ based on emissions changes between the base and future modeling scenarios [Wayland, 2014]. Figure 30 depicts a field of RRFs, showing geographic variation, for two model simulations used to calculate columns 5 (Figure 30a) and 6 (Figure 30b) in Table 4. For this analysis, the RRF is the AM8O₃ for 2018 divided by the AM8O₃ for 2011. The design value for the future model scenario (DVF) is the RRF multiplied by the DVB at each evaluated monitoring site. This value is compared to the NAAQS standards to determine whether the location of the monitoring site, in the simulated attainment strategy, is in attainment [Wayland, 2014].

Table 4: Observed and modeled design values calculated for several modeling scenarios at AQS sites in CT/NY/MD. Bolded values in last two columns show largest Δ DVF.

AIR QUALITY ATTAINMENT STRATEGY (SCENARIO 4A) ANALYSIS							
AQS Site	DVB	DVF Baseline	DVF 4A Baseline	DVF c3 Science	DVF 4A c3 Science	ΔDVF 4A Base - Base	ΔDVF 4A c3 Science – c3 Science
Greenwich Point, CT	80.3	76.37	76.12	71.54	71.15	-0.25	-0.39
Stratford Point, CT	84.3	77.30	76.76	76.53	75.98	-0.54	-0.55
Sherwood Island, CT	83.7	82.95	82.51	78.88	78.43	-0.44	-0.45
Hammonasset, CT	85.7	75.69	75.32	76.25	75.88	-0.37	-0.37
Pfizer Lab, NY	74.0	71.02	70.63	66.36	65.93	-0.39	-0.43
Queens College, NY	78.0	73.93	73.44	69.97	69.56	-0.49	-0.41
Babylon, NY	83.3	76.86	76.59	73.15	72.87	-0.27	-0.28
White Plains, NY	75.3	73.20	72.79	67.73	67.25	-0.41	-0.48
Davidsonville, MD	83.0	71.10	70.28	73.71	72.78	-0.82	-0.93
Padonia, MD	79.0	70.44	69.27	71.98	70.69	-1.17	-1.29

Essex, MD	80.7	73.24	72.85	72.24	71.73	-0.39	-0.51
Fair Hill, MD	83.0	73.78	72.64	75.97	74.62	-1.14	-1.35
Southern Maryland, MD	79.0	70.58	69.40	72.19	70.91	-1.18	-1.28
Fredrick Airport, MD	76.3	67.10	65.42	69.18	67.43	-1.68	-1.75
Edgewood, MD	90.0	82.02	81.47	82.41	81.67	-0.55	-0.74
Aldino, MD	79.3	70.14	69.41	72.22	71.34	-0.73	-0.88
Millington, MD	78.7	69.53	68.66	71.62	70.65	-0.87	-0.97
Rockville, MD	75.7	65.69	64.91	67.08	65.98	-0.78	-1.10
HU-Beltsville, MD	79.0	67.70	66.82	69.46	68.44	-0.88	-1.02
PG Equestrian Center, MD	82.3	70.52	69.64	72.85	71.84	-0.88	-1.01
Beltsville, MD	80.0	68.59	67.65	71.22	70.14	-0.94	-1.08

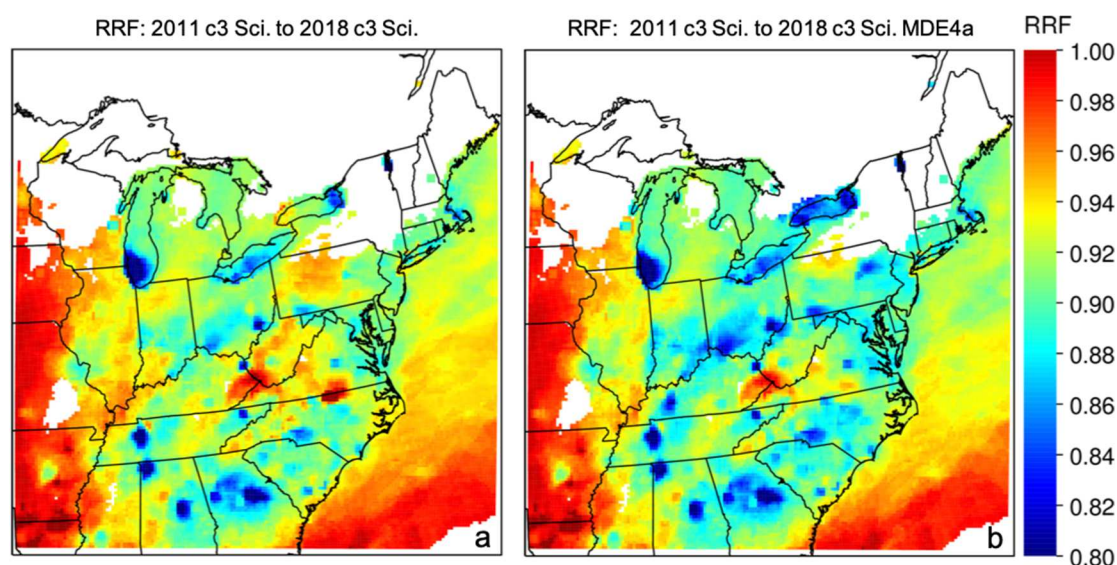


Figure 30: Contours of calculated Relative Response Factors (RRF) showing how surface O₃ is expected to change between 2011 and 2018 in the (a) c3 Science model framework, and (b) in the c3 Science model framework when the MDE4A attainment strategy is implemented for 2018. The blue colors indicate reductions of O₃, while the red indicate very little change, or an increase (dark red).

We have calculated DVs for four modeling scenarios. The scenarios are:

1. 2011 baseline to 2018 baseline
2. 2011 baseline to 2018 baseline with scenario 4A emissions reductions
3. 2011 c3 Science to 2018 c3 Science

4. 2011 c3 Science to 2018 c3 Science with scenario 4A emissions reductions

The DVs at selected AQS sites are shown in columns 3-6 of Table 4. Results shown in Table 4 indicate the effectiveness of the scenario 4A power plant emissions reductions in both the baseline and c3 Science model scenarios (columns 7 and 8). As expected, all DVs are negative in the two columns, meaning the 4A scenario is effective at reducing surface O₃ at all monitoring sites. Focusing on the italicized values in Table 4 columns 7 and 8 that denote the larger of the two differences for each site, we see scenario 4A emissions reductions are more effective at reducing surface O₃ in the c3 Science model scenario than in the Baseline scenario for almost all AQS locations. This result has significant relevance to the people responsible for developing air quality policy to control surface O₃. Improper representation of emissions from c3 Marine vessels in the CMAQ model framework underrepresents the impact of emission control scenarios in almost all near-coastal regions affected by shipping emissions [Ring *et al.*, 2018].

3.7: Conclusions

In this study, we examine the impact of Class 3 Commercial Marine Vessels (c3 Marine) emissions on air quality of coastal regions along the eastern US. Class 3 vessels are the largest within the global shipping fleet and have diesel engines with fuel displacement of at least 30 liters/cylinder. We have adjusted near-shore c3 Marine emissions to reflect a more realistic and consistent vertical distribution of pollutants. Model results that include the adjusted c3 Marine emissions show a decrease of AM8O₃

at Maryland AQS sites near the Chesapeake Bay, such as Essex, MD: ~6.5 ppb, where photochemical O₃ production is NO_x-sensitive. In areas such as Long Island Sound, the vertical distribution adjustment of marine emissions has increased AM8O₃ by ~3.5 ppb. This result, driven by a reduction of NO_x near the surface within the model, is due to the primary local production of O₃ being within the VOC-sensitive regime. Additionally, elevation of near-shore c3 Marine emissions off the surface allows the model to more accurately represent pollution dispersion and transportation associated with the c3 Marine emissions. This increases the lifetime of these chemical species within the model because they are no longer remaining near the emission source, over-producing O₃ in most regions and/or being removed from the atmosphere via wet or dry deposition [Castellanos *et al.*, 2011; Castellanos *et al.*, 2009]. Therefore, pollution transport from the Chesapeake Bay to the NY Metropolitan region could also contribute to increased surface O₃ production [Ring *et al.*, 2018].

Column HCHO and NO₂ retrievals from the Ozone Monitoring Instrument (OMI) are used to calculate HCHO/NO₂ ratios to determine the O₃ production regime in the atmospheric column above the surface for observations and model simulations. In the Baseline CMAQ model scenario, the HCHO/NO₂ ratio shows a larger area over the New York metropolitan area is transitioning from VOC-sensitive to NO_x- sensitive than is observed from the satellite. When empirically based model improvements are incorporated, known as the c3 Science scenario (see Table 2), a greater area over the New York metropolitan is NO_x- sensitive, improving model performance in relation to the observed satellite HCHO/NO₂ ratio and more accurately representing tropospheric conditions [Ring *et al.*, 2018].

We also examine the impact of these model improvements on future (2018) modeling scenarios. When comparing the 2018 Baseline to the c3 Science scenario, AM8O₃ reductions are larger for the Chesapeake Bay (Essex, MD: ~13.6 ppb), and smaller increases are shown for the New York metropolitan area than in 2011. Some regions around Long Island Sound show reductions in surface O₃, indicating the area has transitioned to NO_x- sensitive in the 2018 model simulation [*Ring et al.*, 2018].

Analysis of a NAAQS attainment strategy developed by MDE that simulates EGUs operating with optimal emissions rates shows that the strategy is more effective within the c3 Science scenario. Simulations using this improved model framework predict greater decreases in surface O₃, indicating legislation aimed at limiting O₃ precursors should be more effective than is demonstrated by the Baseline CMAQ model simulation [*Ring et al.*, 2018].

Chapter 4: Anthropogenic VOCs in the Long Island Sound, NY Airshed and their Role in Ozone Production (Ring et al. in prep)

4.1: Introduction

Over the past few decades, maximum tropospheric O₃ concentrations in the US have decreased [Bloomer et al., 2009; Goldberg et al., 2015; Simon et al., 2015] coinciding with anthropogenic emissions reductions, however low-level O₃ or background O₃ concentrations have increased [Cooper et al., 2012; Goldberg et al., 2015; Parrish et al., 2009; Simon et al., 2015]. The study by Simon et al. [2015] suggests the observed trend of decreasing difference between the lowest and highest O₃ concentrations may be a result of US environmental policy. Regulations targeting anthropogenic NO_x have significantly reduced emissions as reflected in the NEIs. This emission reduction has contributed to the decreasing maximum tropospheric O₃ trend [Simon et al., 2015]. Conversely, reported VOC emissions have remained relatively stagnant or increased in some regions of the US, and may explain the observed rise in background tropospheric O₃ [Simon et al., 2015]. To address the converging trends and reduce background tropospheric O₃ levels, control strategies targeting other O₃ precursors like VOCs may be necessary [Simon et al., 2015].

Many studies have observed high concentrations of tropospheric O₃ over near-coastal bodies of water [Cleary et al., 2015; Goldberg et al., 2014; Hastie et al., 1999]. The small-scale on-shore winds known as bay-breezes or lake-breezes move polluted air over coastal land and cities [Loughner et al., 2011; Loughner et al., 2014; Mazzuca et al., 2017; Sillman et al., 1993; Stauffer et al., 2015] impacting air quality in the coastal regions. Higher resolution modeling of the coastal marine boundary layer would

improve representation of these impactful circulation events [*Angevine et al.*, 2006; *Sillman et al.*, 1993] in air quality models, and better inform air quality policy.

The entire state of Connecticut (CT) is currently designated as non-attainment for the NAAQS O₃ standard (Figure 4) [*EPA*, 2018a; b] and is located downwind from New York City (NYC), a significant source of anthropogenic emissions. Connecticut also receives on-shore atmospheric circulation from the heavily trafficked Long Island Sound (LIS) as shown in Figure 34. In order to develop effective air quality attainment strategies, model and observational studies are needed to identify the meteorological conditions and emission sources directly contributing to the poor air quality in CT. Three research flights discussed in this chapter were conducted over the LIS and CT on May 17 and 18, 2017 and provide measurements of various atmospheric species during an O₃ exceedance event.

The estimated VOC/NO_x ratio derived from data and model comparison provides information about the atmospheric conditions favorable for O₃ production and the emissions adjustments and regulation that are necessary to improve air quality. To analyze the O₃ production VOC/NO_x ratio, *Duncan et al.* [2010] used a photochemical box model to calculate HCHO/NO₂ ratio for the Los Angeles, CA basin from modeled tropospheric columns. The geographic distribution of the model ratio field was compared to the HCHO/NO₂ ratio calculated using corresponding OMI tropospheric column observations. Satellite comparison of the OMI tropospheric column HCHO and NO₂ performed for the CONUS showed HCHO/NO₂ ratio values in major cities (including New York) within the 1-2 range. *Duncan et al.* [2010] cautioned that this

analysis (only based on modeling in the LA basin) may not be representative of regime transitions across the US.

Recently, *Schroeder et al.* [2017] performed box model simulations constrained to DISCOVER-AQ data for the Baltimore-Washington (BW) (2011), Houston (2013) and Colorado (2014) campaigns. They found a chemical transition point calculated with near-surface *in situ* data from the Colorado campaign to be 0.35 and a transition range of 0.9-1.8 when calculated with column profiles of HCHO and NO₂. They believe this discrepancy is due to the vertical distribution of HCHO in the lower free troposphere compared to NO₂ (Figure 31). Partitioning between NO and NO₂ favors NO at lower temperatures, so NO₂ concentrations decrease faster with altitude than HCHO [Schroeder et al., 2017]. Evidence for regional variation in the transition region was found when HCHO/NO₂ calculations were performed using vertical column densities for Houston and BW regions [Schroeder et al., 2017].

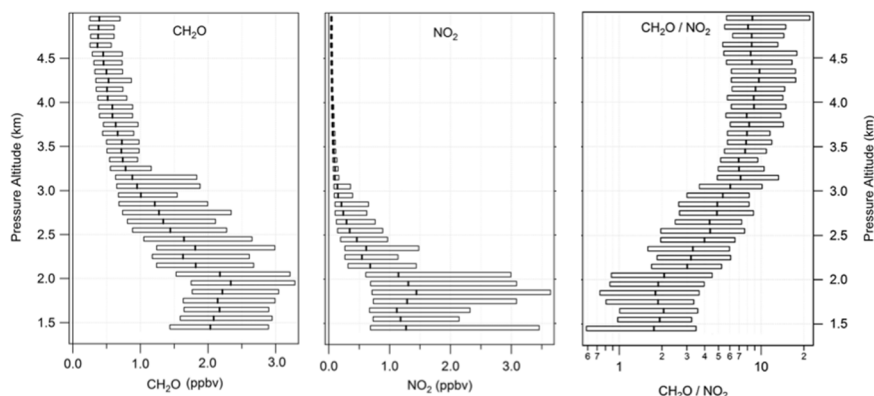


Figure 31: Colorado DISCOVER-AQ (a) HCHO data, (b) NO₂ data and the (c) ratio HCHO/NO₂ binned every 100m. The 25th and 75th percentiles are shown for each bin. (Figure 8 from [Schroeder et al., 2017]).

The binned concentrations of HCHO (CH₂O) and NO₂ as a function of height in Figure 31 depict possible limitations of using a tropospheric column product, like

the OMI satellite retrievals, to calculate the HCHO/NO₂ ratio. Tropospheric column satellite products and the global models used to calculate the *a priori* profiles may have difficulty representing the vertical structure of HCHO and NO₂ in the PBL [Schroeder *et al.*, 2017]. Therefore, use of these products to drive surface air quality policy legislation may not be adequate.

To develop effective air quality control policies, it is necessary to examine the relative contribution of individual VOCs to photochemical O₃ production [Avery, 2006; Carter, 1994; Jenkin *et al.*, 2017; Kamal *et al.*, 2016; Zaveri *et al.*, 2003]. A metric known as the Maximum Incremental Reactivity (MIR) represents the relative importance of individual VOCs to surface O₃ production [Carter, 1994; 2010b]. The MIR scale is calculated from photochemical box-modeling scenarios where the VOC/NO_x ratio is VOC-sensitive, meaning VOCs have the maximum impact on O₃ production [Carter, 1994]. To calculate MIR, Carter [1994] performed two model simulations: the baseline VOC-sensitive simulation, and one with added VOC concentrations that produce an increase of O₃ that follows along a linear trend [Carter, 1994]. The difference in produced O₃ between the two simulations, divided by the amount of added VOC, yields the MIR (O₃ produced / g of VOC) [Carter, 1994]. This metric is a useful tool to quantify the O₃ produced from specific VOC species and therefore emission sources. The MIR values were developed for use in the Statewide Air Pollution Research Center (SAPRC) chemical mechanism [Carter, 1994]. Based on guidance from Carter [2010a], the California Air Resources Board instituted controls on VOC emissions according to their associated MIR value, focusing air quality mitigation efforts on the VOCs that produce the most O₃ [Carter, 2010b].

A similar metric to the MIR called the Photochemical Ozone Creation Potential (POCP) is used in the Master Chemical Mechanism (MCM) [Carter, 2010a; Jenkin *et al.*, 2017]. The POCP equation from Jenkin *et al.* [2017] is as follows:

$$POCP_{VOC} = \frac{(O_{3VOC} - O_{3base})}{(O_{3ethene} - O_{3base})} \times 100. \quad (34)$$

In Equation (34), the change in O_3 due to the addition of the test VOC is divided by the change in O_3 when a reference VOC (ethene in this case) is added. Equal masses of the ethene and test VOC are added to the respective model simulations and the resulting relative change in O_3 is calculated [Jenkin *et al.*, 2017]. Both metrics are quite useful for describing the test VOC's capability of producing O_3 .

In this chapter we use the RAMMPP aircraft data and CMAQ model simulations to examine the importance of VOCs in the LIS region to O_3 production. Analysis of *in situ* VOC data provides policy-relevant information about VOC concentration and potential contribution to O_3 production for a LIS case study. Evaluation of RAMMPP with DISCOVER-AQ 2011 data from the Baltimore-Washington region shows general agreement concerning concentration and atmospheric profile of isoprene, NO_2 and O_3 , but important differences are found. Evidence for long range transport of SO_2 and O_3 into the LIS is presented. Forecast CMAQ simulations for the LIS O_3 episode are compared to measured data to highlight model successes and limitations and provide possible reasons for missed O_3 forecasts. Finally, analysis of the largest contributing model inventory sector to the modeled surface O_3 in the LIS are performed. This study focuses on those emissions sources and VOC species most important to the photochemical production of O_3 in the LIS.

4.2: DISCOVER-AQ 2011: Baltimore-Washington Campaign

For this analysis we compare O_3 , NO_2 , and isoprene collected in 2017 over the Long Island Sound from RAMMPP flights to measurements collected during the July 2011 NASA air quality campaign: Deriving Information on Surface Conditions from Column and VERTically resolved observations relevant to Air Quality (DISCOVER-AQ) conducted in the Baltimore-Washington (BW) region. All three species were measured onboard NASA's P-3B aircraft (<https://discover-aq.larc.nasa.gov/instruments.html>) during DISCOVER-AQ. The P-3B flight paths for the entire campaign are shown in Figure 32.

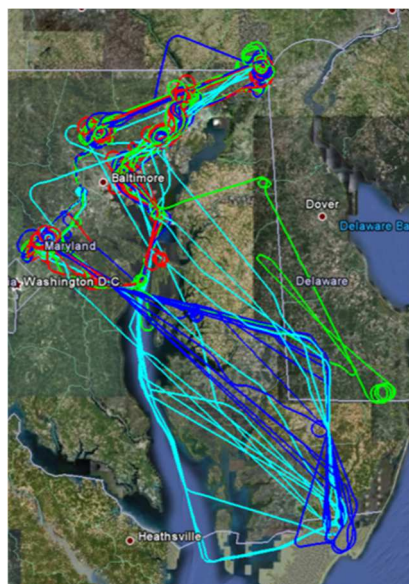


Figure 32: All of the flights for NASA's P-3B aircraft are shown. Flights were conducted from July 1st to July 29th, 2011 during the DISCOVER-AQ campaign. Spirals over designated areas were conducted to measure atmospheric profiles of various pollutants. (Figure is from <https://www-air.larc.nasa.gov/missions/discover-aq/discover-aq.html>)

Isoprene was measured by Armin Wisthaler's group from the University of Innsbruck with a Proton-Transfer-Reaction Mass Spectrometer. The abundance of NO_2 was measured using two instruments: the Thermal Dissociation-Laser Induced Fluorescence (TD-LIF) by Ronald Cohen's group from the University of California – Berkley, and the National Center for Atmospheric Research (NCAR) 4-channel

Chemiluminescence Instrument by Andrew Weinheimer’s group. O₃ was also measured using the 4-channel Chemiluminescence Instrument by Andrew Weinheimer’s group. All data were quality controlled and are publicly available at <https://www-air.larc.nasa.gov/cgi-bin/ArcView/discover-aq.dc-2011>.

4.3: VOC Can Analysis

As described in Section 1.4.3, aircraft flights were conducted over the Long Island Sound on May 17 and May 18 of 2017. A total of 35 canisters (hereafter called “cans”) of air samples were collected during the three research flights and 50 VOCs were measured in each can. To correct for any contamination from the can interior or sample inlets, two reference or blank cans were collected during the sample period. All VOC data were corrected by first calculating the average reference can concentration for each VOC species. This value was then subtracted from the measured VOC concentration in each sample can.

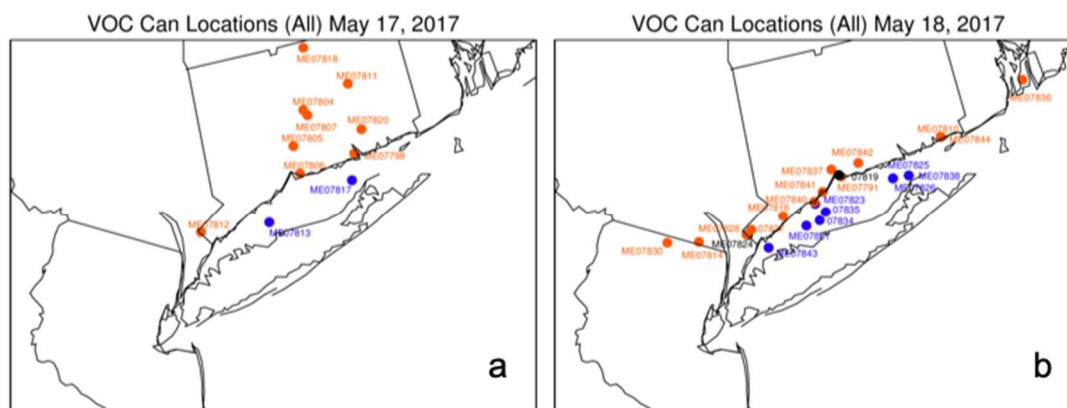


Figure 33: Geographic distribution of all VOC cans on (a) May 17, 2017 and (b) May 18, 2017. Cans were collected in flight and at various altitudes. Color distinctions represent cans collected over land (orange) and cans collected over water (blue). Cans with black colors overlap with other cans – color distinction purely for readability.

Figure 33 shows the spatial distribution of the cans collected on the two flight days. Most cans collected on May 18 were over the LIS or collected when flying along the CT shore of the LIS, sampling air downwind of New York City. Surface winds were westerly to south-westerly moving the New York City pollution plume over the LIS and CT (Figure 34).

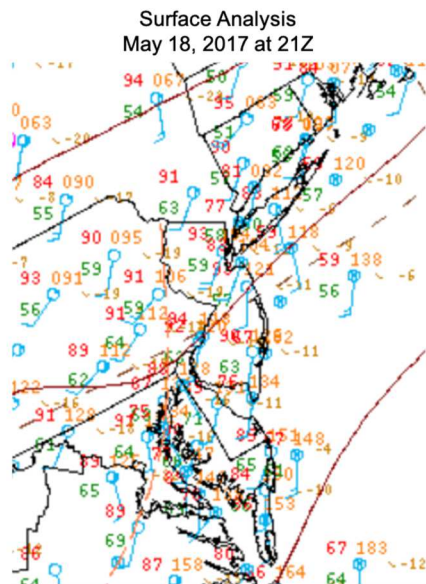


Figure 34: Surface observations of temperature, dew point, sky cover, sea-level pressure, pressure trend, wind speed and wind direction for the Long Island Sound region on May 18, 2017 (21Z).

(Image taken from the NOAA website:
<https://www.wpc.ncep.noaa.gov/html/sfc-zoom.php>)

4.3.1: Reactivity with OH

Reaction of VOCs with OH initiates a series of chemical reactions that lead to photochemical production of O_3 and production of oxidants [Kansal, 2009]. Reaction rates of VOCs with OH (k_{OH}) are the primary factors for determining individual contributions of various VOCs, and therefore importance to O_3 formation [Carter, 1994; Kansal, 2009].

All 50 VOC species measured by the RAMMPP aircraft are ranked by their reaction with OH ($k_{OH} \times [VOC] (s^{-1})$) to show the most important VOCs to O_3 formation in the LIS. The 12 VOCs with the highest ($k_{OH} \times [VOC] (s^{-1})$) values are

shown ranked by median (Figure 35a) and ranked by average (Figure 35b). The reaction rate constants used for this analysis are included in Appendix A.

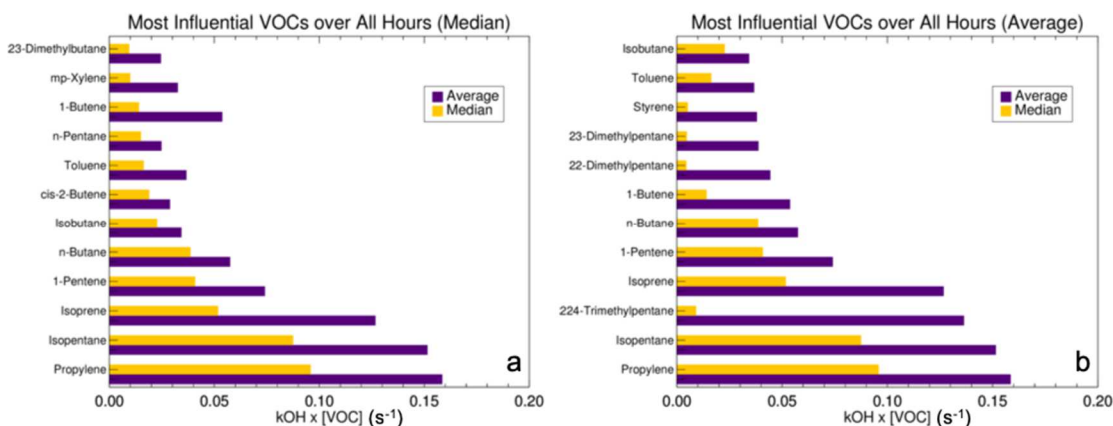


Figure 35: VOCs ranked by their reactivity with OH. Sorted by (a) median (gold bars) and by (b) average (purple bars).

Regardless of the sorting technique, propylene and isopentane are the two most influential VOCs in the LIS measured during the RAMMPP campaign. Isoprene is the 3rd most important VOC when ranked by the median, and the 4th when ranked by the average. One VOC canister contained concentrations of ~20 ppb for 224-trimethylpentane, explaining the large average $k_{OH} \times [VOC]$ value but relatively small median.

When sorting by time of day, it becomes clear that the reactivity of anthropogenic VOCs dominates biogenic isoprene in the morning (Figure 36a). Eight species rank higher than isoprene in Figure 36a compared to only two species in Figure 36b. This shows that morning VOC emissions may be contributing to the high afternoon O₃ concentrations in the LIS and downwind areas. Controlling the emission of anthropogenic VOCs, especially in the morning, may reduce the maximum O₃ concentrations reached in the afternoon. For example, propylene has a MIR value of

11.66 compared to 10.61 for isoprene [Carter, 2010b], and therefore could be responsible for significant amounts of tropospheric O₃ in the LIS. Many of the VOCs measured during the research flights (RFs) have large MIR values and therefore significant O₃ formation potential (Appendix A). Further analysis is necessary to determine the impact of these VOC emissions on surface air quality.

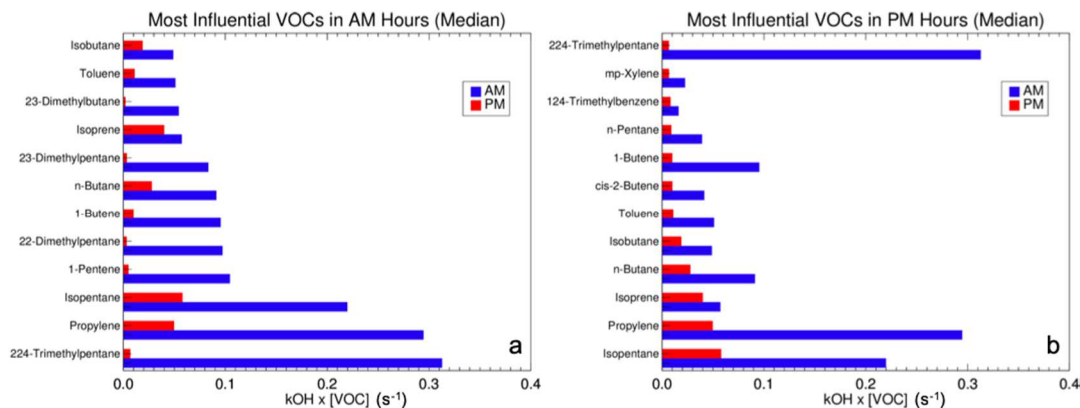


Figure 36: VOCs ranked by their reactivity with OH. Sorted by median values for cans collected during (a) AM (blue bars) and (b) PM (red bars) hours.

The lifetime of propylene is ~5 hours [Atkinson, 2000]. It is mainly emitted from petrochemical manufacturing facilities and diesel exhaust [Buzcu and Fraser, 2006; Hocking, 2005; Liu *et al.*, 2008]. Isopentane has a longer lifetime of ~1 day and is mainly emitted from gasoline vapors and engine exhaust emissions [Buzcu and Fraser, 2006; Liu *et al.*, 2008]. The compound 224-trimethylpentane has a lifetime of ~3 days with respect to reaction with OH [Atkinson, 2000] and sources mostly from industrial activity [Buzcu and Fraser, 2006]. Comparatively, isoprene (biogenic) has an atmospheric lifetime of ~1-2 hours [Atkinson, 2000].

In Figure 37, the OH reactivity of all measured species from the canisters are calculated and sorted for each VOC. The cans with reactivity values in the highest 10%

of the distribution are averaged and the cans with reactivity values in the middle 20% are averaged for each VOC species. The average reactivity values for each VOC in the highest 10% and middle 20% are grouped into their representative categories shown in Figure 37. This shows the relative importance of the various VOC groups to O₃ production in the NYC/LIS region. The values reported in Figure 37 are in relative agreement with the *Kleinman et al.* [2005] values, however more VOCs are present in the RAMMPP observations than in observations collected by *Kleinman et al.* [2005] in 1996. The large reactivity for the alkane group is because 26 alkanes were measured (over half of all VOCs), and isopentane is prominent in the VOC observations.

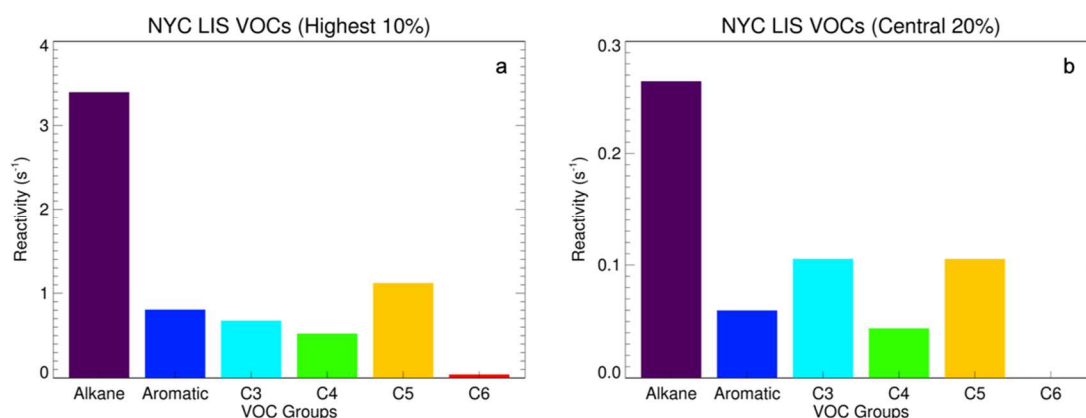


Figure 37: OH-VOC reactivity for the categories of hydrocarbons for the (a) top 10% of measurements and the (b) middle 20% of measurements for all VOCs measured throughout the RAMMPP 2017 flight campaign. Analysis is based upon Figure 12 from *Kleinman et al.* [2005].

In recent personal communication with Joel Dreessen from the Maryland Department of the Environment (MDE), we discovered that high concentrations of isopentane associated with high levels of O₃ are observed in the Chesapeake Bay. The winds carrying these polluted air masses generally come from the south, directly over

the Bay. The sources for the VOC-polluted air are not immediately known at this time. A forthcoming collaboration with Joel Dreessen and MDE will address these findings.

These results provide evidence that anthropogenic VOCs are playing a more important role in O₃ production in this region than was expected. However, given the sample size and short duration (2-days) of these flights, further observations are needed to confirm this result. Continued monitoring of NO_x, O₃, and VOCs in the LIS would improve confidence in the conclusion that emission controls targeting the most important anthropogenically produced VOCs would improve air quality in the LIS/CT region.

In a study by *Sillman et al.* [1993], model simulations incorporating a combination of polluted air, stagnant meteorology (usually associated with high pressure systems) and suppressed PBL height, indicate near-surface air over large bodies of water like the LIS or Lake Michigan (downwind from Chicago) is VOC-sensitive. It is likely that air aloft in the vertical column, over the LIS, is decoupled from the terrestrial biosphere and therefore low in isoprene concentrations due to the short lifetime ($\tau = \sim 1\text{-}2$ hours) [*Sillman et al.*, 2002]. Surface measurements of VOCs and NO_x may not explain the photochemical O₃ production regime conditions aloft (500 to 1000 m) [*Sillman*, 1999], therefore measurement of the urban plume from aircraft in subsequent campaigns would provide valuable information about the photochemical O₃ production regime downwind from NYC. This has important implications for states downwind as vertical mixing of the boundary layer can bring O₃ produced in VOC-sensitive conditions down to the surface [*Sillman*, 1999; 2002; *Sillman et al.*, 1993].

4.4: Evaluating RAMMPP and DISCOVER-AQ measurements

We compare RAMMPP and DISCOVER-AQ data to examine the average atmospheric profiles for O₃, NO₂, and isoprene for the duration of the respective flight campaigns. In the analysis from Chapter 3, we determine that in the CMAQ model framework the BW region and the LIS region are controlled by two different photochemical O₃ production regimes, evidenced by the behavior of surface O₃ (Figure 28). Despite measurement collection during different years, and in different locations, comparison of the data is valuable. We see notable differences in the vertical profile of isoprene (Figure 38a), similarities with NO₂ (Figure 38b) and differences again with O₃ (Figure 38c). Measurements of isoprene detected in the RAMMPP canisters collected over the LIS are much smaller than DISCOVER-AQ measurements collected over the BW region.

There is more NO₂ in the sampled atmospheric column over the LIS during the RAMMPP campaign than in the BW region during July, but measurements follow a similar profile trend reaching minimum mixing ratios by the 3000 to 3500 m height bin (Figure 38).

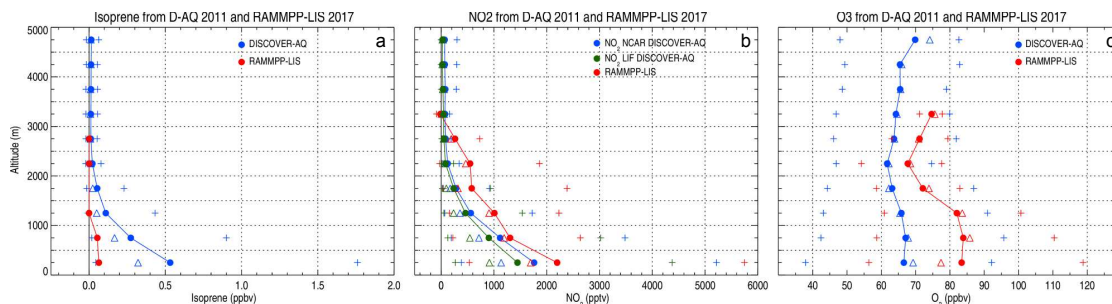


Figure 38: Comparisons between DISCOVER-AQ measurements from the P-3B in 2011 and RAMMPP measurements from the Cessna in 2017. Panel (a) shows the vertical distribution of Isoprene binned every 500 m over the entire sampling period, (b) shows the vertical distribution of NO₂ and (c) shows

the vertical distribution of O_3 . The triangles represent the 50th percentile, the plus signs represent the 5th and 95th percentiles, and the solid circles represent the average for each binned layer.

Figure 39 shows the vertical distribution of the 12 most influential VOCs from Figure 35b binned every 500 m. Average concentrations of these anthropogenic VOCs are generally larger in each altitude bin than isoprene.

Based on the NO_2 shown in Figure 38b and VOC concentrations shown in Figure 39, it is possible that VOC-sensitive conditions near the surface and aloft could be responsible for some of the O_3 produced over the LIS. The increasing concentrations of O_3 aloft could also indicate a possible up-wind source of pollution advected into the LIS region. More discussion of this will follow in the next section.

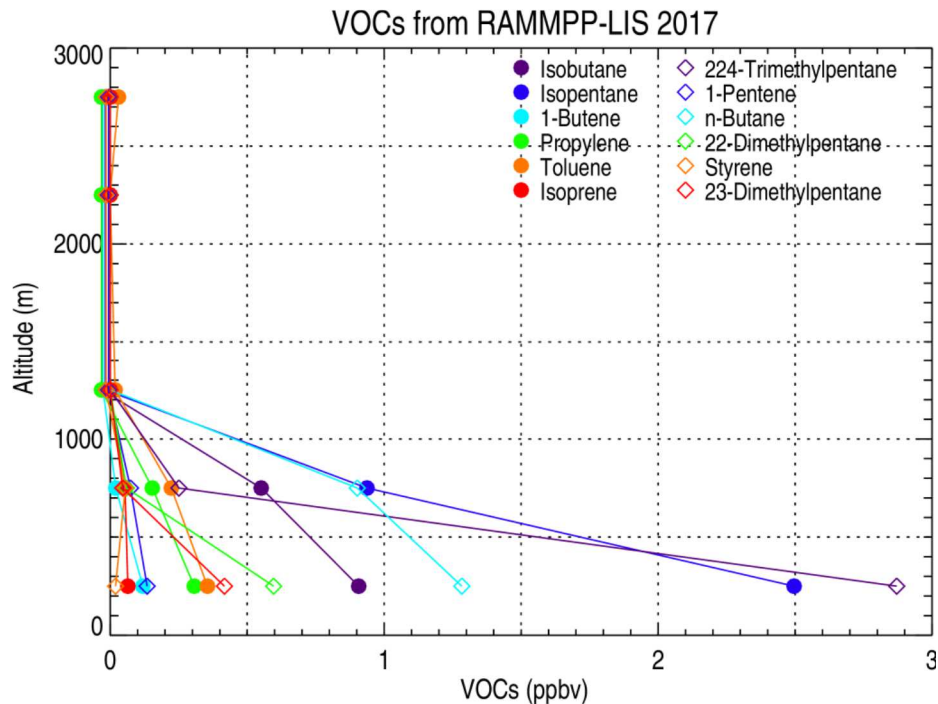


Figure 39: The vertical distribution of the 12 most influential VOCs binned every 500 m over the LIS sampling period. Each symbol represents the average for the binned layer.

4.5: Long-Range Transport

The long-range transport of O_3 and O_3 precursors into the Mid-Atlantic and Northeast regions of the US has long been an area of active research [Altshuiler, 1976; Angevine *et al.*, 2006; Bergin *et al.*, 2007; Cleveland and Graedel, 1979; Endlich *et al.*, 1984; Frost *et al.*, 2006; He *et al.*, 2016b; McDonald-Buller *et al.*, 2016; Ryan *et al.*, 1998; Stehr *et al.*, 2000; Taubman *et al.*, 2006; Vukovich, 1994; Wolff *et al.*, 1977]. Pollution transport from upwind sources bring air parcels with elevated O_3 in the lower free troposphere into the Northeast US and can contribute to local surface O_3 exceedance events if mixed into the planetary boundary layer [Bergin *et al.*, 2007; Cleveland and Graedel, 1979; Frost *et al.*, 2006; He *et al.*, 2016b; Lippmann, 1989; McDonald-Buller *et al.*, 2016; Ryan *et al.*, 1998]. As shown in Figure 40a, high O_3 levels were measured throughout the atmospheric column, with a minimum in the 2000 to 2500 m bin. In Figure 40b, we see low surface SO_2 concentrations that steadily increase with height, until the highest bin (3000 to 3500 m) where the average SO_2 concentration nearly doubles. Temperature for the aircraft profile that measured the highest SO_2 during RF3 is shown in Figure 40c. A temperature inversion, a common feature associated with pollution events [Moghani *et al.*, 2018; Ryan *et al.*, 1998], can be seen at about 2500m which can keep air aloft of the inversion from mixing with the air below meaning the O_3 above the inversion is likely not from local sources. The combination of the O_3 , SO_2 and temperature profiles in Figure 40 indicates O_3 and SO_2 from an upwind source (or sources) has been transported into the LIS region.

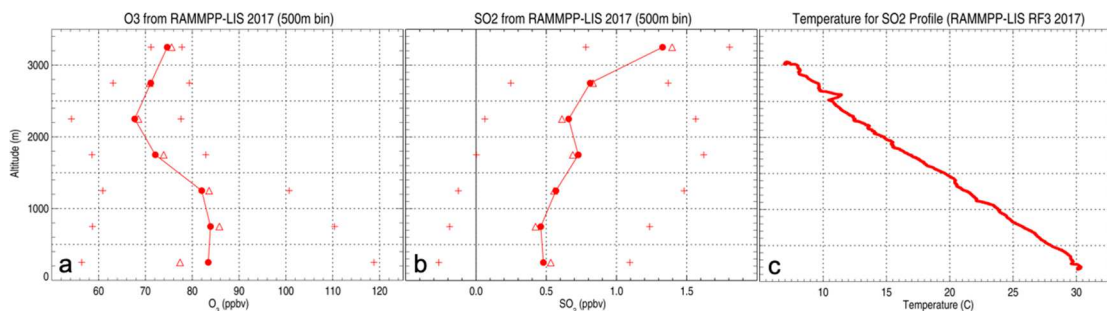


Figure 40: O_3 and SO_2 concentrations from all 3 research flights during the period binned every 500 m. (a) Average O_3 for each 500 m bin is represented by the filled circles, 50th percentile for each 500 m bin is represented by the open triangles, and the 5th and 95th percentiles for each 500 m bin are represented by the +. (b) The same as in (a) but for SO_2 . (c) The temperature profile along the aircraft spiral during RF3 that recorded the highest SO_2 measurement. The spiral began ~14:25 LST and ended ~14:51 LST on May 18, 2017.

We conduct ensemble HYSPLIT back-trajectories to examine the source of the measured SO_2 over the LIS (Figure 40b). Simulations were conducted for 2 days using archived 3 km High Resolution Rapid Refresh (HRRR) meteorological input (<https://rapidrefresh.noaa.gov/hrrr>). The ensemble back-trajectories end over the LIS at 3016 m (Figure 41b), the height of the highest SO_2 measurement. All SO_2 measurements were taken on May 18th between 14:48 and 14:50 LST. The black dots in Figure 41 along the back-trajectories represent power plants in operation on May 17th and 18th according to the Continuous Emissions Measurement Systems (CEMS) data. To monitor emissions, the EPA requires installation of CEMS instruments on power plant exhaust towers. Species important for air quality are measured, such as SO_2 , NO_x , Particulate Matter, and other Hazardous Air Pollutants (HAPs) (<https://www.epa.gov/emc/emc-continuous-emission-monitoring-systems>). We use these data to highlight active point sources of SO_2 along the back-trajectory paths, indicating possible sources for the high SO_2 and O_3 aloft over the LIS.

To confirm the presence of synoptic meteorological conditions favorable for atmospheric transport, we conducted two additional ensemble back-trajectories, from a height of 2500 m (Figure 41a) and 3500 m (Figure 41c) respectively. A consistent pattern is discernable among all three trajectory figures.

Ensemble Trajectories for Long Island Sound beginning at 20:00 UTC on 5/18/2017

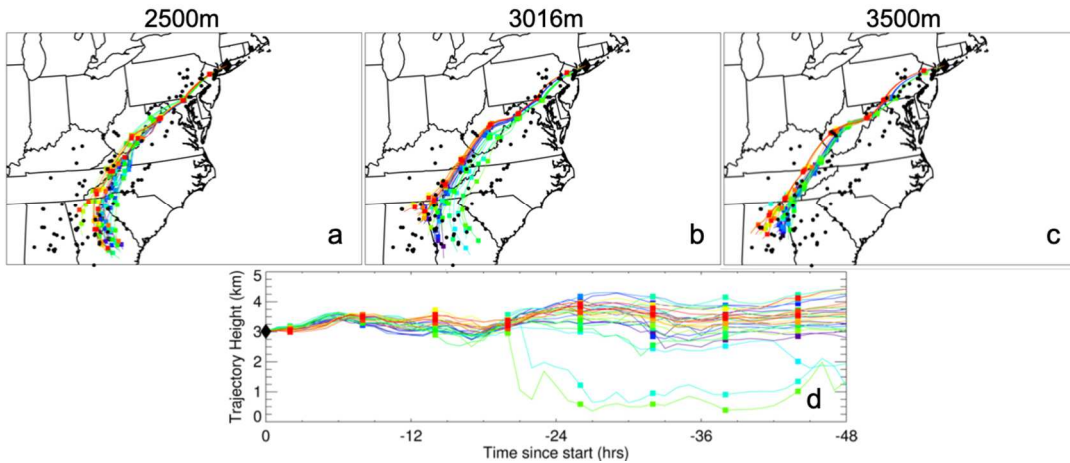


Figure 41: 2-day ensemble HYSPLIT back-trajectories ending at the location of the highest SO₂ measurement over Long Island Sound. The trajectories end at 3 different heights: (a) 2500 m, (b) 3016 m, and (c) 3500 m. (d) Trajectory height shown for the trajectory ending at 3016m (b) over the 48-hour back trajectory period. The black dots represent power plants with CEMS reported SO₂ emissions during the time period.

Review of a NASA Worldview true-color image of the US East Coast on May 17, 2017 at ~13:30 LST, shows fair-weather cumulus along the Appalachian Mountains, following the HYSPLIT back-trajectory paths (Figure 42). This image is approximately 24-hours prior to the highest SO₂ measurement in the LIS. The fair-weather cumulus clouds indicate vertical mixing in the atmosphere [Chen *et al.*, 2012; Ching and Alkezweeny, 1986; Seigel, 2014] where emissions from power-plants (black dots in trajectory figures) could be vented to an elevation shown in the HYSPLIT back trajectories and thus transported to the LIS.

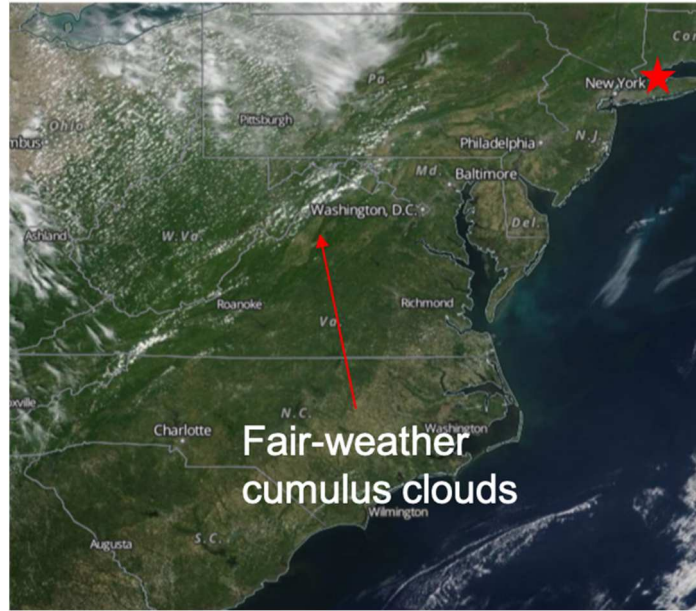


Figure 42: NASA Worldview true color image from Suomi NPP/VIIRS for May 17, 2017. The red star represents the approximate origin of the back-trajectories. Fair weather cumulus clouds are present along the HYSPLIT trajectory. Imagery can be found at: <https://worldview.earthdata.nasa.gov/>.

Additional comparison of the trajectories with archived NAM surface observations for May 17, 2017 at 18z and 21z shown in Figure 43, and archived Storm Prediction Center (SPC) analysis from May 17, 2017 at 12z and from May 18, 2017 at 00z shown in Figure 44 confirm the presence of a high-pressure system centered just east of the US coast near Bermuda. Surface observations in Figure 43 show southerly winds over the Appalachian Mountains, clear sky conditions, and warm surface temps (85-90°F); atmospheric conditions that are prime for O₃ production and transport. The upper air maps in Figure 44 further support the southerly to south-westerly winds along the Appalachian Mountains into the LIS region.

This atmospheric pattern is known for bringing pollution from upwind sources into the Mid-Atlantic and Northeast regions of the US. Emissions of SO₂ from power plants coincide with emissions of NO which produce O₃ relatively close to the emission

source, once the balance between NO and NO₂ has shifted towards NO₂ [Cleveland and Graedel, 1979]. Due to their short atmospheric lifetime, concentrations of NO (and NO₂) are not transported like SO₂ and the resulting O₃ [Cleveland and Graedel, 1979; He et al., 2016b; Ryan et al., 1998; Stehr et al., 2000] which explains elevated levels of O₃ (Figure 40a) and SO₂ (Figure 40b) in the highest 500m bins, with little NO₂ (Figure 38b).

In summary, analysis of atmospheric conditions during May 17, 2017 strongly indicate long-range transport of pollution contributed to the elevated measurements of SO₂ and O₃ aloft in the LIS region on May 18, 2017.

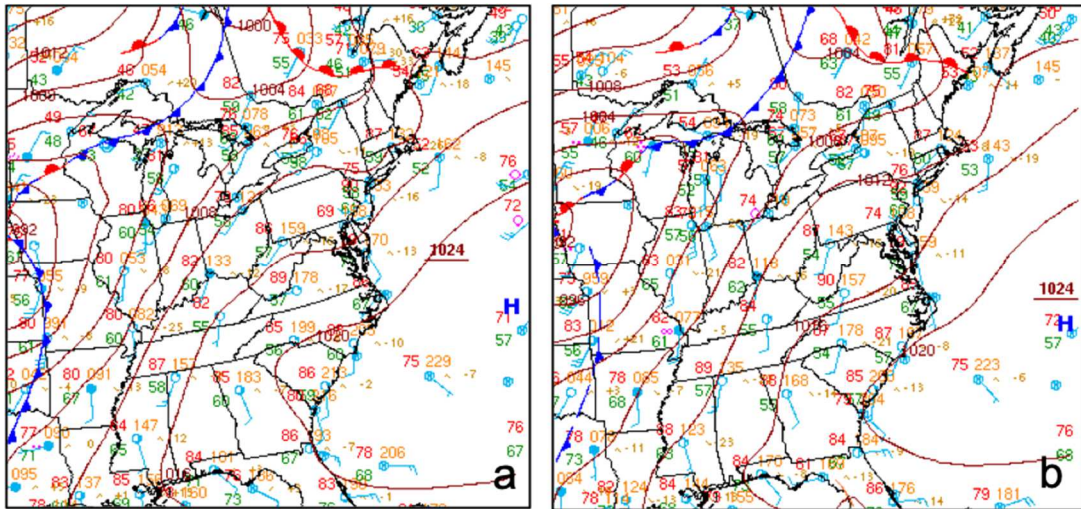


Figure 43: Archived surface analysis with observations from the NOAA Weather Prediction Center (WPC) for May 17, 2017 (a) at 18z and (b) 21z. Both maps are zoomed in over the Eastern US, highlighting the region of the HYSPLIT back-trajectories. Archived surface maps can be found here: https://www.wpc.ncep.noaa.gov/archives/web_pages/sfc/sfc_archive.php.

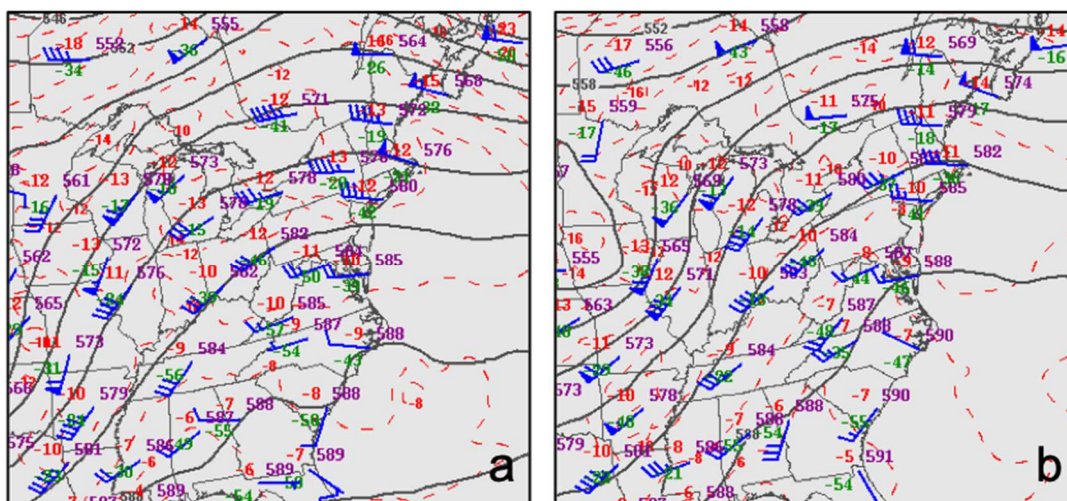


Figure 44: Archived 500mb upper air maps with observations from the NOAA Storm Prediction Center (SPC) for (a) May 17, 2017 at 12z and (b) May 18, 2017 at 00z. Both maps are zoomed in over the Eastern US, highlighting the region of the HYSPLIT back-trajectories. Archived upper air maps can be found here: <https://www.spc.noaa.gov/obswx/maps/>.

4.6: Comparison of RAMMPP Data with Forecast CMAQ Output

4.6.1: Measured and Modeled O₃, NO₂, and CO

Research flights were conducted on May 17 and 18, 2017 because the CMAQ air quality forecasts from NOAA indicated an O₃ exceedance event would occur. These forecast model simulations of O₃ are directly compared to observational data for all RFs (Figure 45). The CMAQ model platform used by NOAA is described in Section 1.5.2.

Figure 45 (a, b, and c) show O₃ measured along each RF plotted over modeled O₃. To create the curtain of modeled O₃, we find the closest surface CMAQ grid cell spatially and temporally to the aircraft observation and extract O₃ from all the grid cells above that point. The extracted O₃ is then plotted as a function of time and height for comparison to RAMMPP observations. This method allows the modeled atmosphere

to be sampled in a similar manner to how the aircraft observed the actual atmosphere. Figure 45 (d, e, and f) show the average of the 1st km of O₃ from CMAQ at hour 17 EST (d), hour 10 EST (e) and hour 15 EST (f) each corresponding to the figure above, with measured O₃ from the flight path again over plotted.

For RF1 (Figure 45a) the model and data show good agreement above 1.5 km but indicate differences by as much as 50 ppb near the surface (~17:15 and 17:45 EST). The spiral at 17:15 EST occurred just outside the model surface plume in Figure 45d, in the middle of CT. The spiral at 17:45 EST occurred at the CT border with Massachusetts shown in Figure 45d, again outside the CMAQ surface plume. Comparing Figure 45a and Figure 45d we notice CMAQ produced an O₃ plume mostly capturing the vertical structure of the high O₃ concentrations (between hours 16 and 17 for example), but missed the magnitude and horizontal extent of the plume. Measured near-surface air quality was in excess of 100 ppb for the spiral at 17:45 EST, but CMAQ only modeled concentrations between 50 and 60 ppb.

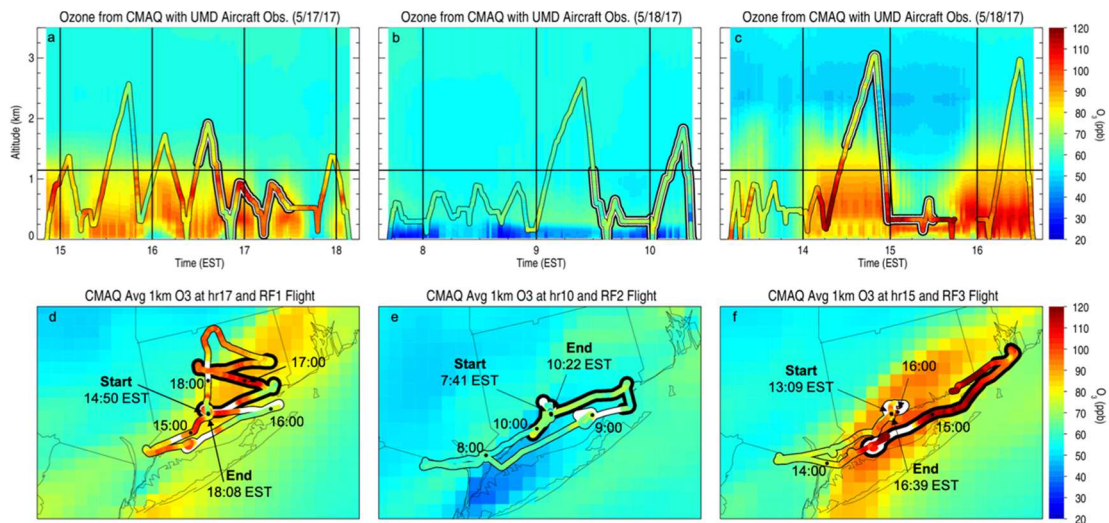


Figure 45: The top three panels show RAMMPP aircraft measured O₃ plotted over curtains of modeled O₃ concentrations from the closest CMAQ grid point for (a) RF1 (b) RF2 (c) RF3. The bottom three panels show RAMMPP aircraft measured O₃ plotted over average 1 km CMAQ O₃ at (d) hour 17 EST of RF1, (e) hour 10 EST of RF2 and (f) hour 15 EST of RF3. Hour demarcations are shown in all plots

for convenience. The bolded portion of the flight data (outlined in black) in the top three panels is also shown in its corresponding panel below (a and d, b and e, c and f). The bolded portion is centered around the hour shown in the bottom three panels. The white portion of the flight path in the bottom three panels corresponds to flight observations at altitudes higher than ~1 km.

For RF2 (Figure 45b), measured and modeled O₃ show similar concentrations aloft of ~250 m. At hours 9 and 10 the aircraft again measured near surface O₃ to be ~50 ppb higher than modeled values. In the 1st km over the LIS near NYC (Figure 45e), CMAQ produces O₃ concentrations ~ 40 ppb, much smaller than measured values. Overall, RF2 O₃ and average 1 km CMAQ O₃ are similar along much of the flight track.

Comparison of modeled and measured O₃ (Figure 45c) from RF3 shows similar results to RF1, however the differences are even greater (see ~14:15 EST). CMAQ does produce an O₃ plume at hour 16 (not pictured) that corresponds well with the aircraft measurements. Additionally, the model resolves a thin filament of high O₃ between hours 15 and 16 in Figure 45c corresponding to measured aircraft values but shows a slight vertical displacement. Clear model underestimations of O₃ are shown in Figure 45f, especially along the flight track over the LIS. This magnitude discrepancy is quite large (~30 to 40 ppb) and corresponds to areas where the measured O₃ levels are well above modeled values. Overall for this observational campaign, CMAQ forecasts generally capture the vertical and temporal extent of O₃ plumes but are below (in some cases severely below) the observed magnitudes.

Correlations of the measured and modeled O₃ for each RF are shown in Figure 46. These plots indicate poor or no correlation of the measurement with its corresponding closest CMAQ grid point and vertical level for all three RFs. All three panels indicate a general underestimation of modeled compared to measured O₃ shown

as negative normal mean bias (NMB) values and mean ratios < 1 , which agrees well with panels a-c in Figure 45. For modeled and measured O_3 to have good agreement, CMAQ must produce the appropriate magnitude of O_3 in the correct grid cell; a significant challenge for an air quality model with a resolution of 12×12 km grid boxes and 35 vertical levels. Based on Figure 45, we see CMAQ produced high O_3 but usually in the wrong location. This explains the poor correlations in Figure 46.

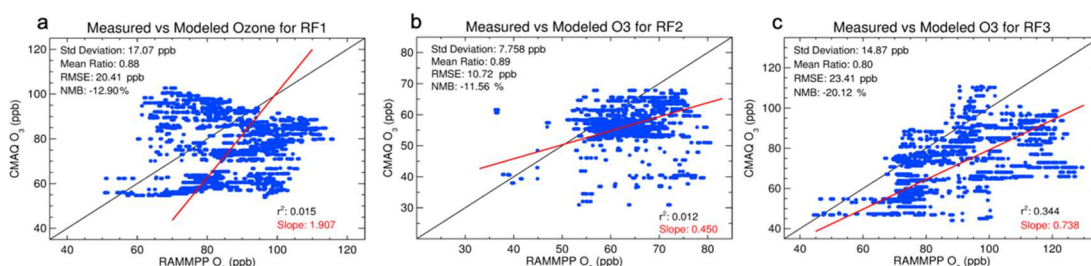


Figure 46: Scatter plots comparing measured O_3 from the RAMMPP aircraft and modeled CMAQ from the closest grid point and model layer for (a) RF1, (b) RF2, (c) RF3. The standard deviation, mean ratio (CMAQ O_3 /RAMMPP O_3), root mean square error (RMSE), normal mean bias (NMB), liner fit, and r^2 are calculated for each flight.

In this CMAQ configuration, the first model layer is ~ 40 m (necessary for computational speed) compared to the UMD CMAQ configuration that has a first model layer of ~ 20 m. The lower vertical resolution of the NOAA CMAQ model may explain some of the vertical displacement of O_3 . Small-scale boundary layer dynamics are known to impact surface O_3 concentrations [Loughner *et al.*, 2011; Loughner *et al.*, 2014; Mazzuca *et al.*, 2017; Sillman *et al.*, 1993; Stauffer *et al.*, 2015] so improving model resolution near the surface may help correct this issue.

Comparisons of measured and modeled NO_2 were also performed. NO_2 is critical to O_3 production and therefore requires appropriate representation within the CMAQ framework to accurately model O_3 . Figure 47 panels a, b, and c show measured

NO₂ along the RAMMPP flight track over a curtain of modeled CMAQ NO₂ from the closest model grid point. Near-surface modeled NO₂ values are generally high compared to measured, especially in Figure 47b just before 9 EST and Figure 47c at ~14:15 EST. In both cases, measured NO₂ is much lower than modeled NO₂.

Comparison of Figure 45b and Figure 47b show low O₃ collocated with very high NO₂ which may indicate near-surface VOC-sensitivity in the morning hours over the LIS within the CMAQ model framework. The corresponding measurements of O₃ and NO₂ in these regions do not agree with the model, therefore model adjustments are necessary to rectify this discrepancy with the data.

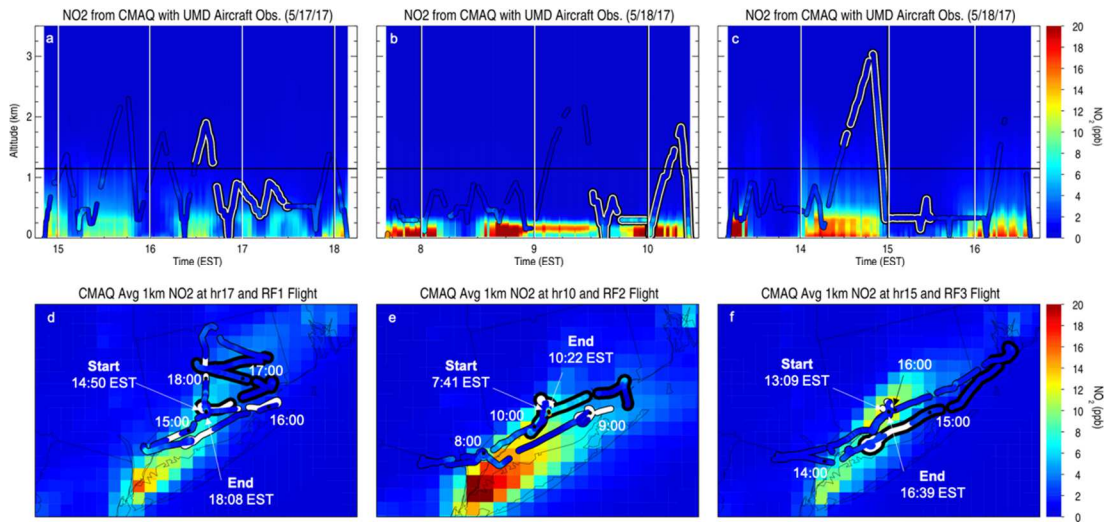


Figure 47: Same as Figure 45 except for NO₂.

Panels d, e, and f of Figure 47 show NO₂ measured by the RAMMPP aircraft (below ~1 km) over average ~1 km CMAQ NO₂ for the indicated hour. Measured NO₂ values agree well with modeled NO₂ outside of the plume downwind of NYC over the LIS (most evident in Figure 47e and f). Elevated levels of NO₂ are shown along the

model plume over the Connecticut coastline in Figure 47e, however measurements directly over the LIS within the urban plume are much lower than modeled NO₂.

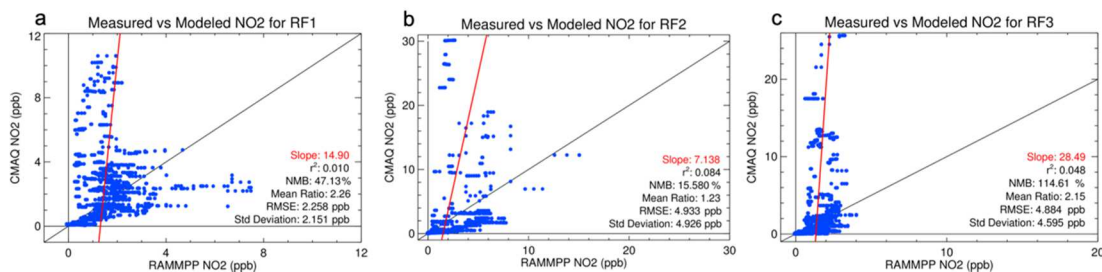


Figure 48: Same as Figure 46 except for NO₂.

Poor correlations of measured and modeled NO₂ are shown in Figure 48, along with significant overestimation of CMAQ NO₂ compared to measured NO₂. Based on Figure 47 and Figure 48, it is evident that the measurements and model agree well at low NO₂ mixing ratios, aloft of the near-surface region. The outliers shown in Figure 48 are from near-surface regions within the model. It is important to note that this forecast CMAQ simulation does not include the necessary 50% reduction of mobile NO_x first discovered by *Anderson et al.* [2014]. Implementation of the 50% emissions reduction of mobile NO_x will likely not be enough to offset the large overestimation of model near-surface NO₂ in the noted plumes shown in Figure 47. Therefore, it is possible that the NO₂ emissions in some model grid cells are much too high and/or vertical mixing of pollutants within CMAQ is not representative of the actual atmosphere.

Comparison of measured and modeled carbon monoxide (CO) are performed to examine the vertical mixing within the forecast CMAQ model. Figure 49 panels a, b, and c show the CO measured along the respective flights over the curtain of model

output from the closest CMAQ grid point. Overall, the model produces much larger CO mixing ratios near the surface than are measured.

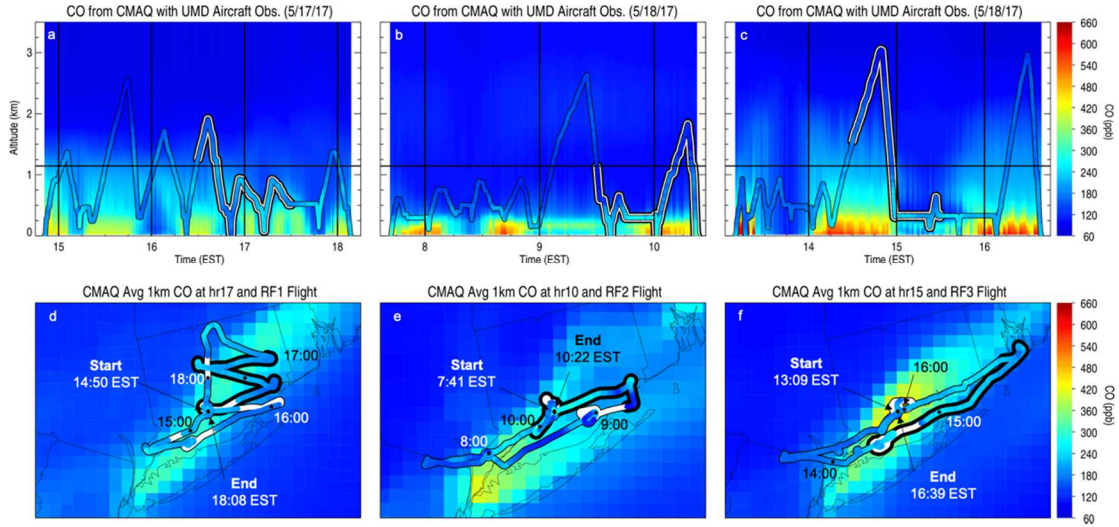


Figure 49: Same as Figure 45 except for CO.

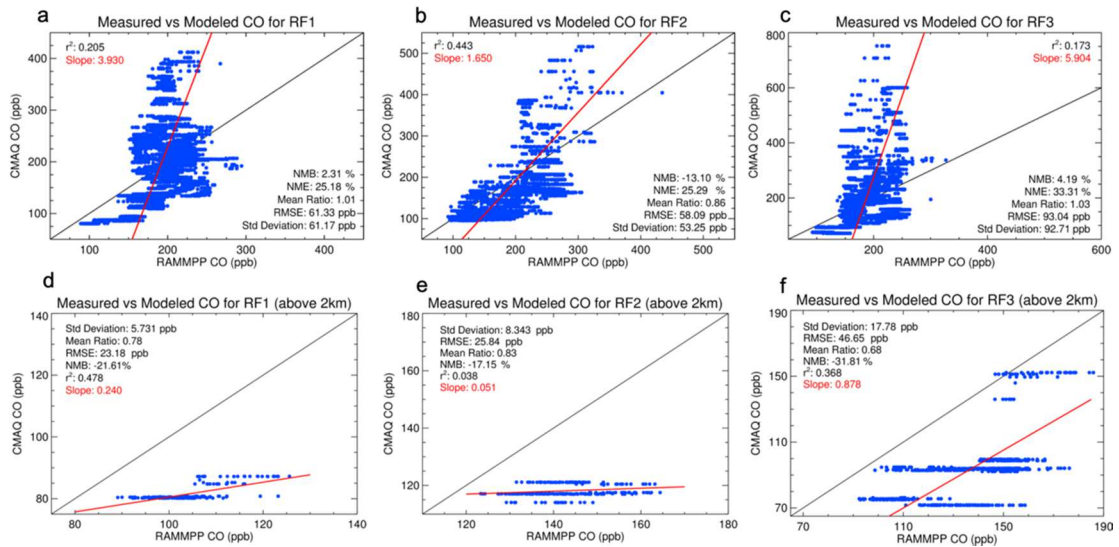


Figure 50: Same as Figure 46 except for CO (panels a, b, and c). Panels d, e, and f compare values above 2 km.

Figure 50 panels a, b, and c show scatter plots comparing measured and modeled CO for the respective RFs. CMAQ produces much larger CO values than are

measured in the near-surface region, especially for RF3 shown in panel c. Panels d, e, and f of Figure 50 compare measured and modeled CO at altitudes above 2 km and should mostly represent background CO values. In this region, CO from CMAQ is consistently low compared to observations. The overestimation at the surface coupled with the underestimation above 2 km could indicate vertical mixing within the model is too slow (i.e. CO is not mixing aloft fast enough), and/or it could indicate CO emissions are too high in the model.

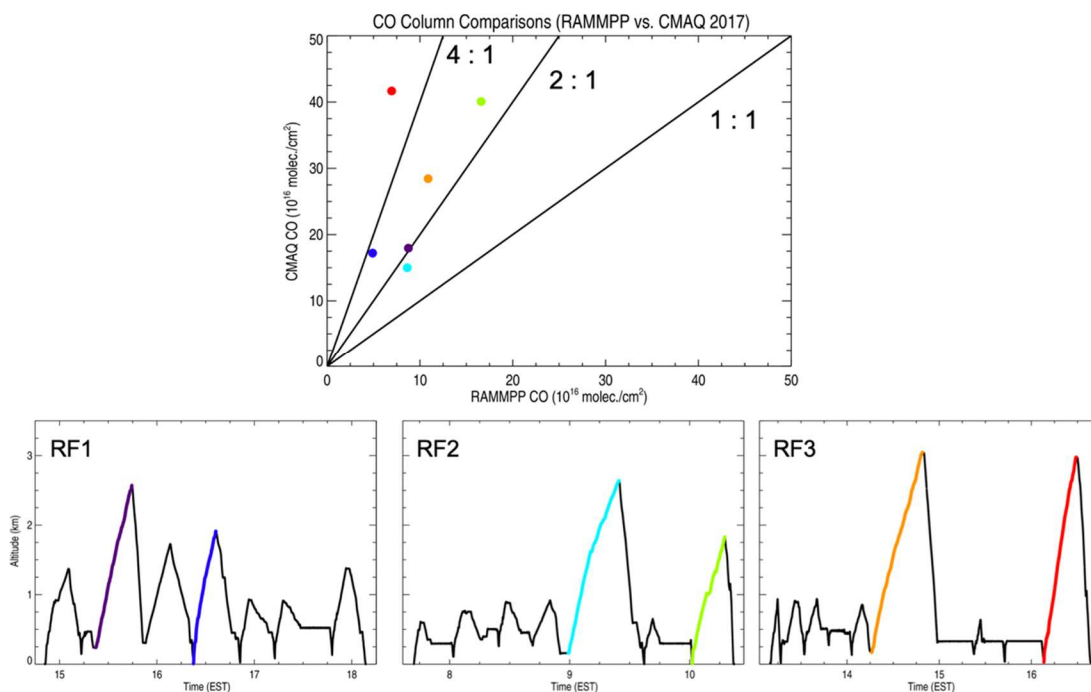


Figure 51: Comparison of measured and modeled column CO for six color coded spirals (three panels below). The 1:1, 2:1 and 4:1 lines show the degree of modeled CO overestimation compared to measured CO.

Integrated column concentrations of measured and modeled CO are calculated along six spirals (highlighted in Figure 51). The closest CMAQ grid point to each RAMMPP measurement is used in the column integration. The total column concentration for each spiral is shown in the top panel of Figure 51. Based on this

analysis, the modeled column CO is high by a factor of 2 for some spirals and a factor of 4 for others. The underestimation of modeled CO above 2 km coupled with the large overestimation of the integrated column CO provides evidence that vertical mixing within the model is too slow. Further analysis is necessary to definitively prove this point.

4.6.2: Measured and Modeled VOCs

It is difficult to compare observed and modeled VOCs because CMAQ uses the CB05 chemical mechanism and therefore produces concentrations for VOC groups rather than individual VOC species. We reference the CB05 documentation (Table 1 of *Yarwood et al.* [2005]) to speciate the measured VOCs into CB05 groups. It is important to note that we are only speciating the 50 measured VOCs into the appropriate CB05 groups; CB05 speciates *many* more VOCs into these representative groups. Therefore, comparisons of magnitude between measured and modeled VOC groups are inappropriate. Instead, we compare the relative composition of the VOC groups in the collected canisters to the composition of VOC groups in the closest CMAQ grid point at the appropriate vertical level. Once speciated, we use the OH reaction rates (k_{OH}) for each VOC group from CB05 (see Appendix B) to compare their relative importance to O₃ formation.

Comparisons for two VOC cans are shown in Figure 52. The left side of Figure 52 shows the relative composition of the collected VOC cans, and the right side shows the relative composition of the corresponding CMAQ grid cell. The time and altitude for the measurement and corresponding CMAQ grid cell are provided. For can ME07812 (top row of Figure 52), relative contribution of the species groups is in good

agreement, however CMAQ has more contribution of isoprene (ISOP) and less olefin (OLE) than was measured in the can. For can ME07817, much more paraffin (PAR) is attributed to the CMAQ grid cell than was measured. It is possible that other VOCs missing from the 50 that were measured contribute significantly to the PAR group, explaining the discrepancy in the composition. It is also possible that the VOCs are speciated incorrectly in CB05 or emissions are incorrectly attributed to the PAR group. Drawing any significant conclusions about CB05 speciation based on composition assessments of 35 VOC cans from a single case study is difficult. We can say however that these cans include more VOCs contributing to the OLE group than are present in the CMAQ output. Further analysis of the speciating algorithm with the newer chemical mechanism (CB6) that includes 26 more species may improve this discrepancy.

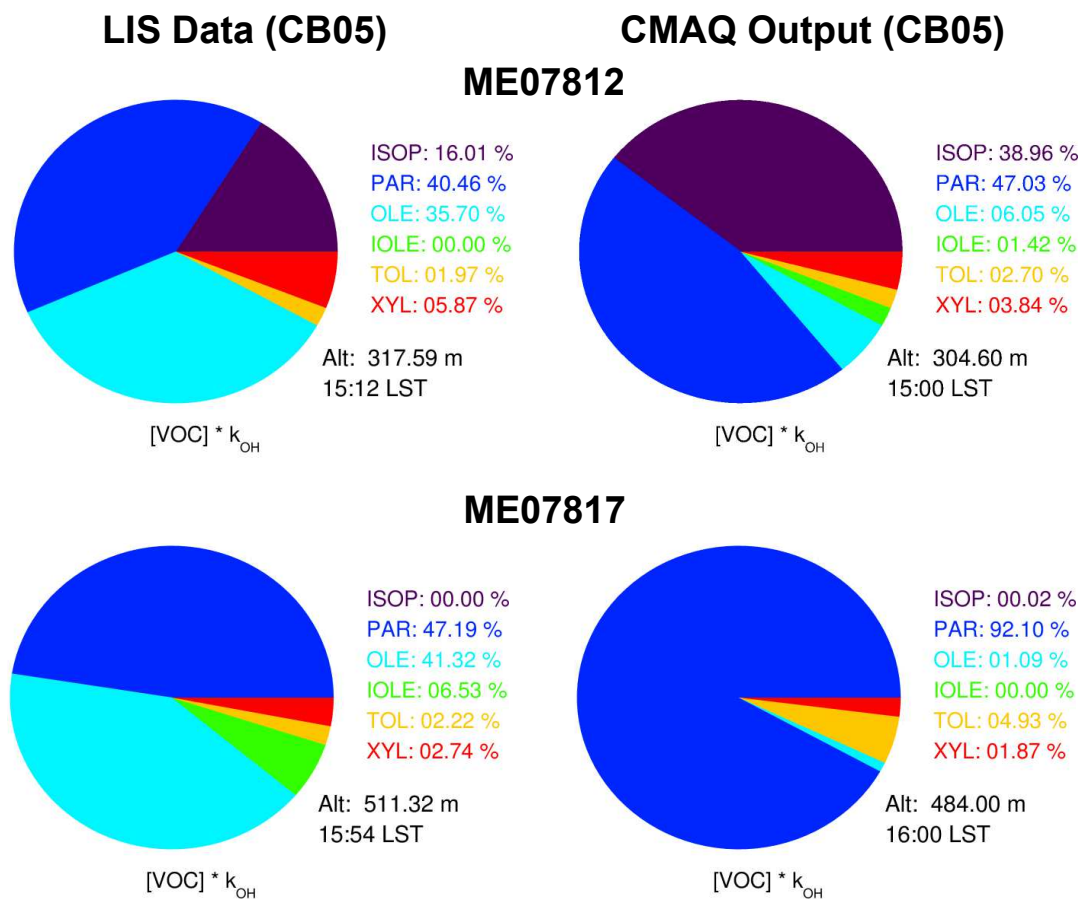


Figure 52: Composition comparison for two example canisters. Top: Can ME07812 data grouped as CB05 (left) and corresponding CMAQ grid point and height (right). Bottom: Same as the top except for can ME07817. Data are expressed in terms of the CB05 reaction rate for each CB05 VOC speciated group.

A recent study by *Zhao et al.* [2019] compared high resolution WRF-Chem modeled and observed 8-hour average O_3 and calculated MDA8 for the May 17 and 18, 2017 air quality episode in the LIS. They state that the biogenic emissions inventory Model of Emissions of Gases and Aerosols from Nature (MEGANv2.04) underestimates isoprene emissions by a factor of two in the LIS. This finding is justified by comparing the correlations between two respective model simulations (unadjusted isoprene and doubled isoprene) with measured surface O_3 . *Zhao et al.* [2019] reported that the simulation with doubled isoprene emissions correlates better with surface O_3

observations. *Canty et al.* [2015] determined CMAQ simulations using MEGANv2.1 rather than MEGANv2.04 biogenic emissions produced better agreement between modeled and OMI observed tropospheric column NO₂. The updated MEGANv2.1 inventory reduces isoprene emissions by ~25% compared to MEGANv2.04 [*Canty et al.*, 2015]. According to *Goldberg et al.* [2016], comparison of DISCOVER-AQ BW measurements of isoprene to Comprehensive Air Quality Model with Extensions (CAMx) simulations using MEGANv2.1 biogenic emissions shows significant overestimation of modeled isoprene. Correlation of modeled and isoprene observed during DISCOVER-AQ further improved when CAMx simulations using BEISv3.61 (the biogenic inventory used for all modeling in this analysis) were performed [*Goldberg et al.*, 2016]. Based on these previous studies, it is unlikely that MEGANv2.04 underestimates isoprene emissions by a factor of two. It is possible however that the improved correlation between WRF-Chem simulated and observed O₃ found by *Zhao et al.* [2019] is due to the overall increase in total (highly reactive) VOCs in the region. Their simulations generally underestimated O₃, so increasing concentrations of a highly reactive VOC with a high production of O₃ potential (isoprene in this case), may be the underlying reason for the improved O₃ agreement. The findings by *Zhao et al.* [2019] would be further strengthened if modeled concentrations of isoprene were compared to observations. The RAMMPP VOC observations from the same O₃ event measure little isoprene in the atmosphere compared to other anthropogenic VOCs during the study period.

Figure 53 shows the comparison of measured and modeled isoprene. VOCs were only measured from the air collected in the 35 cans throughout the three flights, limiting the available data for comparison. In

Figure 53 panels a, b and c, a curtain of modeled isoprene is plotted along the flight tracks, with reported isoprene from the VOC cans along the flight path (solid circles). Panels d, e, and f of

Figure 53 show modeled average ~1 km isoprene with the corresponding measured isoprene.

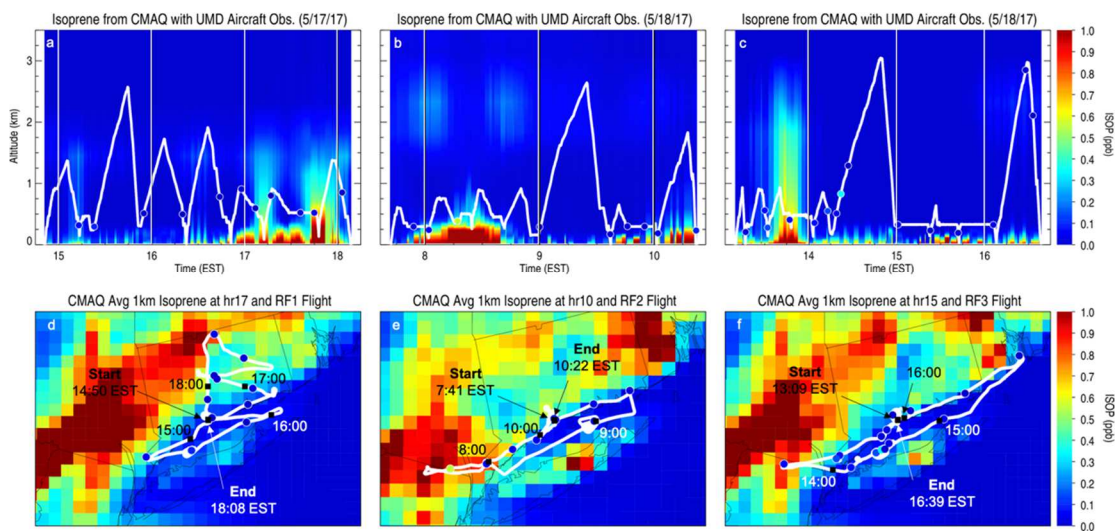


Figure 53: Same as Fig 45 but for Isoprene. The flight tracks are indicated in white with VOC can data shown as circles. Time-steps are shown as black squares in panels d, e, and f.

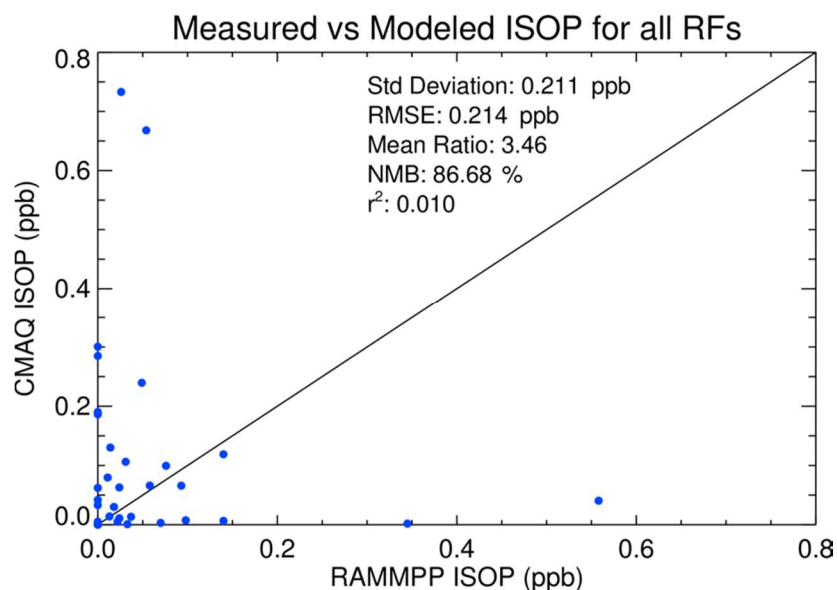


Figure 54: Comparison of measured and modeled isoprene for all three RFs during the May 17 and 18, 2017 flights.

Figure 54 shows no correlation between measured and modeled isoprene. The mean ratio of 3.46 indicates isoprene is generally high in the model compared to observations, but this high bias is not always true.

Figure 53 and Figure 54 compare observations and model output for the same O₃ event as examined by the *Zhao et al.* [2019] study and provide further evidence that the *Zhao et al.* [2019] findings (isoprene emissions are underestimated by a factor of 2) are questionable.

4.7: Modeled Sources of VOCs in Long Island Sound

We use CMAQ (configuration described in Section 1.5.1) to examine the various sources of anthropogenic VOCs in the model. All findings discussed in this section reference how the model sees the atmosphere rather than the actual atmosphere,

therefore this section comments only on VOC emission allocation within the CMAQ model framework, not on emission values.

4.7.1: Analysis with the Alpha2 Inventory

Motivated by the modeled surface O₃ dis-benefit when c3 Marine emissions are modified (discussed in Chapter 3 and exemplified by Figure 28), CMAQ simulations with perturbed VOC emissions values were performed. Reducing all anthropogenically produced VOC emissions in the model should show surface O₃ reductions in the same region as the O₃ dis-benefit in Figure 28 assuming the O₃ dis-benefit is because the region is VOC-sensitive. The initial CMAQ analysis is performed using emissions from the available Alpha2 inventory (described in section 1.5.3) because the Beta2 inventory emissions (described in section 1.5.4) were being gathered at the time of this analysis. CMAQ simulations with the Beta2 inventory are discussed in the next section.

To begin, all anthropogenic VOC emissions are reduced by 50%. This is executed in 3 steps:

1. Remove the biogenic emission file from the fully merged model-ready emission file using SMOKE;
2. Reduce all VOC species by half;
3. Re-merge the biogenic emissions with the adjusted VOC file to create the modified model-ready emissions file using SMOKE.

The difference in Average Maximum 8-hr O₃ (AM8O₃) (described in Section 3.3) between a CMAQ simulation with 50% reduced anthropogenic VOCs and one without adjustment to the anthropogenic VOCs in the c3 Science framework (Section

3.3, Table 2) is shown in Figure 55. The yellow star indicates the grid box with the largest change in surface O₃; a reduction of ~50 ppb.

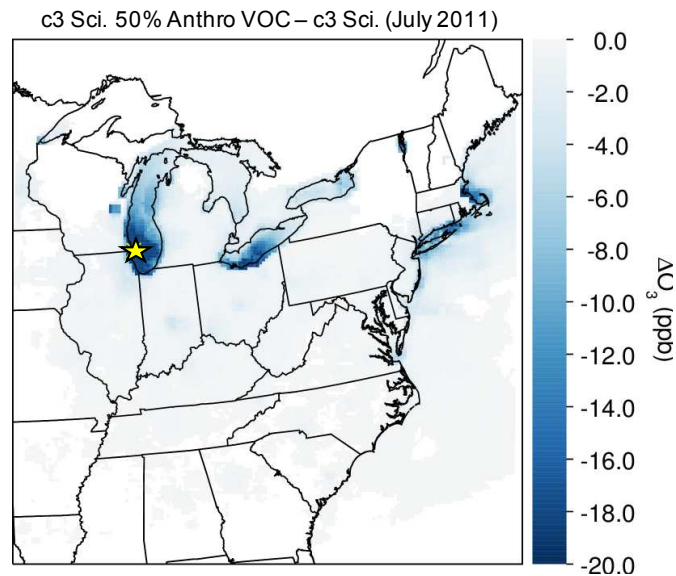


Figure 55: Change in surface O₃ due to reducing anthropogenic VOCs by 50% in the c3 Science (Table 2) model framework. Surface O₃ is represented as Average Maximum 8-hr O₃ (AM8O₃), a metric used by the EPA and described in Section 3.3. The yellow star marks the grid box with the largest change: 50 ppb.

The largest surface O₃ concentrations are mostly concentrated over large bodies of water, near major cities like Chicago, New York, and Boston. As expected, surface O₃ reductions are shown in coastal NY, CT and MA, the same locations as the O₃ dis-benefit shown in the analysis in Chapter 3. To determine which inventory sector (or sectors) contributes most to this ozone reduction, model simulations with VOC perturbations in specific source sectors are executed. Due to the spatial location of the largest O₃ reductions, we reduced the c3 Marine VOC emissions by 50%. This modification produced an undetectable change in surface O₃ throughout the domain (not shown).

Emissions from power plants (EGUs) and cars have historically been the focus of air quality regulation due to their large emission rates and ubiquity throughout the model domain. Again, based on the spatial distribution of the change in surface O_3 shown in Figure 55 being predominantly over water and near-coastal areas coupled with the small VOC emission values in the EGU inventory files, we determine that EGUs are unlikely to be the predominant sector contributing to the surface O_3 change.

The mobile sector consists of both on-road and non-road sources and has concentrated emissions in the same general regions as the large surface O_3 reductions in Figure 55; densely populated areas like Chicago, New York, and Boston. Model simulations were performed with 50% reductions of VOCs for the entire mobile sector (Figure 56a), the on-road mobile sector only (Figure 56b), and the non-road mobile sector only (Figure 56c). The yellow star in each figure represents the grid box with the largest reduction in surface O_3 . Reducing the VOCs from the entire mobile sector results in a maximum surface O_3 reduction of ~ 43 ppb. The on-road mobile sector adjustment results in a ~ 3 ppb reduction, and the non-road mobile sector adjustment results in a ~ 42 ppb reduction of surface O_3 . Based on Figure 56a and c, it is likely that the non-road mobile sector is responsible for a considerable amount of surface O_3 within the model framework. Further analyses with the improved Beta2 emissions are necessary to confirm that this holds true with the updated inventory.

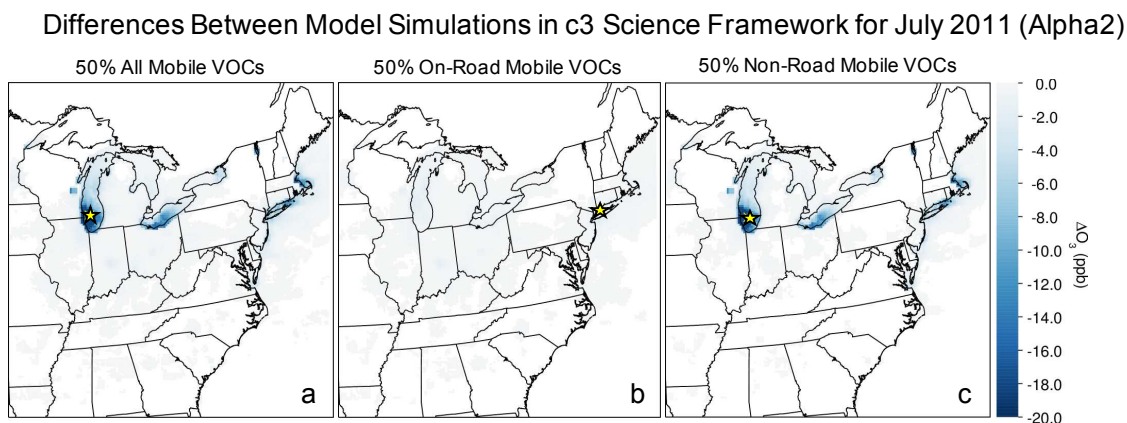


Figure 56: Change in surface O₃ due to reducing VOCs by 50% in the c3 Science (Table 2) model framework for the (a) entire mobile sector (b) on-road mobile only (c) non-road mobile only. Surface O₃ is represented as AM8O₃. The yellow star highlights grid box with largest surface O₃ change.

For reference, the non-road inventory represents emissions from vehicles used for recreational purposes, construction, lawn maintenance, industrial and commercial activities, logging, underground mining practices, pleasure boating, and railroad/railyard equipment (not locomotives) [McDill *et al.*, 2015; 2017]. This sector represents emissions from gasoline, diesel, compressed natural gas, and liquefied petroleum gas fuel types [McDill *et al.*, 2015; 2017]. The diverse nature of the non-road sector and its current representation (all sources lumped together) in air quality models makes discerning the specific fuel and/or non-road engine type most responsible for producing the problematic VOCs difficult.

4.7.2: Analysis with the Beta2 Inventory

Model simulations for the month of July in 2011 are performed with the updated Beta2 emissions inventory (described in Section 1.5.4). Anthropogenic VOC emission perturbations are executed as described in the previous section. Figure 57a and b show contours of AM8O₃ from model simulations in the c3 Science framework and the c3

Science framework with 50% reduced anthropogenic VOCs. The white areas within the model domain indicate locations where the minimum threshold of 6 days during the month with Maximum Daily Average 8-hour O_3 (MDA8) above 60 ppb was not met.

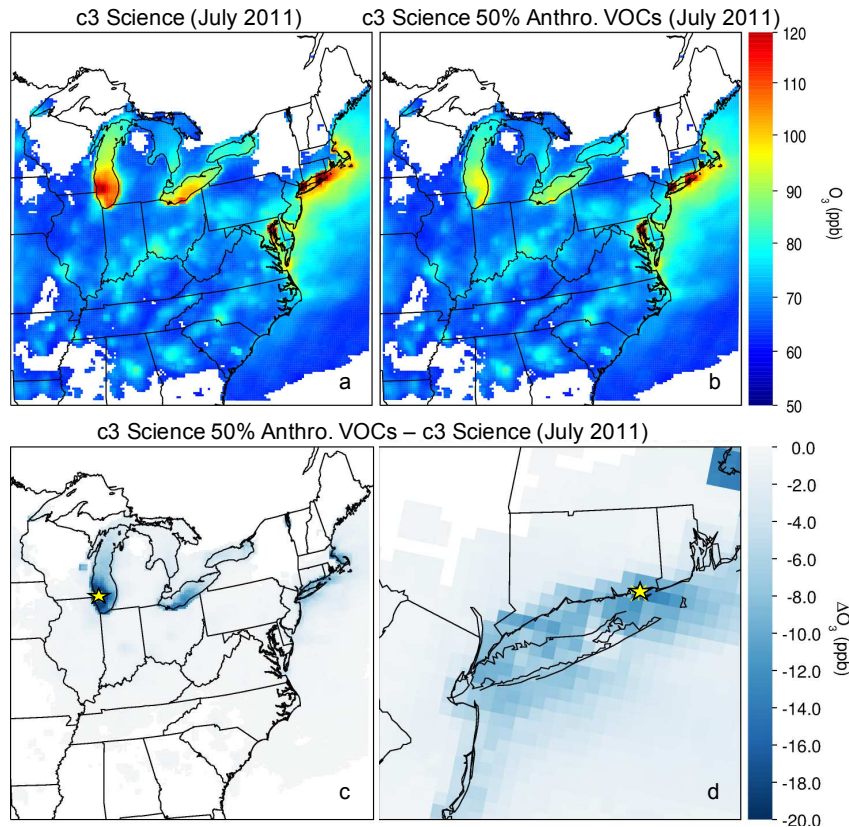


Figure 57: Contours of surface O_3 (represented as AM8 O_3) for July 2011 CMAQ model simulations of the (a) c3 Science model platform (b) c3 Science platform with 50% reduced anthropogenic VOC emissions. (c) The change in surface O_3 between the c3 Science and c3 Science 50% Anthro. VOC model simulations. (d) Same as (c) but zoomed in to LIS region. Yellow stars on (c) and (d) mark the grid cell with the largest O_3 reduction for the displayed domain.

Figure 57c and d show the difference between the model simulations shown in Figure 57a and b, with Figure 57d zoomed in over the LIS. We find consistent spatial distribution of surface O_3 reductions near coastal regions over large bodies of water as with the Alpha2 model simulations. The maximum surface O_3 change in the entire model domain is again over Lake Michigan but with a value of ~ 36 ppb. The maximum

change in the LIS region is ~15 ppb. The smaller maximum change in the Lake Michigan area using the Beta2 emissions compared to the Alpha2 emissions is due to the inventory updates discussed in section. One notable improvement between the two versions is for the non-road sector. In the Alpha2 inventory, the non-road emissions derive from the NEIv1 data, in the Beta2 the non-road emission derive from the NEIv2 data.

Focusing on the mobile sector, we conduct three model simulations with mobile VOC emissions reduced by 50% in the c3 Science framework. Figure 58 shows the surface O₃ reductions for the three VOC perturbation simulations. For the 50% reduced on-road mobile VOC emissions scenario, the maximum change is ~3 ppb in the LIS (Figure 58a and d). For the 50% reduced non-road mobile VOC emissions scenario, the maximum change is ~32 ppb over Lake Michigan (Figure 58b) and ~13 ppb over the LIS (Figure 58d). For the 50% reduced all-road mobile VOC emissions scenario, the maximum change is ~33 ppb over Lake Michigan (Figure 58c) and ~14 ppb over the LIS (Figure 58f).

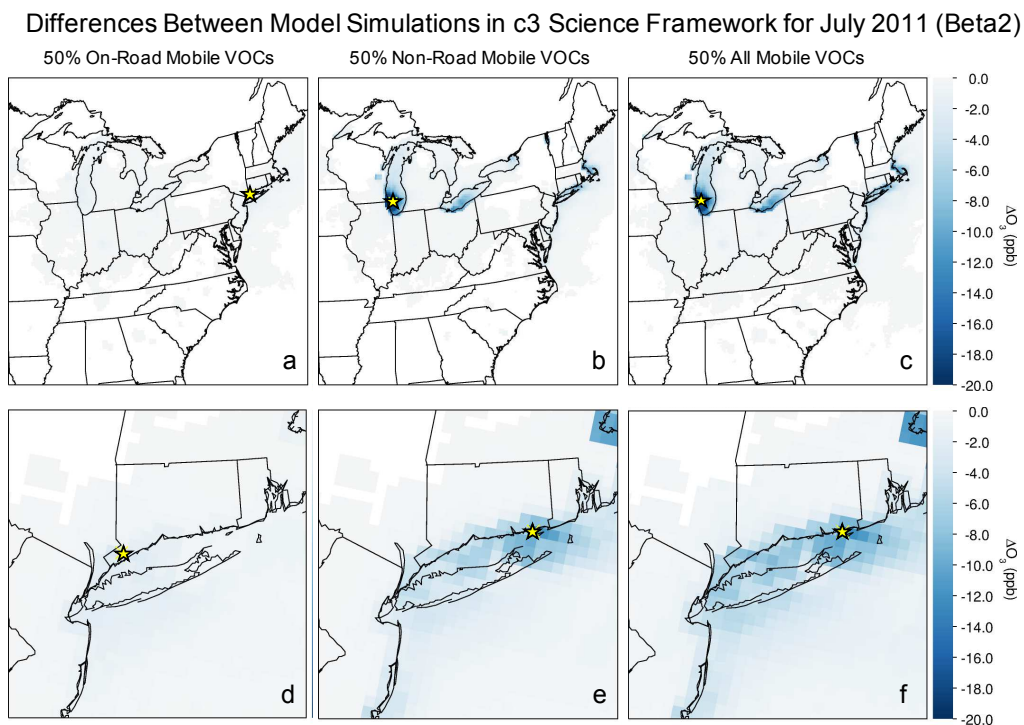


Figure 58: Same as Figure 56 but with Beta2 emissions. (a) c3 Science with 50% reduced on-road mobile VOC emissions, (b) c3 Science with 50% reduce non-road mobile VOC emissions, (c) c3 Science with 50% reduce VOC emissions for all mobile sectors. (d) Same as (a) but zoomed in to LIS, (e) same as (b) but zoomed in to LIS, (f) same as (c) but zoomed in to LIS region. Yellow stars on all panels mark the grid cell with the largest O₃ reduction for the displayed domain.

Figure 59 shows the relative contribution of various modeling scenarios with perturbed mobile VOC emissions to modeled surface O₃ reductions shown in the scenario with 50% reduce anthropogenic VOCs. The percent of the total surface O₃ change due to the 50% reduction of on-road mobile VOC emissions is shown in Figure 59a. The percent due to the emission reduction for the non-road sector is shown in Figure 59b, and percent due to the 50% adjustment applied to both mobile sectors is shown in Figure 59c. Together, these illustrate the impact of VOCs from the mobile emissions sectors to surface O₃ reductions domain wide.

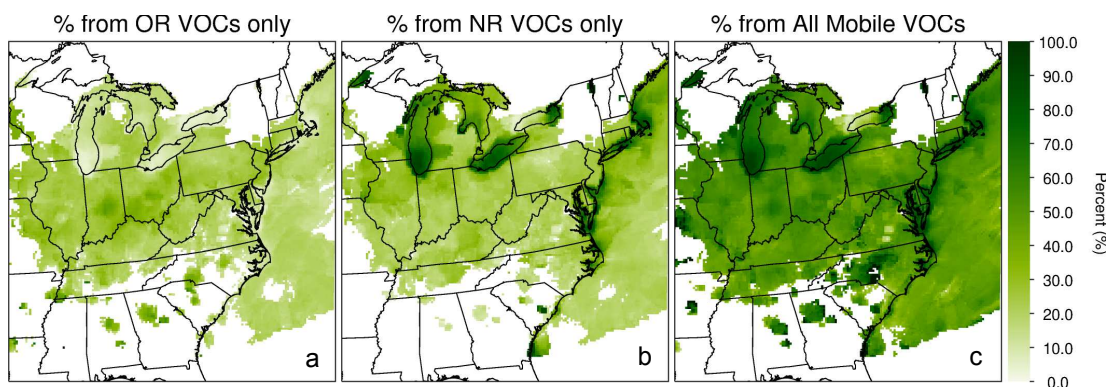


Figure 59: Percent of surface O_3 reductions in the c3 Science model framework attributed to the (a) on-road mobile sector only, (b) non-road mobile sector only (c) non-road and on-road mobile sectors combined, when compared to the simulation with 50% reduction of all anthropogenic VOCs.

It is evident from these simulations that VOC emission reductions from the non-road sector would greatly improve modeled surface O_3 in heavily populated coastal areas. Efforts aimed at ensuring emission inventories contain accurate data for this sector and are allocated correctly, along with further division of this broad modeling sector into more appropriate sources would help discern which specific non-road combustion engines are producing the problematic VOCs. A concerted effort to focus on this sector would greatly benefit air quality policy.

4.8: Conclusions

In this chapter we analyze observations from the RAMMPP aircraft campaign conducted on May 17 and 18, 2017 over the Long Island Sound (LIS), and compare these measurements to DISCOVER-AQ Baltimore-Washington observations. We also compare the RAMMPP data to forecast CMAQ model simulations conducted by NOAA for May 2017. Finally, we conduct model analysis using an in-house CMAQ model platform to discern the emissions sectors most responsible for O_3 formation in the LIS.

We find that atmospheric concentrations of anthropogenic VOCs dominate over isoprene in the NYC area and downwind over LIS throughout the May 17 and 18 O₃ exceedance event. Morning concentrations of VOCs are large compared to afternoon concentrations and may contribute to the high concentrations of O₃ measured later in the day. Propylene and isopentane (produced by the petrochemical industry and evaporative emissions) are shown to be the most important VOCs for O₃ production based on their concentrations and reactivity with OH. Efforts should be made to determine the specific sources of these problematic VOCs.

Comparison of DISCOVER-AQ isoprene, NO₂ and O₃ from the 2011 campaign in the Baltimore-Washington area show that significantly less isoprene and more NO₂ are present in the LIS region than in Maryland. It is possible that the atmosphere in the LIS may be VOC-sensitive at certain times and in certain locations. Elevated levels of O₃ and SO₂ above 2500 m over the LIS are shown to be due to long-range pollution transport. HYSPLIT back trajectories are performed and CEMS data are used to show active power plants along the trajectories. A Worldview image shows cumulus cloud formation along the HYSPLIT trajectories in areas of active power plants, indicating venting of the emissions due to atmospheric convection. Archived surface analysis and 500 mb maps show the presence of a high-pressure system off the east coast of the US during this time producing atmospheric flow consistent with the HYSPLIT trajectories.

CMAQ results from the NOAA National Air Quality Forecast Capability system for May 17 and 18 were compared to RAMMPP aircraft observations. Overall, the model produces concentrations of O₃ that support issuing an air quality alert, however it misses the full magnitude of the measured O₃ and often displaces the O₃

horizontally and vertically. Higher resolution modeling of the near-surface atmosphere may help to improve the distribution of O₃ within CMAQ. Comparison of measured and modeled CO indicate emissions of CO are too high by as much as a factor of 4 in the NYC/LIS region. Initial analysis indicates vertical mixing within the forecast CMAQ model framework is too slow.

Finally, model simulations conducted with two EPA emissions inventories indicate the non-road mobile sector contributes significantly to surface O₃ production in the LIS. The non-road sector consists of a variety of sources including recreational boats, construction vehicles, and railyard vehicles (not locomotives). Surface O₃ is significantly reduced when model simulations with 50% reductions in VOC emissions from the non-road mobile sector are performed. Areas with the largest reductions are over bodies of water like Lake Michigan (~32 ppb) and the Long Island Sound (~13 ppb). We suggest that air quality policies should consider controlling VOC emissions from member sources to the non-road sector. Inventory developers should consider providing model-ready emissions files that further categorize the non-road mobile inventory so model simulations can test attainment strategies that target specific sectors like recreational boating. A detailed, data-driven analysis of the various sub-categories that make up the non-road inventory is needed to verify the emission strength and attribution to this sector.

Chapter 5: Concluding Remarks

5.1: Conclusions from this Research

The work presented in this thesis provides the following contribution to the air quality community by answering the questions presented in Section 1.6:

1. Does the regulatory CMAQ modeling platform accurately represent the atmosphere? If not, what modifications can be made to improve this representation? How does this impact modeled ozone?

We found near-shore emissions from the largest shipping vessels (class 3 Commercial Marine Vessel (c3 Marine)) to be represented incorrectly in the regulatory CMAQ modeling platform. Modifications were performed to generate consistency between the off-shore and near-shore emissions files for the c3 Marine inventory sector. Comparison of baseline (unadjusted) and modified (c3 Adjust) simulations were performed. Model results show a decrease of Average Maximum 8-hr O₃ (AM8O₃) at Maryland AQS sites along the Chesapeake Bay where photochemical O₃ production is NO_x-sensitive (e.g. a decrease of ~6.5 ppb of O₃ at Essex, MD). This adjustment had the adverse effect of increasing AM8O₃ by ~3.5 ppb in the Long Island Sound (LIS). This result suggested local production of O₃ being within the VOC-sensitive regime. Vertical distribution of the near-shore c3 Marine emissions also improved model pollution dispersion and transport associated with these emissions; they are no longer remaining near the emission source, over-producing O₃ locally. Based on these changes, we conclude the CMAQ model represents the actual atmosphere more accurately.

Emissions for future year simulations are grown from a designated base year. In the 2018 modeling scenarios, AM8O₃ reductions are larger for the Chesapeake Bay (e.g. a decrease of ~13.6 ppb O₃ at Essex, MD). Surface O₃ in the LIS shows smaller increases than in the 2011 modeling, indicating the area may be transitioning to NO_x-sensitive conditions in the future. Analysis of a Maryland Department of the Environment (MDE) attainment strategy indicates legislation aimed at limiting O₃ precursors should be more effective in the adjusted model scenarios with the than in the baseline CMAQ simulation. This has important policy implications because incorrect emissions inventories for future years may obscure the true impacts of proposed attainment strategies.

We also examine accuracy of the NOAA forecast CMAQ model platform. We compare aircraft data to model output for the May 17 and 18, 2017 O₃ event over the LIS. Overall, the model reproduces an O₃ exceedance event, however it misses the full magnitude of the measured O₃. Additionally, this O₃ is often displaced from measured values. Limitations of model resolution may be to blame for this discrepancy.

2. Can we use CMAQ and relevant data sources to determine photochemical O₃ production regimes across the model domain?

We use column HCHO and NO₂ retrievals from the Ozone Monitoring Instrument (OMI) to calculate HCHO/NO₂ ratios across the CMAQ model domain. Baseline CMAQ ratios shows a larger area over the New York metropolitan area is within the VOC-sensitive to NO_x-sensitive transition region than is observed by the satellite. When empirically based model improvements are incorporated, known as the

c3 Science scenario (see Table 2), model performance improves compared to the satellite. Other metropolitan areas like Baltimore-Washington (MD), Chicago (IL), and Detroit (MI) are analyzed. Chicago and Detroit metro areas are within the transition region in the model framework; the satellite shows them as NO_x-sensitive. For the Baltimore-Washington region the model agrees well with the satellite; model and observations show the area to be NO_x-sensitive. In all cases, the proposed model adjustments improve performance in relation to the observed satellite HCHO/NO₂ ratio and more accurately represent tropospheric conditions. This comparison is relevant for *column* ratio calculations, not surface calculations for photochemical O₃ production regime. Care must be taken to keep from over-attributing column calculated ratios to surface conditions.

3. How does the photochemistry in the New York City pollution plume impact heavily populated areas downwind?

Analysis of aircraft observations for the May 2017 episode shows anthropogenic VOC concentrations are more important than isoprene over NYC and the LIS in the Planetary Boundary Layer. This goes against the conventional wisdom that isoprene, from biogenic sources, is the dominant VOC in the Eastern US. Air parcels decoupled from the surface biosphere will quickly lose isoprene ($\tau = \sim 1\text{-}2$ hours) relative to longer lived anthropogenic hydrocarbons.

Diurnal variation in VOC concentrations are large, suggesting morning concentrations may contribute to high concentrations of O₃ observed later in the day. Propylene and isopentane are shown to be the most important VOCs for O₃ production

in the LIS for these flights. Comparison of isoprene, NO₂ and O₃ between the RAMMPP and DISCOVER-AQ campaigns show significantly less isoprene and more NO₂ are present in the LIS than in Maryland. It is possible that at certain times and locations, atmospheric conditions may be VOC-sensitive in the LIS. Determination of the specific sources for the anthropogenic VOCs measured over the LIS, along with the temporal emissions rates would improve air quality attainment strategy development.

4. What are the important emissions sectors contributing to O₃ production in heavily populated, near-coastal, urban centers within the model domain?

Model simulations performed with the EPA Alpha2 and Beta2 emissions inventories indicate the non-road mobile sector contributes significantly to surface O₃ production in near-coastal urban areas. The non-road sector is a catch-all for fossil fuel burning sources including pleasure craft boats, recreational vehicles, lawn care equipment, and construction vehicles. Surface O₃ is significantly reduced when model simulations are performed with 50% reductions in VOC emissions from the non-road mobile sector. This reduction is not feasible for current attainment strategies, but it serves to highlight the importance of this inventory sector to modeled surface O₃ in areas with high modeled surface O₃ like the LIS (~13 ppb) and downwind of Chicago (~32 ppb) (see Figure 58e and b respectively).

Surface O₃ decreases of this magnitude are considerable compared to other attainment strategy modeling efforts performed by the University of Maryland MDE modeling group. Focus for attainment strategy development for the LIS has been given mostly to peaking units or auxiliary power units; electricity generation from small

sources used during high-demand days [McDill *et al.*, 2017]. This effort has provided evidence and motivation for development of attainment strategies focusing on the non-road sector. Emission strength and attribution from the member sources should be confirmed with data.

State Implementation Plans (SIPs) are complete outlines for how the State proposes to meet the required NAAQS standards. These include the full cost analysis of the proposed emissions reductions [EPA, 2018d]. Air directors for each state must therefore provide evidence that the proposed control measures are necessary and cost-effective. Justification for the proposed emissions reductions stems directly from modeling efforts, like those described in this thesis. Therefore, improving model representation of the actual atmosphere is of paramount importance, and providing direction for effective attainment strategies is critical for the success of future SIP development.

The main findings about air quality from this work can be applied worldwide, especially to coastal urban centers. Approximately 40% of the world's population lives within 60 miles of the coast [UN, 2017] and is impacted by coastal activity. Pollution from large urban centers and shipping ports therefore impacts significant populations worldwide. Mitigation of pollution in these areas is paramount to the health and well-being of human populations.

5.2: Needs for the Air Quality Community

5.2.1: Model and Emission Inventory Improvements

The modeling in this research is performed with the regulatory air quality model used by the state of Maryland to develop air quality attainment strategies. This CMAQ platform is for 12×12 km resolution over the Eastern half of the US. Based on lake/bay breeze studies discussed previously [Loughner *et al.*, 2014; Mazzuca *et al.*, 2017; Sillman *et al.*, 1993; Stauffer *et al.*, 2015], it is necessary to increase model resolution to resolve these small-scale events. Using a nested modeling domain over the targeted regions may improve the dynamical representation of these events. This is important for near-coastal areas as high O₃ events are often associated with this atmospheric flow [Loughner *et al.*, 2014; Mazzuca *et al.*, 2017; Sillman *et al.*, 1993; Stauffer *et al.*, 2015]. Incorporation of this nested, high-resolution modeling scheme into the commonly used regulatory modeling platform would improve attainment strategy development.

Source-apportionment model analysis would be improved by the division of catch-all inventory sectors like the non-road mobile and non-point inventories [McDill *et al.*, 2017]. The current non-road mobile inventory consists of a variety of combustion sources, some over water and some over land. Further separation of these sources would provide the ability to examine the contribution of each specific combustion source to surface O₃ production. Analysis of the pollution associated with marine pleasure craft will be useful in the LIS, Chesapeake Bay, and Great Lakes region.

Finally, more focus on particulate matter (PM_{2.5} and PM₁₀) modeling efforts are necessary. Particulate pollutants can be emitted directly into the atmosphere from sources like wildfires, smoke stacks, engine exhaust, and construction sites, or formed

in the atmosphere as secondary pollutants through reactions with, for example, SO₂ or NO_x compounds [EPA, 2018c]. Fine particulate matter is a source of human health issues including impaired lung function with inhaled into the lungs, and cardiovascular problems when particles (PM_{2.5}) cross the alveolar barrier into the blood stream [EPA, 2018c; Fann *et al.*, 2018; Pope *et al.*, 2002; Pope and Dockery, 2006]. Urban centers have large populations and emissions of PM associated with construction sites and combustion sources (cars, trucks, buses, shipping vessels, ferries, etc.) are numerous. City-scale modeling of these sources would improve understanding of the full extent of common exposure levels to PM. Larger scale modeling would highlight upwind sources of PM and the transport of PM throughout the domain. Air quality policy modeling efforts have mostly been focused on O₃; it is time to broaden the scope to include PM.

5.2.2: Field Campaigns and Measurements

This body of work has focused on the air quality in large coastal population centers where people are exposed to poor air quality episodes specifically related to O₃. Data for these areas, especially the New York City/Long Island Sound (LIS) region are scarce, relying heavily on available ground-based monitoring networks. A robust field campaign like the previous DISCOVER-AQ campaigns in Baltimore-Washington (2011), Houston (2013), California (2013), and Colorado (2014) would enhance the understanding of pollution in this region.

As shown earlier in Figure 4, the NJ/NYC/CT area currently has the worst air quality on the East Coast. A field mission dedicated to discerning the specific causes for this poor air quality is necessary. Measurement of air in the NYC/NJ urban centers,

air upwind, and air downwind would help to characterize emission flux for the region. Examination of the boundary layer dynamical processes over the LIS would provide valuable information about the circulation, impacts of the nocturnal inversions over the region, and the features of the small-scale on-shore atmospheric flow (“sound breeze”) that contribute to high measured O₃ levels along the CT coast.

Exploration of the various sources of O₃ precursors in the NJ/NYC metropolitan area, and sources over the LIS, would improve analysis of model results. Specific attention to the emissions sources contained within the non-road mobile inventory would assist with development of effective State Implementation Plans for the region. Additional examination of emissions from other over-water sources such as those classified in the class 1 and class 2 commercial marine vessels (ferries, fishing boats, etc.) [McDill *et al.*, 2017] would provide data for inventory comparison.

Finally, measurements of important VOCs like isoprene and propylene would significantly expand the community’s understanding of the role VOCs have in producing O₃ over the LIS. Quantifying the isoprene flux for the region would provide data for inventory comparison and assist with estimating the biogenic VOC contribution to the area. As we have shown in the analysis in Chapter 4, isoprene is not the dominant VOC (at least during this sampling period) in the region, therefore we cannot assume the effects of anthropogenic VOCs are minimal compared to biogenic VOCs. Measurements of the most important VOCs to O₃ production (Based on the MIR/POCP scales) would allow policy makers to target the most important VOCs, and design effective attainment strategies.

Findings from this proposed field campaign in the NYC/LIS region would be applicable in other heavily populated coastal areas like Chicago and the Great Lakes region. Modeling analyses performed in this body of work have shown that large concentrations of O₃ are produced over near-coastal regions. Comparison of the model results to data gathered during the field mission described above would assist with determining if the models are correct, and if so, why this is the case.

5.3: Halogens and Ozone: Future Research

CMAQ often produces extremely high values of modeled AM8O₃ (> 140ppb) for some locations, especially along large bodies of water (Figure 21a and b from Chapter 3). These high values are much greater than measured surface O₃ in 2011 and are unrealistic. A possible source of this model/measurement discrepancy is the simplified representation of the chemistry within the chemical mechanism used by CMAQ. The CAMx model can use a variety of chemical mechanisms including the most recently released version of the Carbon Bond mechanism, CB6r4 [Emery *et al.*, 2016; Yarwood *et al.*, 2005]. Halogens such as iodine, chlorine, and bromine are known to impact O₃ production [Chang *et al.*, 2002]. However, halogen chemistry is limited in the chemical mechanisms used for regulatory air quality modeling.

One important halogen mechanism is the heterogeneous reaction of N₂O₅ on an aerosol particle to form either HNO₃ or ClNO₂ [Riedel *et al.*, 2013; Riedel *et al.*, 2014; Sarwar *et al.*, 2014; Simon *et al.*, 2009; Thornton *et al.*, 2010]. This is reaction with direct influence on the production of O₃ because HNO₃ is considered a reservoir species for NO₂, leading to its loss from the atmosphere. In mechanisms with no halogen

chemistry, N_2O_5 is assumed to always produce HNO_3 , which accounts for approximately 50% of loss of NO_x from the atmosphere [Riedel *et al.*, 2012]. Appropriate reaction partitioning, including the formation of ClNO_2 from the heterogeneous reaction of N_2O_5 on a chlorine containing aerosol particle would reduce the loss of NO_x from the atmosphere via N_2O_5 . This is because ClNO_2 photolyzes just after sunrise to release NO_2 and Cl back into the atmosphere, enhancing the impact of NO_2 on the production of O_3 [Riedel *et al.*, 2012; Riedel *et al.*, 2013; Riedel *et al.*, 2014; Simon *et al.*, 2009; Thornton *et al.*, 2010].

In addition to the research focusing on Cl and ClNO_2 , studies have shown that bromine contributes to O_3 removal in the free troposphere [Schmidt *et al.*, 2016] and in the arctic surface regions [Salawitch *et al.*, 2010]. Concentrations of bromine compounds increase significantly with altitude, reflecting the longer lifetimes of these species and emphasizing their impact in the free troposphere [Sherwen *et al.*, 2016]. Analysis of aircraft and satellite observations during the NASA ARCTAS (Arctic Research of the Composition of the Troposphere from Aircraft and Satellites) mission combined with 3-D model simulations indicate the presence of significant amount of BrO in the stratosphere [Choi *et al.*, 2012; Salawitch *et al.*, 2010]. Studies such as these have helped improve the representation of tropospheric BrO in models.

In a study using the GEOS-Chem global model, Sherwen *et al.* [2016] found that inclusion of halogen chemistry accounts for ~20% of the total loss of O_3 . Incorporating various halogen related O_3 production and loss processes into the regulatory air quality modeling using currently studies as guidance would improve ozone representation within the model, especially for near-coastal areas.

5.3.1: Implementation of Halogen Chemistry in Air Quality Models

A previous study by *Sarwar et al.* [2012] used a continental US CMAQ domain and added halogen chemistry to the CB05 mechanism to analyze the impact of ClNO_2 from N_2O_5 on air quality. They showed that monthly mean O_3 increases throughout the eastern US when heterogeneous production of ClNO_2 is included [*Sarwar et al.*, 2012]. This group also assessed the impact of heterogeneous production of ClNO_2 for the Northern Hemisphere [*Sarwar et al.*, 2014]. They found significant levels of ClNO_2 in winter over highly populated areas like Europe and China, contributing to increased O_3 levels. Inclusion of iodine chemistry and examination of ozone deposition velocities over near-coastal salt water will better represent modeled surface O_3 in air quality models [*Gantt et al.*, 2017; *Wong and Cheng*, 2008].

To accurately model halogens, emission inventories must be updated and/or developed. Point source emissions of HCl , Cl_2 and other chlorine sources are shown to be potentially large and are currently not well characterized in emissions inventories [*Crisp et al.*, 2014; *Lawler et al.*, 2011]. Concentrations of HCl tend to be highest in polluted coastal regions [*Crisp et al.*, 2014; *Keene et al.*, 2007], like the Chesapeake Bay and Long Island Sound because of elevated sea salt concentrations. Additionally, air polluted with HCl is advected towards the East Coast from source regions like the Midwest [*Keene et al.*, 2007]. HCl is the dominant form of chlorine in the troposphere, and besides direct emission from coal combustion, it is also produced via reaction of Cl with VOCs [*Hossaini et al.*, 2016]. Studies suggest that halogen emissions can be approximated based on the EPA's Toxic Release Inventory (TRI), and the most recent National Emissions Inventory (NEI), both developed by the US EPA. For context,

important sources of atmospheric Cl are industrial point sources, cooling towers, swimming pools, and sea salt [*Chang et al.*, 2002; *Riedel et al.*, 2013; *Spicer et al.*, 1998; *Tanaka et al.*, 2000].

Evaluation of air quality models with data is necessary to ensure model chemistry and emissions inventories are within reason. Various monitoring networks measure halogen species including the Clean Air Status and Trends Network (CASTNET), National Atmospheric Deposition Program (NADP), and observations from the Interagency Monitoring of Protected Visual Environments (IMPROVE) network. Observational campaigns conducted throughout the US like SENEX 2013 (<https://www.esrl.noaa.gov/csd/projects/senex/>) focused mostly over the southeast portion of the US, and the WINTER 2015 campaign centered mostly over the Mid-Atlantic US (<https://www.esrl.noaa.gov/csd/groups/csd7/measurements/2015winter/>) provide the opportunity to gain a better understanding of halogen chemistry under different geographic, seasonal and meteorological conditions.

Based on previous studies, implementation of sophisticated halogen chemistry in regulatory air quality models should reduce surface O₃ over large salt water bodies (iodine chemistry) and within the free troposphere, but a more localized increase in O₃ downwind from power plants and in some near-shore polluted regions due to chlorine chemistry is expected. Model results should be compared to available data to examine this hypothesis.

Appendix A

Table 5: Measured VOCs from RAMMPP aircraft campaign, corresponding rate constants for reaction with OH, and their Maximum Incremental Reactivity value.

VOCs	Rate Reaction (k_{OH}) ^a	MIR ^b
Benzene	$k_{OH} = (2.3 \times 10^{-12}) \times e^{(-190/T)}$	0.72
Cyclohexane	$k_{OH} = (2.88 \times 10^{-17}) \times T^2 \times e^{(-309/T)}$	1.25
Cyclopentane	$k_{OH} = (2.67 \times 10^{-17}) \times e^{(-509/T)}$	2.39
Ethylbenzene	$k_{OH} = (7.00 \times 10^{-12})$	3.04
Isobutane	$k_{OH} = (5.4 \times 10^{-12}) \times e^{(-285/T)}$	1.23
Isopentane	$k_{OH} = (3.70 \times 10^{-12})$	1.45
Isoprene	$k_{OH} = (2.70 \times 10^{-11}) \times e^{(390/T)}$	10.61
Methylcyclohexane	$k_{OH} = (9.29 \times 10^{-12})$	1.70
Methylcyclopentane	$k_{OH} = (7.66 \times 10^{-12})$	2.19
Propane	$k_{OH} = (7.60 \times 10^{-12}) \times e^{(-585/T)}$	0.49
Propylene	$k_{OH} = (3.0 \times 10^{-11}) \times e^{(300/T)}$	11.66
Styrene	$k_{OH} = (5.80 \times 10^{-11})$	1.73
Toluene	$k_{OH} = (1.8 \times 10^{-12}) \times e^{(340/T)}$	4.00
cis-2-Butene	$k_{OH} = (1.10 \times 10^{-11}) \times e^{(487/T)}$	14.24
cis-2-Pentene	$k_{OH} = (6.54 \times 10^{-11})$	10.38
m-Ethyltoluene	$k_{OH} = (1.86 \times 10^{-11})$	7.39
o-Ethyltoluene	$k_{OH} = (1.19 \times 10^{-11})$	5.59
p-Ethyltoluene	$k_{OH} = (1.18 \times 10^{-11})$	4.44
mp-Xylene	$k_{OH} = (1.65 \times 10^{-11})$	7.80
o-Xylene	$k_{OH} = (1.36 \times 10^{-11})$	7.64
n-Butane	$k_{OH} = (9.8 \times 10^{-12}) \times e^{(-425/T)}$	1.15
n-Decane	$k_{OH} = (3.13 \times 10^{-17}) \times T^2 \times e^{(416/T)}$	0.68
n-Dodecane	$k_{OH} = (1.39 \times 10^{-11})$	0.55
n-Heptane	$k_{OH} = (1.59 \times 10^{-17}) \times T^2 \times e^{(478/T)}$	1.07

n-Hexane	$k_{OH} = (1.53 \times 10^{-17}) \times T^2 \times e^{(414/T)}$	1.24
n-Nonane	$k_{OH} = (2.51 \times 10^{-17}) \times T^2 \times e^{(447/T)}$	0.78
n-Octane	$k_{OH} = (2.76 \times 10^{-17}) \times T^2 \times e^{(378/T)}$	0.90
n-Pentane	$k_{OH} = (2.44 \times 10^{-17}) \times T^2 \times e^{(183/T)}$	1.31
n-Propylbenzene	$k_{OH} = (5.80 \times 10^{-12})$	2.03
n-Undecane	$k_{OH} = (1.29 \times 10^{-11})$	0.61
m-Diethylbenzene	$k_{OH} = (2.31 \times 10^{-11})$	7.10
p-Diethylbenzene	$k_{OH} = (1.43 \times 10^{-11})$	4.43
trans-2-Butene	$k_{OH} = (1.01 \times 10^{-11}) \times e^{(550/T)}$	15.16
trans-2-Pentene	$k_{OH} = (6.69 \times 10^{-11})$	10.56
123-Trimethylbenzene	$k_{OH} = (3.27 \times 10^{-11})$	11.97
124-Trimethylbenzene	$k_{OH} = (3.25 \times 10^{-11})$	8.87
135-Trimethylbenzene	$k_{OH} = (5.67 \times 10^{-11})$	11.76
1-Butene	$k_{OH} = (6.60 \times 10^{-12}) \times e^{(465/T)}$	9.73
1-Hexene	$k_{OH} = (3.70 \times 10^{-11})$	5.49
1-Pentene	$k_{OH} = (5.86 \times 10^{-12}) \times e^{(500/T)}$	7.21
224-Trimethylpentane	$k_{OH} = (2.09 \times 10^{-12}) \times e^{(140/T)}$	1.26
22-Dimethylbutane	$k_{OH} = (3.22 \times 10^{-11}) \times e^{(-781/T)}$	1.17
23-Dimethylbutane	$k_{OH} = (1.24 \times 10^{-17}) \times T^2 \times e^{(494/T)}$	0.97
22-Dimethylpentane	$k_{OH} = (5.20 \times 10^{-12})$	1.12
23-Dimethylpentane	$k_{OH} = (1.95 \times 10^{-11}) \times e^{(-330/T)}$	1.34
2-Methylhexane	$k_{OH} = (6.86 \times 10^{-12})$	1.19
2-Methylpentane	$k_{OH} = (5.30 \times 10^{-12})$	1.50
3-Methylpentane	$k_{OH} = (5.40 \times 10^{-12})$	1.80
3-Methylheptane	$k_{OH} = (6.30 \times 10^{-12})$	1.24
3-Methylhexane	$k_{OH} = (7.15 \times 10^{-12})$	1.61

^a Reaction rates are from MCMv3.3.1 (<http://mcm.leeds.ac.uk/MCM/>)

^b Maximum Incremental Reactivity (MIR) values are from [Carter, 2010b]

Appendix B

Table 6: The species used in the core Carbon Bond 05 (CB05) chemical mechanism. Species for the 16 added chemical reactions, the secondary organic aerosol formation, and chlorine mechanisms in the UMD CMAQ installation (CB05TUCI) are indicated with an *. (Core CB05 species are in Table 2-1 from *Yarwood et al.* [2005].)

Species Name	Description
NO	Nitric oxide
NO2	Nitrogen dioxide
O3	Ozone
O	Oxygen atom in the O ³ (P) electronic state
O1D	Oxygen atom in the O ¹ (D) electronic state
OH	Hydroxyl radical
HO2	Hydroperoxy radical
H2O2	Hydrogen peroxide
NO3	Nitrate radical
N2O5	Dinitrogen pentoxide
HONO	Nitrous acid
HNO3	Nitric acid
PNA	Peroxynitric acid (HNO ₄)
CO	Carbon monoxide
FORM	Formaldehyde
ALD2	Acetaldehyde
C2O3	Acetylperoxy radical
PAN	Peroxyacetyl nitrate
ALDX	Propionaldehyde and higher aldehydes
CXO3	C3 and higher acylperoxy radicals
PANX	C3 and higher peroxyacetyl nitrates
XO2	NO to NO2 conversion from alkylperoxy (RO ₂) radical
XO2N	NO to organic nitrate conversion from alkylperoxy (RO ₂) radical
NTR	Organic nitrate (RNO ₃)
ETOH	Ethanol
CH4	Methane
MEO2	Methylperoxy radical
MEOH	Methanol
MEPX	Methylhydroperoxide
FACD	Formic acid
ETHA	Ethane
ROOH	Higher organic peroxide

AACD	Acetic and higher carboxylic acids
PACD	Peroxyacetic and higher peroxycarboxylic acids
PAR	Paraffin carbon bond (C-C)
ROR	Secondary alkoxy radical
ETH	Ethene
OLE	Terminal olefin carbon bond (R-C=C)
IOLE	Internal olefin carbon bond (R-C=C-R)
ISOP	Isoprene
ISPD	Isoprene product (lumped methacrolein, methyl vinyl ketone, etc.)
TERP	Terpene
TOL	Toluene and other monoalkyl aromatics
XYL	Xylene and other polyalkyl aromatics
CRES	Cresols and higher molecular weight phenols
TO2	Toluene-hydroxyl radical adduct
OPEN	Aromatic ring opening product
CRO	Methylphenoxy radical
MGLY	Methylglyoxal and other aromatic products
SO2	Sulfur dioxide
SULF	Sulfuric acid (gaseous)
*BENZ	Benzene
*CAO2	Peroxy radical from aromatic degradation products
*CAT1	Methyl-catechols
*CRN2	Peroxy radical from nitro-cresol
*CRNO	Alkoxy radical from nitro-cresols
*CRON	Nitro-cresols
*CRPX	Nitro-cresol hydroperoxides
*HCO3	Adduct from HO2 + formaldehyde
*OPAN	Peroxyacyl nitrate (PAN compound) from OPO3
*OPO3	Peroxyacyl radical from OPEN
*SESQ	Sesquiterpene
*SULRXN	Counter species for aerosol from SO2
*TOLRXN	Counter species for aerosol from TOL
*XYLRXN	Counter species for aerosol from XYL
*ISOPRXN	Counter species for aerosol from ISOP
*TRPRXN	Counter species for aerosol from TERP
*TOLRO2	First generation products from TOL that react with NO and NO2
*XYLRO2	First generation products from XYL that react with NO and NO2
*BENZRO2	First generation products from BENZ that react with NO and NO2

*TOLNRXN	Counter species for aerosol from TOL (from reaction with NO)
*TOLHRXN	Counter species for aerosol from TOL (from reaction with HO2)
*XYLNRXN	Counter species for aerosol from XYL (from reaction with NO)
*XYLHRXN	Counter species for aerosol from XYL (from reaction with HO2)
*BNZNRXN	Counter species for aerosol from BENZ (from reaction with NO)
*BNXHRXN	Counter species for aerosol from BENZ (from reaction with HO2)
*SESQRXN	Counter species for aerosol from SESQ

Chlorine Mechanism Species

*CL2	Molecular chlorine
*HOCL	Hypochlorous acid
*CL	Chlorine atom
*CLO	Chlorine monoxide
*FMCL	Formyl chloride (HC(O)Cl)

Constants

H2	Molecular hydrogen fixed constant (0.56)
N2	Molecular nitrogen fixed constant (0.7808E+06)
O2	Molecular oxygen fixed constant (0.2095E+06)
AIR	Fixed constant (1.00E+06)
CH4	Methane fixed constant (1.85)

Table 7: Reactions used in the UMD CMAQ installation (CB05TUCI) that consist of the core CB05 chemical mechanism (156 reactions) with updates from the mechanism definitions file mech.def (16 additional reactions, 10 secondary organic aerosol formation reactions, 23 chlorine reactions). Any reactions with modifications from the core CB05 mechanism are indicated with an *. Interpretation of the rate expression is described below the table. (Core CB05 rates in Table 2-2 from *Yarwood et al.* [2005]).

Rxn	Reactants	Products	Rate Expression (k)
R1	NO2	NO + O	$1.00 \times j_{\text{NO2}}$ (SAPRC 99)
R2	O + O2 + M	O3 + M	$6.00\text{E-}34^{-2.4}$
R3	O3 + NO	NO2	$3.00\text{E-}12 @ 1500$
R4	O + NO2	NO	$5.60\text{E-}12 @ -180$
R5	O + NO2	NO3	$2.50\text{E-}31^{-1.8} \&$ $2.20\text{E-}11^{-0.7}$
R6	O + NO	NO2	$9.00\text{E-}32^{-1.5} \&$ $3.00\text{E-}11$
R7	NO2 + O3	NO3	$1.20\text{E-}13 @ 2450$
R8	O3	O	$1.00 \times j_{\text{O3(O3P)}}$ (IUPAC 05)
R9	O3	O1D	$1.00 \times j_{\text{O3(O1D)}}$ (IUPAC 05)
R10	O1D + M	O + M	$2.10\text{E-}11 @ -102$
R11	O1D + H2O	2*OH	$2.20\text{E-}10$
R12	O3 + OH	HO2	$1.70\text{E-}12 @ 940$
R13	O3 + HO2	OH	$1.00\text{E-}14 @ 490$
R14	NO3	NO2 + O	$1.00 \times j_{\text{NO3(NO2)}}$ (SAPRC99)
R15	NO3	NO	$1.00 \times j_{\text{NO3(NO)}}$ (SAPRC99)
R16	NO3 + NO	2*NO2	$1.50\text{E-}11 @ -170$
R17	NO3 + NO2	NO + NO2	$4.50\text{E-}14 @ 1260$
R18	NO3 + NO2	N2O5	$2.00\text{E-}30^{-4.4} \&$ $1.40\text{E-}12^{-0.7}$
*R19	N2O5 + H2O	2*HNO3	$1.00\text{E-}22$ (IUPAC 10)
*R20	N2O5 + H2O + H2O	2*HNO3	0.00 (IUPAC 10)
R21	N2O5	NO3 + NO2	$1.0\text{E-}03^{-3.5} @ 11000 \&$ $9.70\text{E}14^{0.1} @ 11080 \&$ $0.45 \& 1.0$
R22	NO + NO + O2	2*NO2	$3.30\text{E-}39 @ -530$
R23	NO + NO2 + H2O	2*HONO	$5.00\text{E-}40$

R24	NO + OH	HONO	7.00E-31 ^{-2.6} & 3.60E-11 ^{-0.1}
R25	HONO	NO + OH	1.00 x j _{HONO} (IUPAC 05)
R26	OH + HONO	NO2	1.80E-11 @ 390
R27	HONO + HONO	NO + NO2	1.00E-20
R28	NO2 + OH	HNO3	2.00E-30 ^{-3.0} & 2.50E-11
R29	OH + HNO3	NO3	%2 2.4E-14 @ -460 & 2.70E-17 @ -2199 & 6.50E-34 @ -1335
R30	HO2 + NO	OH + NO2	3.50E-12 @ -250
R31	HO2 + NO2	PNA	1.80E-31 ^{-3.2} & 4.70E-12 & 0.6
R32	PNA	HO2 + NO2	4.10E-5 @ 10650 & 4.80E15 @ 11170 & 0.6
R33	OH + PNA	NO2	1.30E-12 @ -380
R34	HO2 + HO2	H2O2	%3 2.3E-13 @ -600 & 1.70E-33 @ -1000
R35	HO2 + HO2 + H2O	H2O2	%3 3.22E-34 @ -2800 & 2.38E-54 @ -3200
R36	H2O2	2*OH	1.00 x j _{H2O2} (SAPRC99)
R37	OH + H2O2	HO2	2.90E-12 @ 160
R38	O1D + H2	OH + HO2	1.10E-10
R39	OH + H2	HO2	5.50E-12 @ 2000
R40	OH + O	HO2	2.20E-11 @ -120
R41	OH + OH	O	4.20E-12 @ 240
R42	OH + OH	H2O2	6.90E-31 ^{-1.0} & 2.60E-11 ⁰
R43	OH + HO2		4.80E-11 @ -250
R44	HO2 + O	OH	3.00E-11 @ -200
R45	H2O2 + O	OH + HO2	1.40E-12 @ 2000
R46	NO3 + O	NO2	1.00E-11
R47	NO3 + OH	HO2 + NO2	2.20E-11
R48	NO3 + HO2	HNO3	3.50E-12
R49	NO3 + O3	NO2	1.00E-17
R50	NO3 + NO3	2*NO2	8.50E-13 @ 2450
R51	PNA	0.610*HO2 + 0.610*NO2 + 0.390*OH + 0.390*NO3	1.00 x j _{HO2(NO2)} (IUPAC 05)
R52	HNO3	OH + NO2	1.00 x j _{HNO3} (IUPAC 05)

R53	N2O5	NO2 + NO3	1.00 x j _{N2O5} (IUPAC05)
R54	XO2 + NO	NO2	2.60E-12 @ -365
R55	XO2N + NO	NTR	2.60E-12 @ -365
R56	XO2 + HO2	ROOH	7.50E-13 @ -700
R57	XO2N + HO2	ROOH	7.50E-13 @ -700
R58	XO2 + XO2		6.80E-14
R59	XO2N + XO2N		6.80E-14
R60	XO2 + XO2N		6.80E-14
R61	NTR + OH	HNO3 + HO2 + 0.330*FORM + 0.330*ALD2 + 0.330*ALDX - 0.660*PAR	5.90E-13 @ 360
R62	NTR	NO2 + HO2 + 0.330*FORM + 0.330*ALD2 + 0.330*ALDX - 0.660*PAR	1.00 x j _{NTR} (IUPAC 05 isopropyl nitrate)
*R62	NTR	*Adjusted binned values for the absorption cross section to be more representative of NTR species	*NTR adjustment in c3 Science
R63	ROOH + OH	XO2 + 0.500*ALD2 + 0.500*ALDX	3.01E-12 @ -190
R64	ROOH	OH + HO2 + 0.500*ALD2 + 0.500*ALDX	1.00 x j _{COOH} (SAPRC99)
R65	OH + CO	HO2	%3 1.44E-13 @ 0.0 & 3.43E-33 @ 0.0
R66	OH + CH4	MEO2	2.45E-12 @ 1775
R67	MEO2 + NO	FORM + HO2 + NO2	2.80E-12 @ -300
R68	MEO2 + HO2	MEPX	4.10E-13 @ -750
R69	MEO2 + MEO2	1.370*FORM + 0.740*HO2 + 0.630*MEOH	9.50E-14 @ -390
R70	MEPX + OH	0.700*MEO2 + 0.300*XO2 + 0.300*HO2	3.80E-12 @ -200
R71	MEPX	FORM + HO2 + OH	1.00 x j _{COOH} (SAPRC99)
R72	MEOH + OH	FORM + HO2	7.30E-12 @ 620
R73	FORM + OH	HO2 + CO	9.00E-12
*R74	FORM	2*HO2 + CO	1.00 x j _{HCHO}

			(SAPRC99)
R75	FORM	CO	1.00 x j _{HCHO} (SAPRC99)
R76	FORM + O	OH + HO2 + CO	3.40E-11 @ 1600
R77	FORM + NO3	HNO3 + HO2 + CO	5.80E-16
R78	FORM + HO2	HCO3	9.70E-15 @ -625
R79	HCO3	FORM + HO2	2.40E+12 @ 7000
R80	HCO3 + NO	FACD + NO2 + HO2	5.60E-12
R81	HCO3 + HO2	MEPX	5.60E-15 @ -2300
R82	FACD + OH	HO2	4.00E-13
R83	ALD2 + O	C2O3 + OH	1.80E-11 @ 1100
R84	ALD2 + OH	C2O3	5.60E-12 @ -270
R85	ALD2 + NO3	C2O3 + HNO3	1.40E-12 @ 1900
R86	ALD2	MEO2 + CO + HO2	1.00 x j _{CCHO(R)} (SAPRC99)
R87	C2O3 + NO	MEO2 + NO2	8.10E-12 @ -270
R88	C2O3 + NO2	PAN	2.70E-28 ^{-7.1} & 1.20E-11 ^{-0.9} & 0.3
R89	PAN	C2O3 + NO2	4.90E-3 @ 12100 & 5.40E16 @ 13830 & 0.3
R90	PAN	C2O3 + NO2	1.00 x j _{PAN} (IUPAC 05)
R91	C2O3 + HO2	0.800*PACD + 0.200*AACD + 0.200*O3	4.30E-13 @ -1040
R92	C2O3 + MEO2	0.900*MEO2 + 0.900*HO2 + FORM + 0.100*AACD	2.00E-12 @ -500
R93	C2O3+XO2	0.900*MEO2 + 0.100*AACD	4.40E-13 @ -1070
R94	C2O3 + C2O3	2*MEO2	2.90E-12 @ -500
R95	PACD + OH	C2O3	4.00E-13 @ -200
*R96	PACD	MEO2 + OH	1.00 x j _{PACD} (CB05 peroxyacetic acid)
R97	AACD + OH	MEO2	4.00E-13 @ -200
R98	ALDX + O	CXO3 + OH	1.30E-11 @ 870
R99	ALDX + OH	CXO3	5.10E-12 @ -405

R100	ALDX + NO ₃	CXO ₃ + HNO ₃	6.50E-15
R101	ALDX	MEO ₂ + CO + HO ₂	1.00 x j _{C₂CHO} (SAPRC99)
R102	CXO ₃ + NO	ALD ₂ + NO ₂ + HO ₂ + XO ₂	6.70E-12 @ -340
R103	CXO ₃ + NO ₂	PANX	2.70E-28 ^{-7.1} & 1.20E-11 ^{-0.9} & 0.3
R104	PANX	CXO ₃ + NO ₂	4.90E-3 @ 12100 & 5.40E16 @ 13830 & 0.3
R105	PANX	CXO ₃ + NO ₂	1.00 x j _{PAN} (IUPAC 05)
R106	PANX + OH	ALD ₂ + NO ₂	3.00E-13
R107	CXO ₃ + HO ₂	0.800*PACD + 0.200*AACD + 0.200*O ₃	4.30E-13 @ -1040
R108	CXO ₃ + MEO ₂	0.900*ALD ₂ + 0.900*XO ₂ + HO ₂ + 0.100*AACD + 0.100*FORM	2.00E-12 @ -500
R109	CXO ₃ + XO ₂	0.900*ALD ₂ + 0.100*AACD	4.40E-13 @ -1070
R110	CXO ₃ + CXO ₃	2*ALD ₂ + 2*XO ₂ + 2*HO ₂	2.90E-12 @ -500
R111	CXO ₃ + C ₂ O ₃	MEO ₂ + XO ₂ + HO ₂ + ALD ₂	2.90E-12 @ -500
R112	PAR + OH	0.870*XO ₂ + 0.130*XO ₂ N + 0.110*HO ₂ + 0.060*ALD ₂ – 0.110*PAR + 0.760*ROR + 0.050*ALDX	8.10E-13
R113	ROR	0.960*XO ₂ + 0.600*ALD ₂ + 0.940*HO ₂ – 2.100*PAR + 0.040*XO ₂ N + 0.020*ROR + 0.500*ALDX	1.00E+15 @ 8000
R114	ROR	HO ₂	1.60E+03
R115	ROR + NO ₂	NTR	1.50E-11
R116	O + OLE	0.200*ALD ₂ + 0.300*ALDX + 0.300*HO ₂ + 0.200*XO ₂ + 0.200*CO + 0.200*FORM + 0.010*XO ₂ N + 0.200*PAR + 0.100*OH	1.00E-11 @ 280
R117	OH + OLE	0.800*FORM + 0.330*ALD ₂ + 0.620*ALDX + 0.800*XO ₂ + 0.950*HO ₂ – 0.700*PAR	3.20E-11
R118	O ₃ + OLE	0.180*ALD ₂ + 0.740*FORM + 0.320*ALDX + 0.220*XO ₂	6.50E-15 @ 1900

		+ 0.100*OH + 0.330*CO + 0.440*HO2 – PAR	
R119	NO3 + OLE	NO2 + FORM + 0.910*XO2 + 0.090*XO2N + 0.560*ALDX + 0.350*ALD2 – PAR	7.00E-13 @ 2160
R120	O + ETH	FORM + 1.700*HO2 + CO + 0.700*XO2 + 0.300*OH	1.04E-11 @ 792
R121	OH + ETH	XO2 + 1.560*FORM + 0.220*ALDX + HO2	1.00E-28^-0.8 & 8.80E-12
R122	O3 + ETH	FORM + 0.630*CO + 0.130*HO2 + 0.130*OH + 0.370*FACD	1.20E-14 @ 2630
R123	NO3 + ETH	NO2 + XO2 + 2.0*FORM	3.30E-12 @ 2880
R124	IOLE + O	1.240*ALD2 + 0.660*ALDX + 0.100*HO2 + 0.100*XO2 + 0.100*CO + 0.100*PAR	2.30E-11
R125	IOLE + OH	1.300*ALD2 + 0.700*ALDX + HO2 + XO2	1.00E-11 @ -550
R126	IOLE + O3	0.650*ALD2 + 0.350*ALDX + 0.250*FORM + 0.250*CO + 0.500*O + 0.500*OH + 0.500*HO2	8.40E-15 @ 1100
R127	IOLE + NO3	1.180*ALD2 + 0.640*ALDX + HO2 + NO2	9.6E-13 @ 270
*R128	TOL + OH	0.280*HO2 + 0.100*XO2 + 0.180*CRES + 0.650*TO2 + 0.072*OH + 1.0*TOLRO2	1.80E-12 @ -355
*R129	TO2 + NO	0.860*NO2 + 1.20*HO2 + 0.860*OPEN + 0.140*NTR + 0.52*MGLY + 0.336*FORM + 0.336*CO	2.70E-12 @ -360
*R130	TO2 + HO2		1.90E-13 @ -1300
*R131	OH + CRES	0.060*CRO + 0.120*XO2 + 1.120*HO2 + 0.130*OPEN + 0.732*CAT1 + 0.060*CO + 0.060*XO2N + 0.060*FORM	1.70E-12 @ -950
*R132	CRES + NO3	0.300*CRO + HNO3 + 0.600*XO2 + 0.360*HO2 + 0.480*ALDX + 0.240*FORM + 0.240*MGLY + 0.120*OPEN + 0.100*XO2N + 0.240*CO	1.40E-11
*R133	CRO + NO2	CRON	2.10E-12
R134	CRO + HO2	CRES	5.50E-12

*R135	CRON + OH	CRNO	1.53E-12
*R136	CRON + NO3	CRNO + HNO3	3.80E-12
*R137	CRNO + NO2	2.0*NTR	2.10E-12
*R138	CRNO + O3	CRN2	2.86E-13
*R139	CRN2 + NO	CRNO + NO2	2.54E-12 @ -360
*R140	CRN2 + HO2	CRPX	2.40E-13 @ -1300
*R141	CRPX	CRNO + OH	0.01 x j _{NO2} (SAPRC99)
*R142	CRPX + OH	CRN2	1.90E-12 @ -190
*R143	OPEN	OPO3 + HO2 + CO	0.04 x j _{NO2} (SAPRC99)
*R144	OPEN + OH	0.600*OPO3 + 0.400*CAO2	4.40E-11
*R145	OPEN + O3	0.030*ALDX + 0.620*OPO3 + 0.700*FORM + 0.030*XO2 + 0.690*CO + 0.080*OH + 0.760*HO2 + 0.200*MGLY	5.40E-17 @ 500
*R146	OPEN + NO3	OPO3 + HNO3	3.80E-12
*R147	CAT1 + OH	CAO2	7.00E-11
*R148	CAT1 + NO3	CRO + HNO3	1.70E-10
*R149	CAO2 + NO	0.860*NO2 + 1.200*HO2 + 0.344*FORM + 0.344*CO + 0.140*NTR	2.54E-12 @ -360
*R150	CAO2 + HO2		2.40E-13 @ -1300
*R151	OPO3 + NO	NO2 + XO2 + HO2 + ALDX	1.10E-11
*R152	OPO3 + NO2	OPAN	1.10E-11
*R153	OPAN	OPO3 + NO2	1.00E-04
*R154	OH + XYL	0.700*HO2 + 0.500*XO2 + 0.200*CRES + 0.800*MGLY + 1.100*PAR + 0.300*TO2 + 1.0*XYLRO2	1.7E-11 @ -116
*R155	OH + MGLY	XO2 + C2O3	1.80E-11
R156	MGLY	C2O3 + HO2 + CO	1.00 x j _{MGLY} (IUPAC 05)

R157	O + ISOP	0.750*ISPD + 0.500*FORM + 0.250*XO2 + 0.250*HO2 + 0.250*CXO3 + 0.250*PAR	3.60E-11
R158	OH + ISOP	0.912*ISPD + 0.629*FORM + 0.991*XO2 + 0.912*HO2 + 0.088*XO2N	2.54E-11 @ -407.6
R159	O3 + ISOP	0.650*ISPD + 0.600*FORM + 0.200*XO2 + 0.066*HO2 + 0.266*OH + 0.200*CXO3 + 0.150*ALDX + 0.350*PAR + 0.066*CO	7.86E-15 @ 1912
R160	NO3 + ISOP	0.200*ISPD + 0.800*NTR + XO2 + 0.800*HO2 + 0.200*NO2 + 0.800*ALDX + 2.40*PAR	3.03E-12 @ 448
R161	OH + ISPD	1.565*PAR + 0.167*FORM + 0.713*XO2 + 0.503*HO2 + 0.334*CO + 0.168*MGLY + 0.252*ALD2 + 0.210*C2O3 + 0.250*CXO3 + 0.120*ALDX	3.36E-11
R162	O3 + ISPD	0.114*C2O3 + 0.150*FORM + 0.850*MGLY + 0.154*HO2 + 0.268*OH + 0.064*XO2 + 0.020*ALD2 + 0.360*PAR + 0.225*CO	7.10E-18
R163	NO3 + ISPD	0.357*ALDX + 0.282*FORM + 1.282*PAR + 0.925*HO2 + 0.643*CO + 0.850*NTR + 0.075*CXO3 + 0.075*XO2 + 0.150*HNO3	1.00E-15
R164	ISPD	0.333*CO + 0.067*ALD2 + 0.900*FORM + 0.832*PAR + 1.033*HO2 + 0.700*XO2 + 0.967*C2O3	0.0036 x j _{ACROLEIN} (SAPRC99)
*R165	TERP + O	0.150*ALDX + 5.12*PAR + TRPRXN	3.60E-11
*R166	TERP + OH	0.750*HO2 + 1.250*XO2 + 0.250*XO2N + 0.280*FORM + 1.66*PAR + 0.470*ALDX + TRPRXN	1.50E-11 @ -449
*R167	TERP + O3	0.570*OH + 0.070*HO2 + 0.760*XO2 + 0.180*XO2N + 0.240*FORM + 0.001*CO + 7.000*PAR + 0.210*ALDX + 0.390*CXO3 + TRPRXN	1.20E-15 @ 821

*R168	TERP + NO3	0.470*NO2 + 0.280*HO2 + 1.030*XO2 + 0.250*XO2N + 0.470*ALDX + 0.530*NTR + TRPRXN	3.70E-12 @ -175
*R169	SO2 + OH	SULF + HO2 + SULRXN	3.30E-31 ^{-4.3} & 1.60E-12 ^{0.0}
R170	OH + ETOH	HO2 + 0.900*ALD2 + 0.050*ALDX + 0.100*FORM + 0.100*XO2	6.90E-12 @ 230
R171	OH + ETHA	0.991*ALD2 + 0.991*XO2 + 0.009*XO2N + HO2	8.70E-12 @ 1070
R172	NO2 + ISOP	0.200*ISPD + 0.800*NTR + XO2 + 0.800*HO2 + 0.200*NO + 0.800*ALDX + 2.400*PAR	1.50E-19

Chlorine Mechanism

CL1	CL2	2.00*CL	1.00 x j _{Cl} (IUPAC 05)
CL2	HOCL	OH + CL	1.00 x j _{HOCl} (IUPAC 05)
CL3	CL + O3	CLO	2.30E-11 @ 200
CL4	CLO + CLO	0.300*CL2 + 1.400*CL	1.63E-14
CL5	CLO + NO	CL + NO2	6.40E-12 @ -290
CL6	CLO + HO2	HOCL	2.70E-12 @ -220
CL7	OH + FMCL	CL + CO	5.00E-13
CL8	FMCL	CL + CO + HO2	1.00 x j _{FMCL} (IUPAC 05)
CL9	CL + CH4	HCL + MEO2	6.60E-12 @ 1240
CL10	CL + PAR	HCL + 0.870*XO2 + 0.130*XO2N + 0.110*HO2 + 0.060*ALD2 + 0.110*PAR + 0.760*ROR + 0.050*ALDX	5.00E-11
CL11	CL + ETHA	HCL + 0.991*ALD2 + 0.991*XO2 + 0.009*XO2N + HO2	8.30E-11 @ 100
CL12	CL + ETH	FMCL + 2.0-*XO2 + 1.00*HO2 + 1.00*FORM	1.07E-11
CL13	CL + OLE	FMCL + 0.330*ALD2 + 0.670*ALDX + 2.00*XO2 + 1.00*HO2 – 1.00*PAR	2.50E-10
CL14	CL + IOLE	0.300*HCL + 0.700*FMCL + 0.450*ALD2 + 0.550*ALDX + 0.300*OLE + 0.300*PAR + 1.70*XO2 + 1.00*HO2	3.50E-10

CL15	CL + ISOP	0.150*HCL + 1.00*HO2 + 1.00*HO2 + 0.850*FMCL + 1.00*ISPD	4.30E-10
CL16	CL + FORM	HCL + 1.00*HO2 + 1.00*CO	8.20E-11 @ 34
CL17	CL + ALD2	HCL + 1.00*C2O3	7.90E-11
CL18	CL + ALDX	HCL + 1.00*CXO3	1.30E-10
CL19	CL + MEOH	HCL + 1.00*HO2 + 1.00*FORM	5.50E-11
CL20	CL + ETOH	HCL + 1.00*HO2 + 1.00*ALD2	8.20E-11 @ -45
CL21	HCL + OH	CL	6.58E-13 ^{1.16} @ -58
CL22	CL + TOL	HCL + 0.880*XO2 + 0.880*HO2 + 0.12*XO2N	6.10E-11
CL23	CL + XYL	HCL + 0.840*XO2 + 0.840*HO2 + 0.160*XO2N	1.20E-10
Secondary Aerosol Mechanism			
SA1	TOLRO2 + NO	NO + TOLRXN	2.70E-12 @ -360
SA2	TOLRO2 + HO2	HO2 + TOLHRXN	1.90E-13 @ -1300
SA3	XYLRO2 + NO	NO + XYLNRXN	2.70E-12 @ -360
SA4	XYLRO2 + HO2	HO2 + XYLRXN	1.90E-13 @ -1300
SA5	BENZ + OH	OH + 1.00*BENZRO2	2.47E-12 @ 206
SA6	BENZRO2 + NO	NO + BNZNRXN	2.70E-12 @ -360
SA7	BENZRO2 + HO2	HO2 + BNZHRXN	1.90E-13 @ -1300
SA8	SESQ + O3	O3 + SESQRXN	1.16E-14
SA9	SESQ + OH	OH + SESQRXN	1.97E-10
SA10	SESQ + NO3	NO3 + SESQRXN	1.90E-11

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